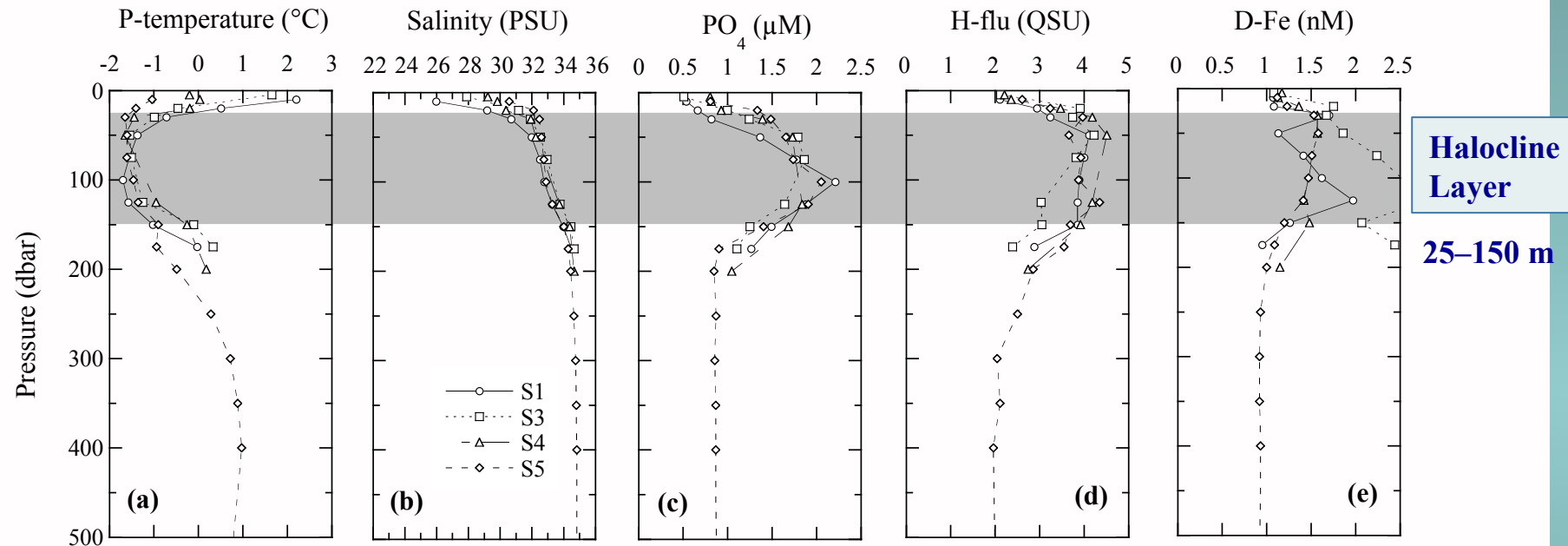


Vertical distributions of T, S, PO₄, humic F-intensity, and [D-Fe] in the surface water (0–500 m) of the slope and shelf regions

Fig. 4



****Subsurface maxima of nutrient, humic F-intensity and [D-Fe] in Upper HL**
Three main processes to form the subsurface maxima of chemical components in Upper HL

(1) Brine rejection during sea ice formation

Formation of dense shelf water with low-temperature and high-salinity

(2) Interactions between brine waters and sediments in the shallower shelves

Supply of nutrient, humic-type FDOM and D-Fe from shelf sediments to the overlying water

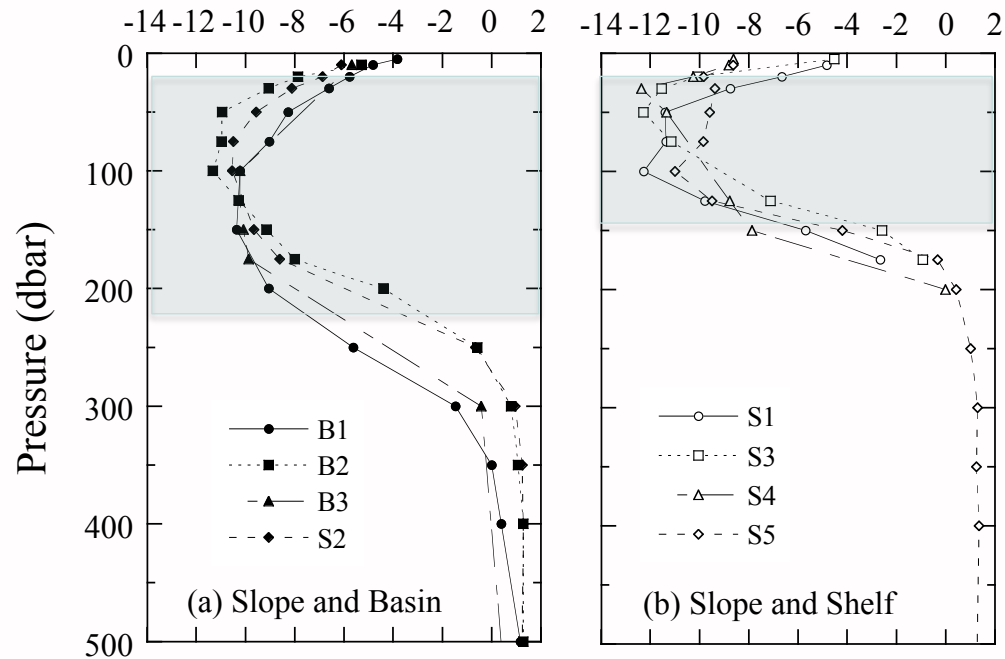
(3) Inflow of the Atlantic water with low nutrient, humic-type FDOM and D-Fe into Lower HL

