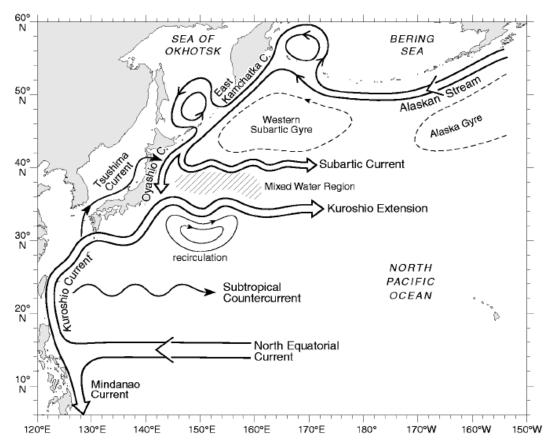
Decadal Variability, Impact, and Prediction of the Kuroshio Extension System

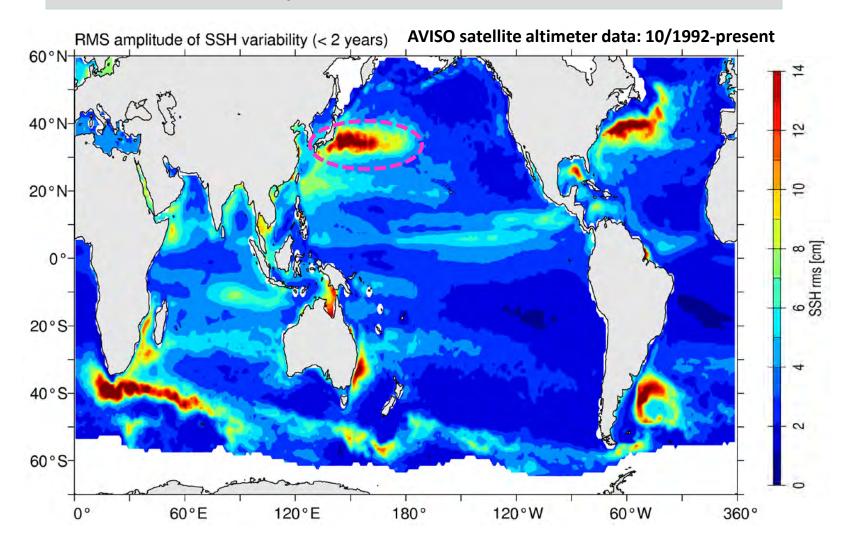
- B. Qiu¹, S. Chen¹, N. Schneider¹ and B. Taguchi²
 - 1. Dept of Oceanography, University of Hawaii, USA
 - 2. Earth Simulator Center, JAMSTEC, Japan

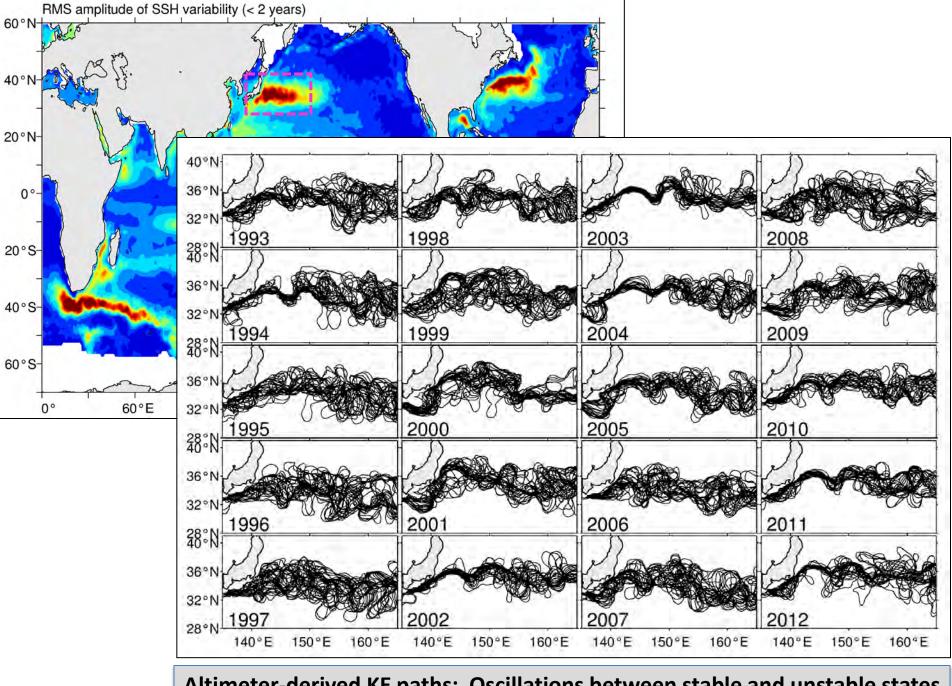


2014 FUTURE Open Science Meeting, Kohala Coast, HI, 14-18 April 2014

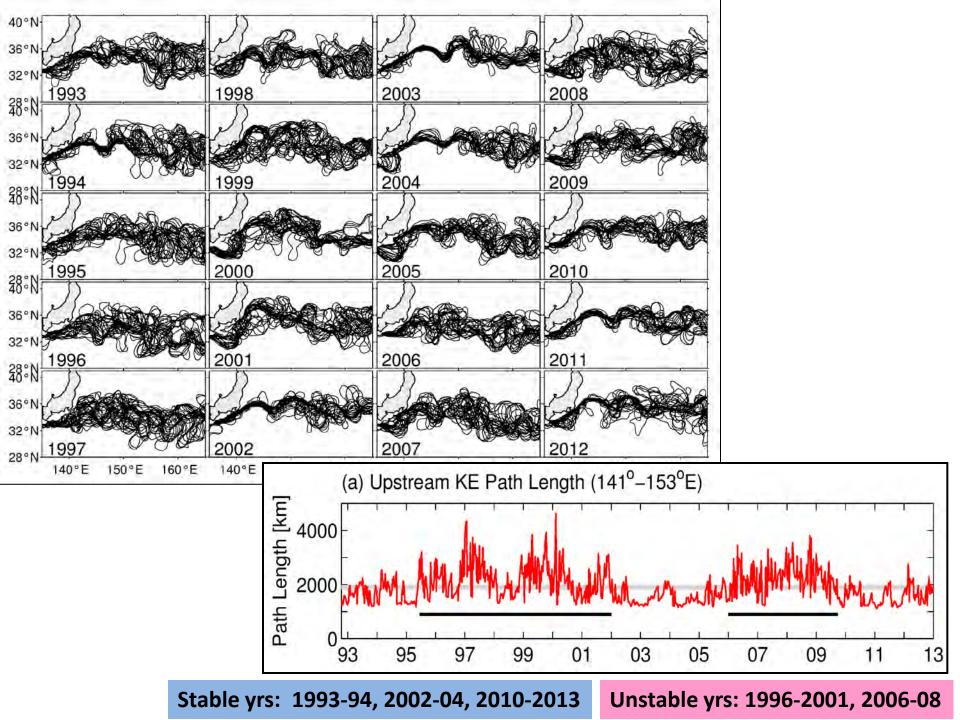
Outlines

- Observed decadal changes in the Kuroshio Extension system
- Impact on SST front and overlying stormtracks
- KE index and its prediction

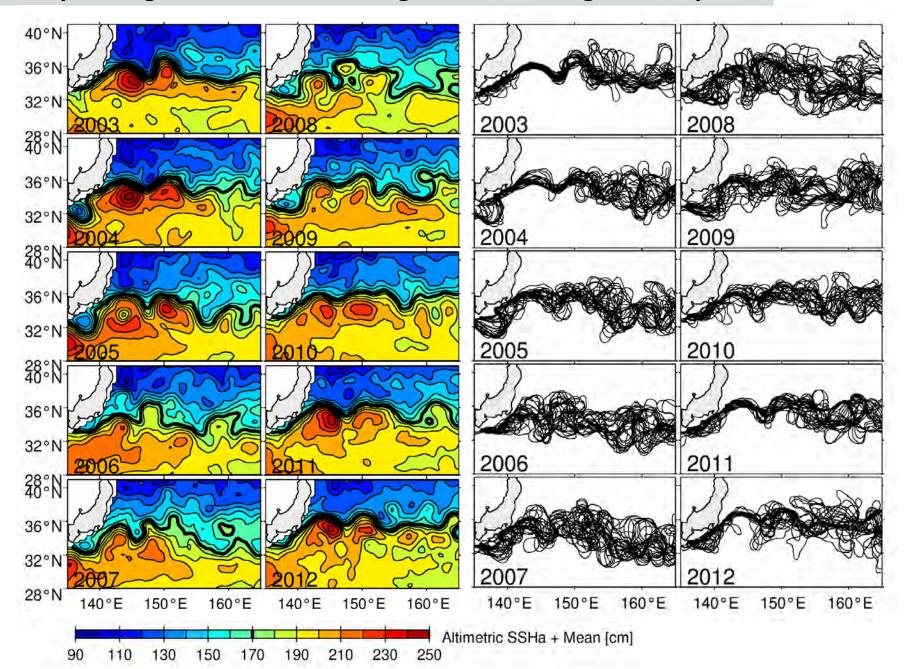




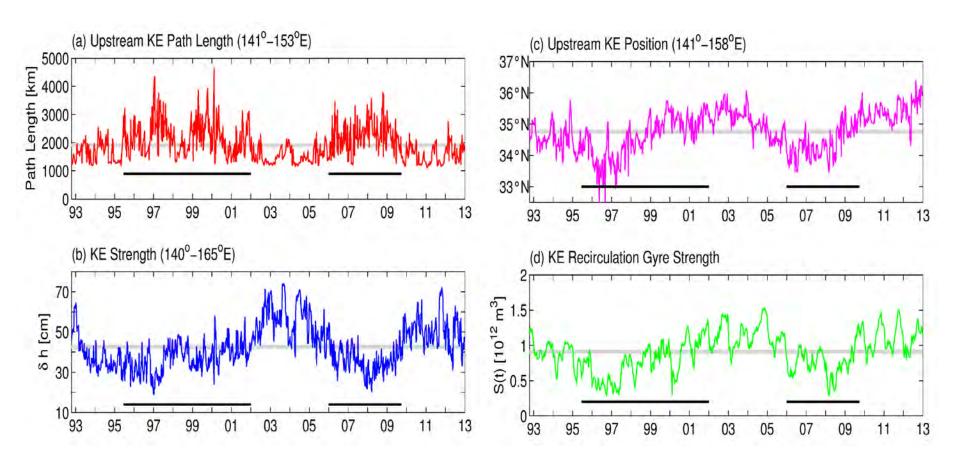
Altimeter-derived KE paths: Oscillations between stable and unstable states