Decrease in nutrient, humic-type FDOM and D-Fe with depth in Lower HL

$$N^* = ([NO_3^-] - 16[PO_4^{3-}] + 2.9) \times 0.87 \quad [\text{Gruber and Sarmiento, 1997}]$$

$$N^* = ([NO_3^-] + [NO_2^-] + [NH_4^+] - 16[PO_4^{3-}] + 2.9) \times 0.87 \quad [\text{Yoshikawa et al., 2006}]$$

$$N^* = ([NO_3^-] + [NO_2^-] + [NH_4^+] - 16[PO_4^{3-}] + 2.9) \times 0.87, \mu\text{M}$$



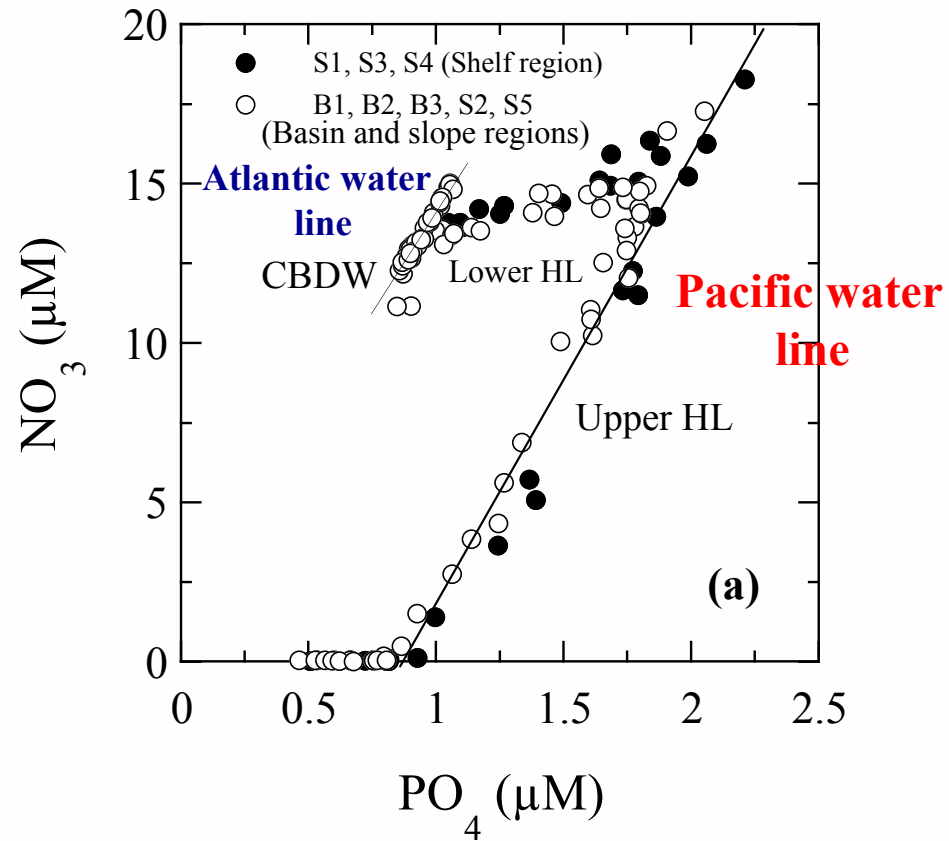
Upper Halocline Layer

$$[NH_4^+] = 0 \sim 3 \mu\text{M} \quad [\text{Gueguen et al., 2007; Wang et al., 2006}]$$