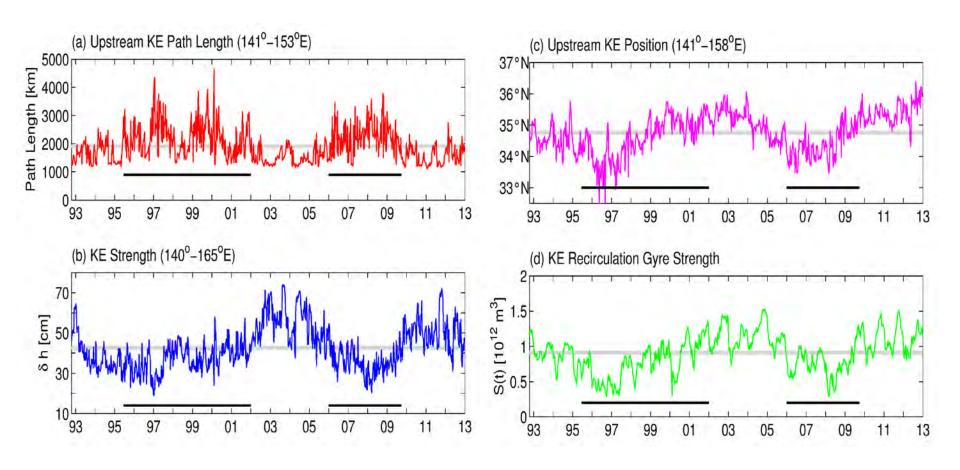
Yearly-averaged SSH field in the region surrounding the KE system



Various dynamic properties representing the decadal KE variability

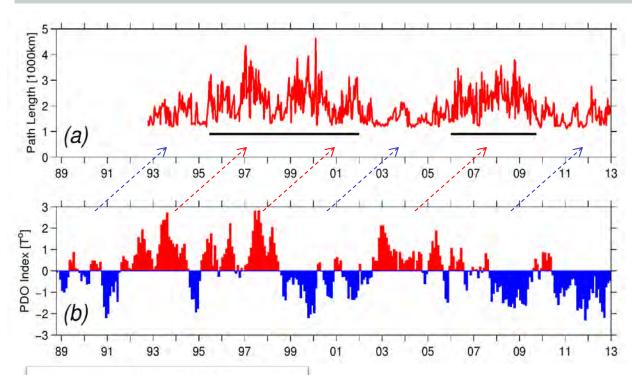


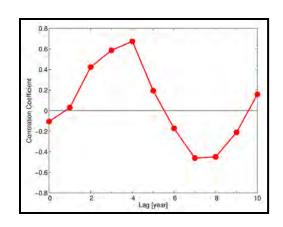
Various dynamic properties representing the decadal KE variability

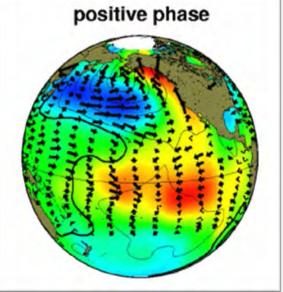


Q1: What causes the transitions between the stable and unstable dynamic states of the KE system?

Decadal KE variability lags the PDO index by 4 yrs (r=0.67)



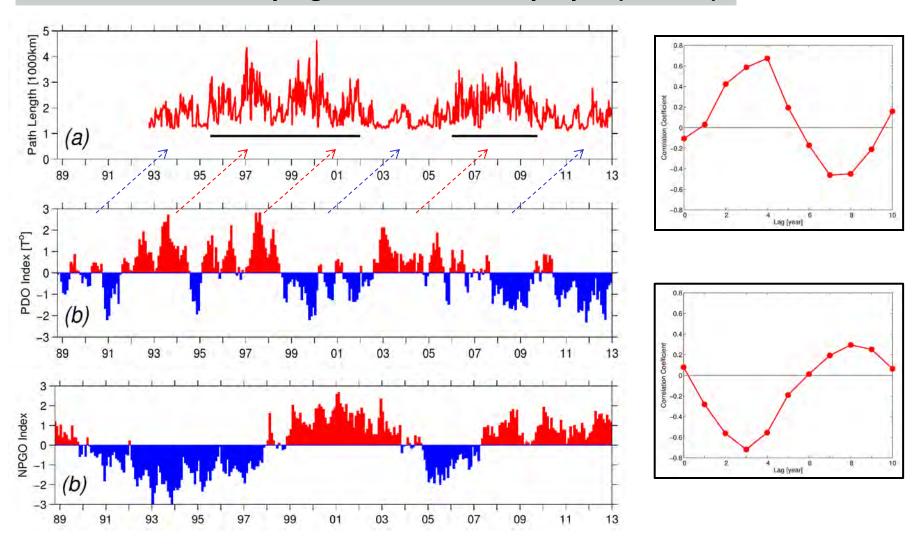




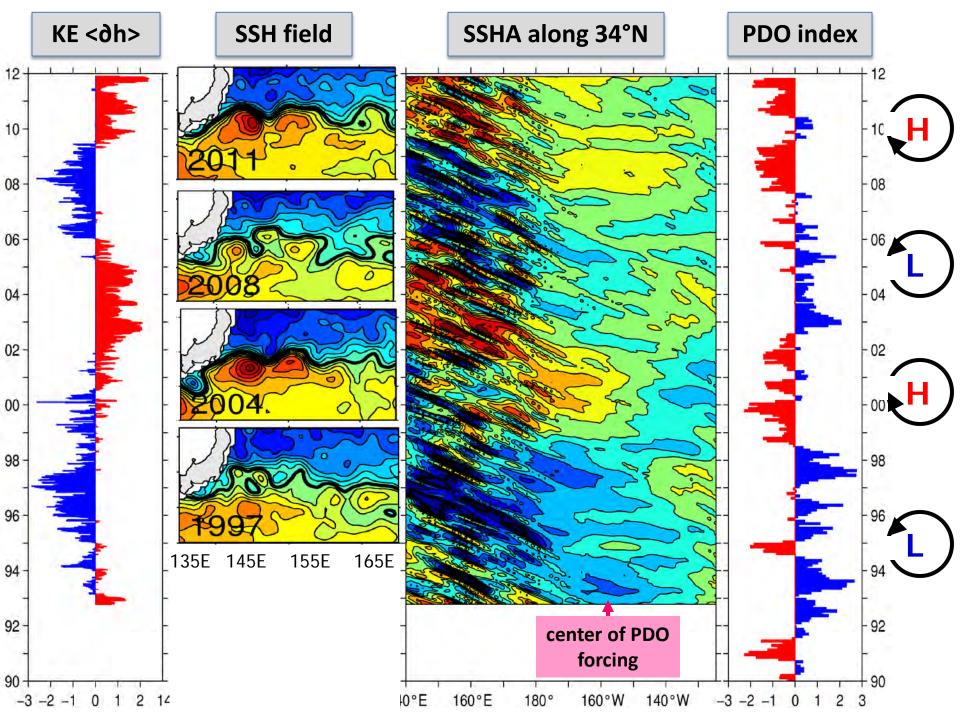
- Center of PDO wind forcing is in the <u>eastern</u> half of North Pacific basin
- + PDO generates negative local SSHAs through Ekman divergence, and vice versa

Mantua et al. (1997, BAMS)

Decadal KE variability lags the PDO index by 4 yrs (r=0.67)



Di Lorenzo et al. (2008, GRL)



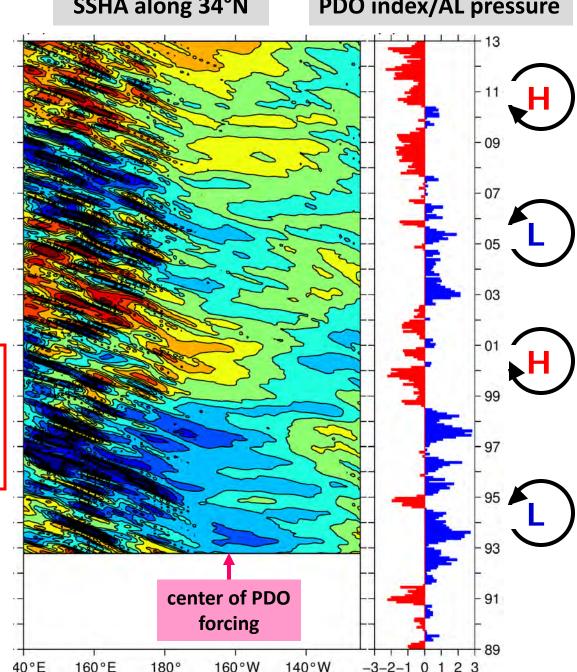
 Large-scale SSH changes are governing by linear vorticity dynamics:

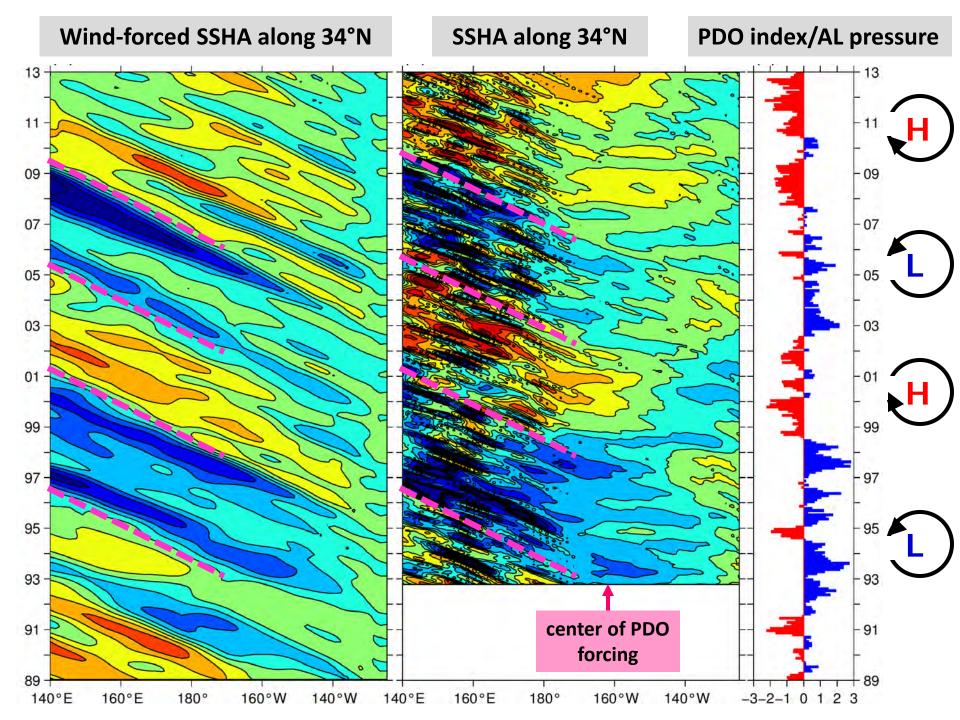
$$rac{\partial h}{\partial t} - c_R rac{\partial h}{\partial x} = -rac{g' \,
abla imes oldsymbol{ au}}{
ho_o g f}$$