Large negative N* value in Upper HL:

Denitrification, consuming nitrate instead of oxygen for bacterial respiration in low-oxygen pore waters in sediments, and supply of nutrient, humic-type FDOM and D-Fe from shelf sediments to the overlying water in the shallower continental shelves

NO₃ versus PO₄ in the Chukchi Sea and Canada Basin in this study (2008)



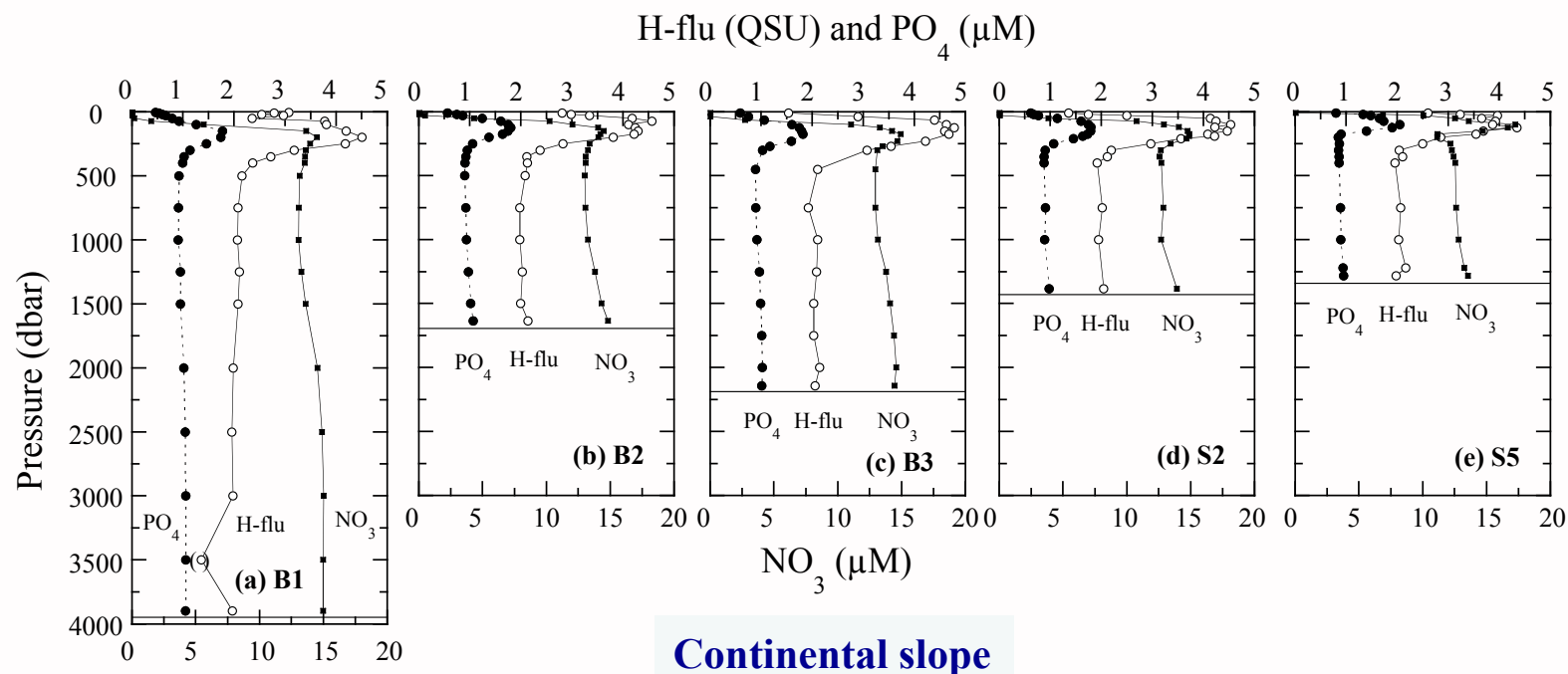
Upper HL: Pacific water via Bering Strait

Lower HL: mixture between Pacific and Atlantic waters due to the inflow of high-salinity Atlantic water into Lower HL

(Jones *et al.*, 1998; Yamamoto-Kawai *et al.*, 2008)

Vertical distributions of PO_4 , NO_3 and humic-type F-intensity in the water column of the slope and basin regions

Fig. 5



Upper HL: Subsurface maxima of nutrient ($\text{NO}_3=15 \mu\text{M}$, $\text{PO}_4=1.8 \mu\text{M}$, $\text{SiO}_2=30 \mu\text{M}$)