 Given the wind forcing, SSH changes can be found along the Rossby wave paths:

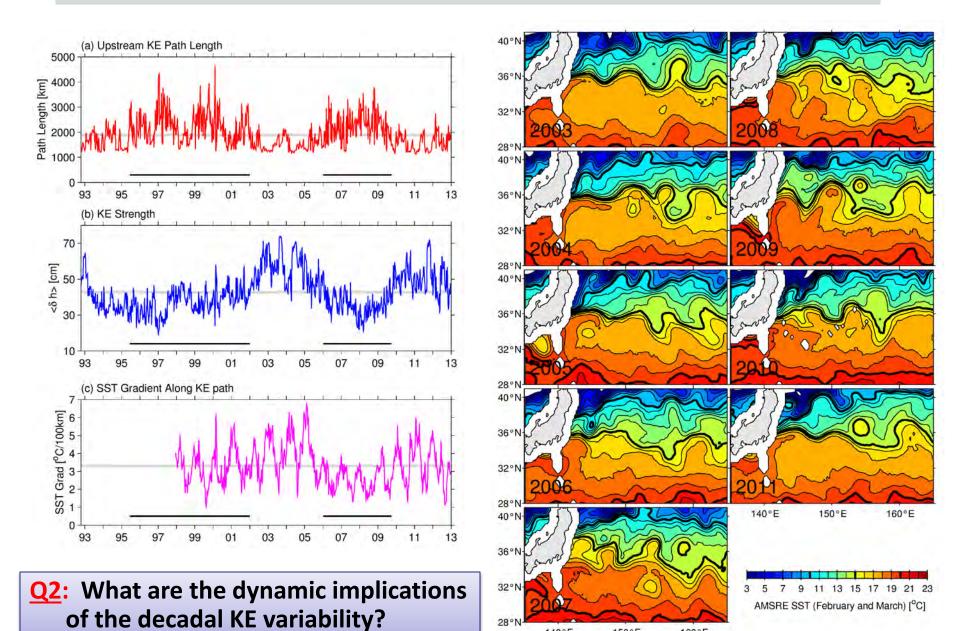
$$egin{aligned} h(x,t) &= h\left(x_e, t + rac{x-x_e}{c_R}
ight) + \ -rac{g'}{
ho_o g f c_R} \int_{x_e}^x
abla imes oldsymbol{ au}\left(x', t + rac{x-x'}{c_R}
ight) dx' \end{aligned}$$

cf. Qiu and Chen (2005); Taguchi et al. (2007); Ceballos et al. (2009); **Sasaki et al. (2013)**





KE dynamic state affects regional SST and cross-front SST gradient

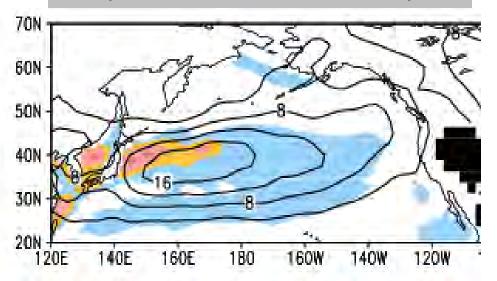


140°E

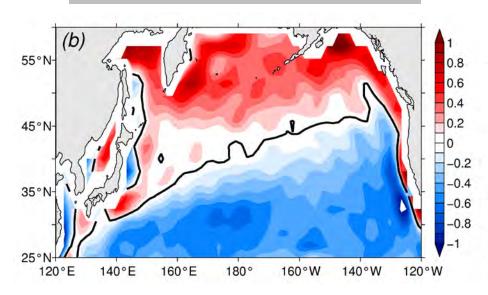
150°E

160°E

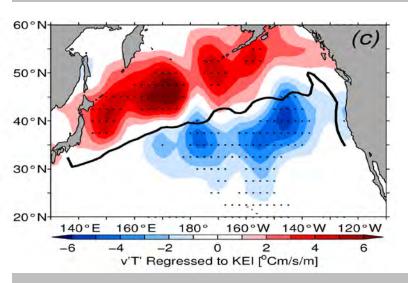
JF rms 850mb v'T' of stormtrack variability (contours; Nakamura et al. 2004)



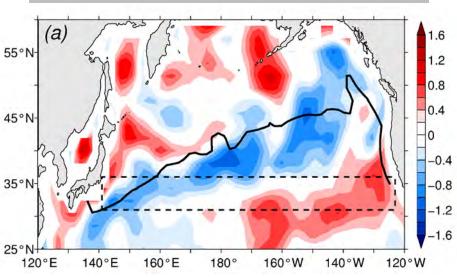
NCEP mean wind stress curl field



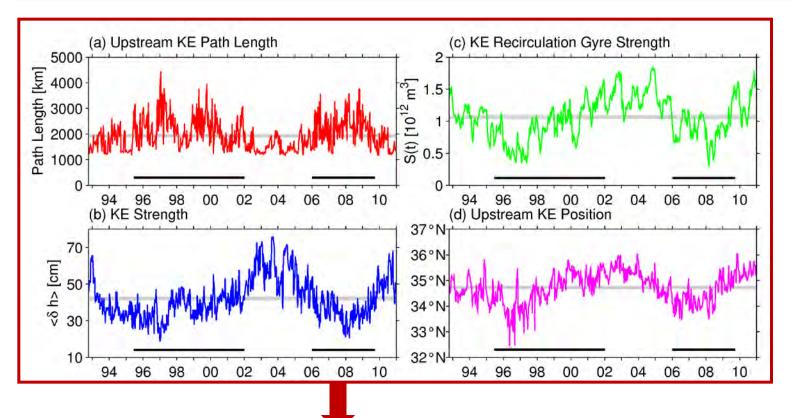
850mb v'T' and Q' regressed to the KE dynamic state in 1977-2012

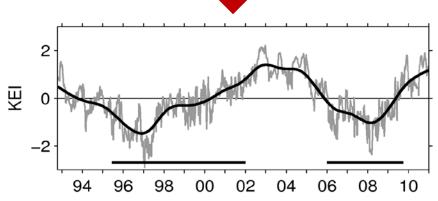


Anomalous wind stress curls regressed to the KE dynamic state in 1977-2012



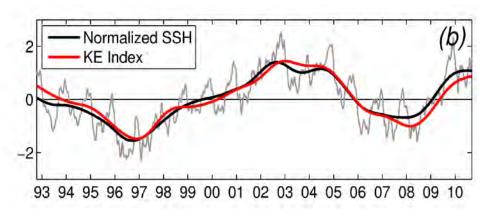
Q3: To what extend can the decadally-modulating KE state be predicted?





KE index: average of the 4 dynamic properties (normalized)

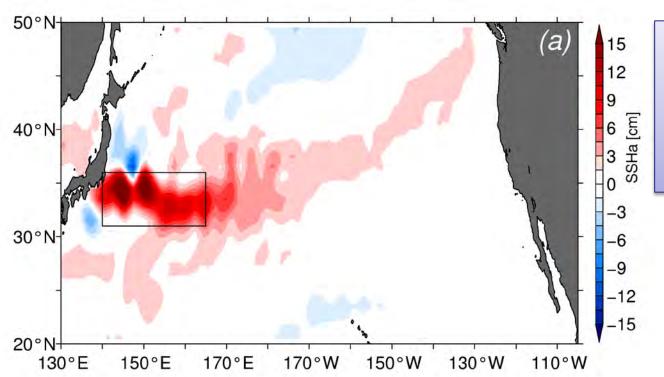
Regression between the KE index and the AVISO SSH anomaly field



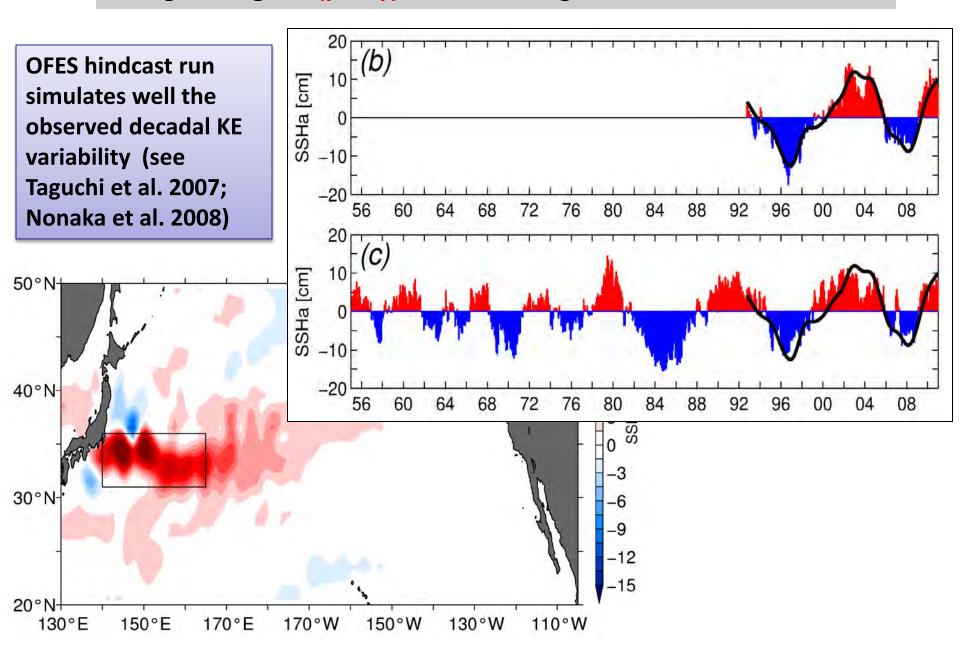
KE index: represented well by SSH anomalies in 31-36°N, 140-165°E