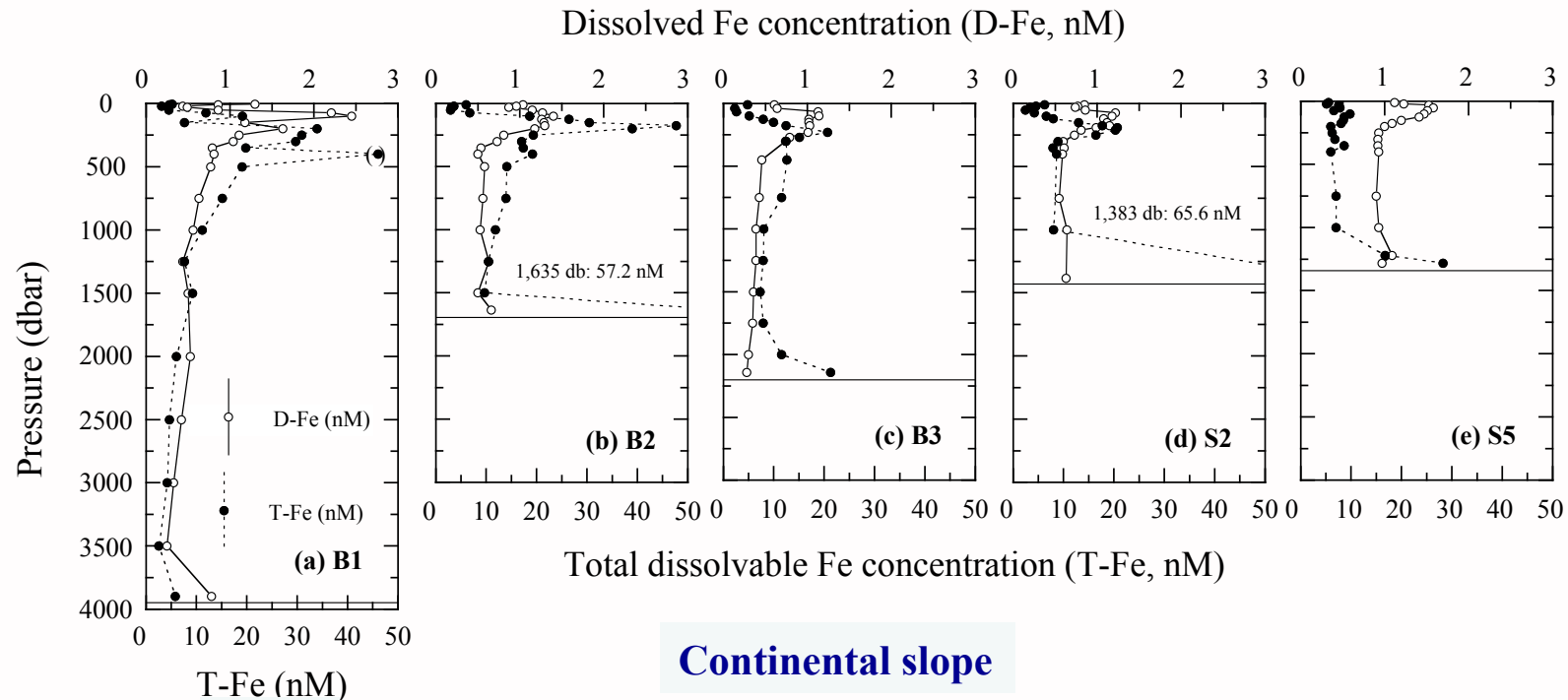
Subsurface maxima of Humic-type F-intensity (4~5 QSU): higher than those (2~2.5 QSU) in the North Pacific Ocean

Deep water (Atlantic layer + CBDW): $\text{NO}_3=12\text{--}15 \mu\text{M}$, $\text{PO}_4=0.8\text{--}1.1 \mu\text{M}$, $\text{SiO}_2=5\text{--}15 \mu\text{M}$

Humic-type F-intensity=2 QSU

Vertical distributions of D-Fe and T-Fe concentrations in the water column of the slope and basin regions

Fig. 8



- * **Upper HL: Subsurface maxima of D-Fe** (D-Fe=1.2–2.5 nM)
 Supply of D-Fe from shelf sediments to the overlying water in the shallower shelves and lateral Fe transport from the shelves to the basin
 Organic Fe complexation with dissolved humic substances, probably controlling D-Fe concentrations in seawater and laterally transporting Fe to the Arctic Ocean interior
- * **Lower HL: Subsurface maxima of T-Fe** (T-Fe=10–50 nM)
 Sediment resuspension on the shelves by the inflow of the Atlantic water into Lower HL