Implications:

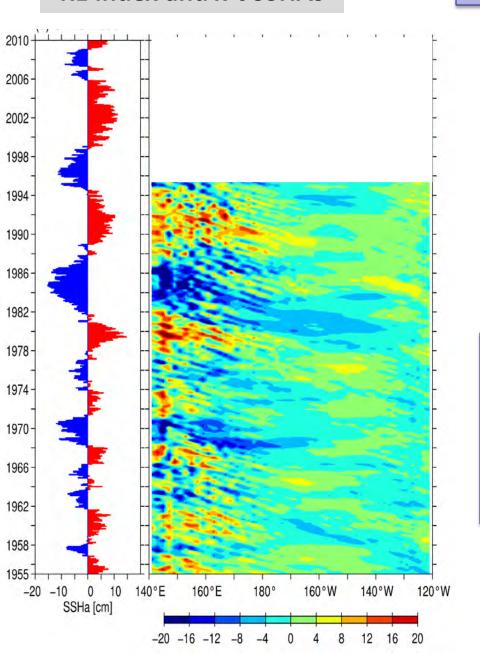
Predicting KE index
becomes equivalent to
predicting SSH
anomalies in this key
box



Lengthening the (proxy) KE index using the OFES hindcast results



KE index and x-t SSHAs



1. Prediction with Rossby wave dynamics

$$h_1(x, t) = h_{obs} [x+c_R(t-t_0), t_0]$$

where

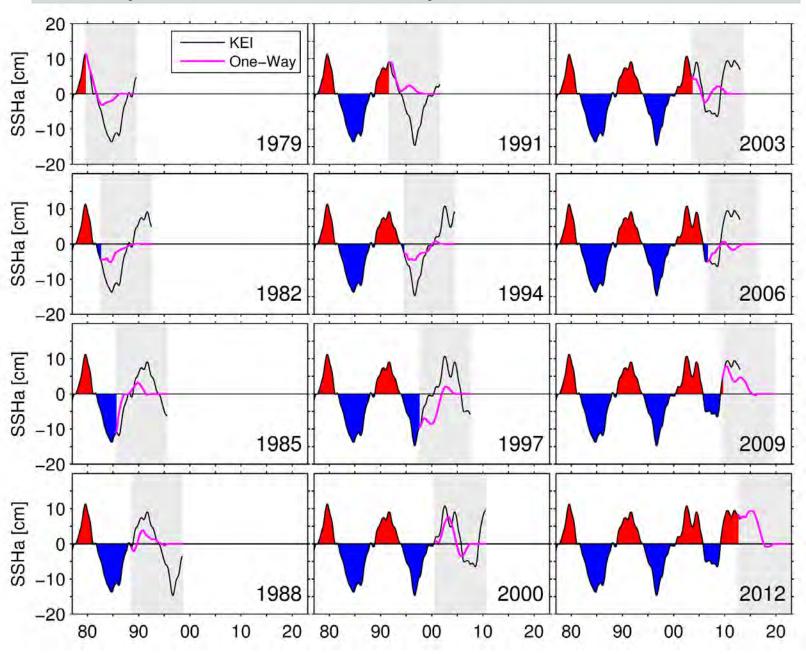
 $h_{obs}(x, t_0)$: initial SSHAs

C_R: Rossby wave speed

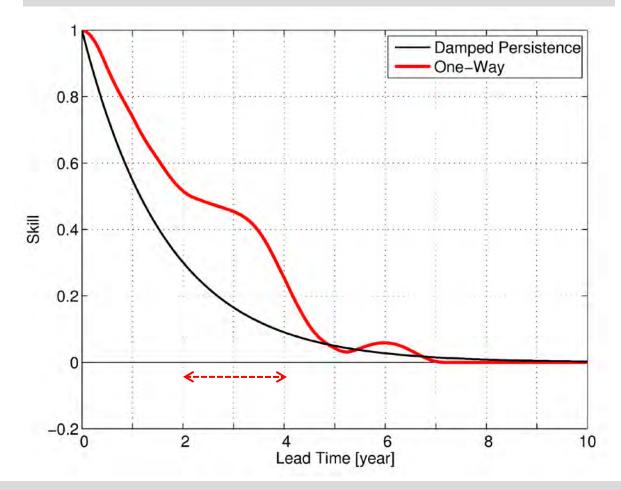
The basis for long-term KE index prediction rests on 2 processes:

1. Oceanic adjustment in mid-latitude North Pacific is via slow, baroclinic Rossby waves

Examples of decadal KE index predictions and verifications

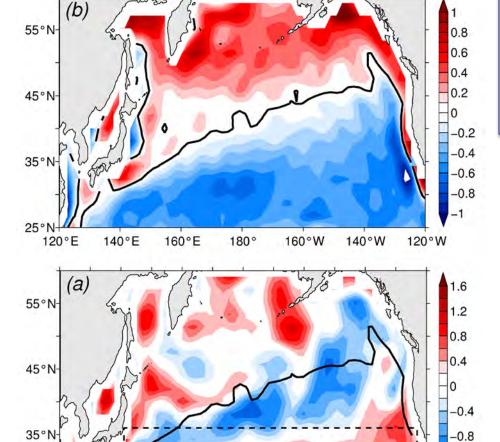


Mean square skill of the predicted KE index



 Compared to damped persistence, wave-carried SSHAs increase predictive skill with a 2~3 yr lead (Schneider and Miller 2001)

NCEP mean wind stress curl field



KE index-regressed curl field (tropical influence removed)

180°

160°W

160°E

140°E

120°E

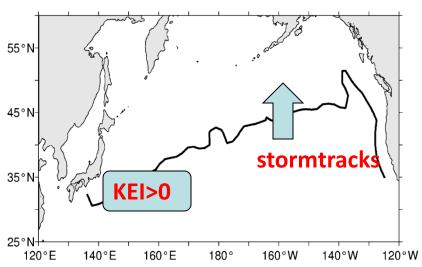
-1.6

120°W

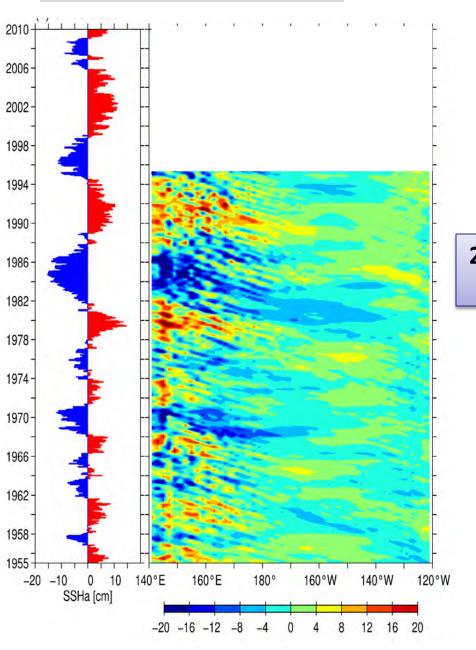
140°W

The basis for long-term KE index prediction rests on 2 processes:

2. The KE decadal variability affects the basin-scale wind stress curl field



KE index and x-t SSHAs



1. Prediction with Rossby wave dynamics

$$h_1(x, t) = h_{obs} [x+c_R(t-t_0), t_0]$$

where

 $h_{obs}(x, t_0)$: initial SSHAs

C_R: Rossby wave speed

2. Prediction with Rossby wave dynamics + KE feedback to wind forcing

$$h_2(x, t) = h_1(x, t) +$$

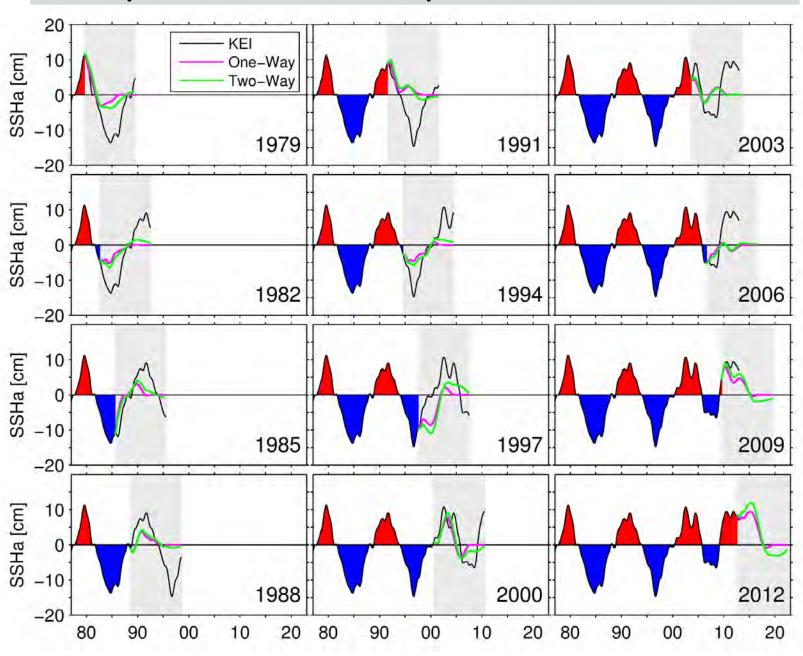
$$\int_{t_0}^{t} b[x+c_R(t'-t_0)] K(t') dt'$$

where

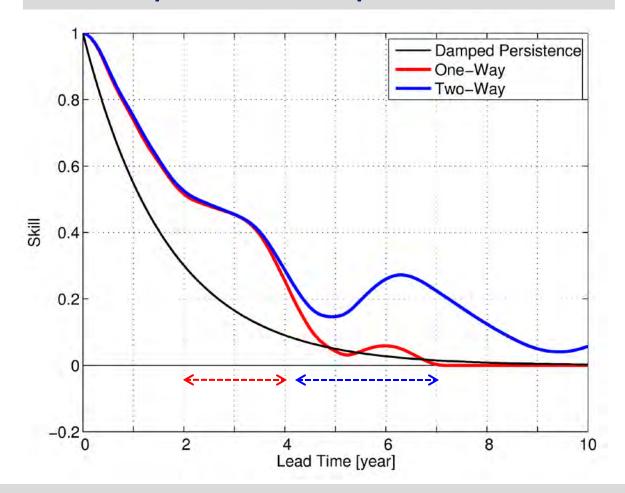
b(x) : feedback coefficient

K(t): forecast KE index

Examples of decadal KE index predictions and verifications

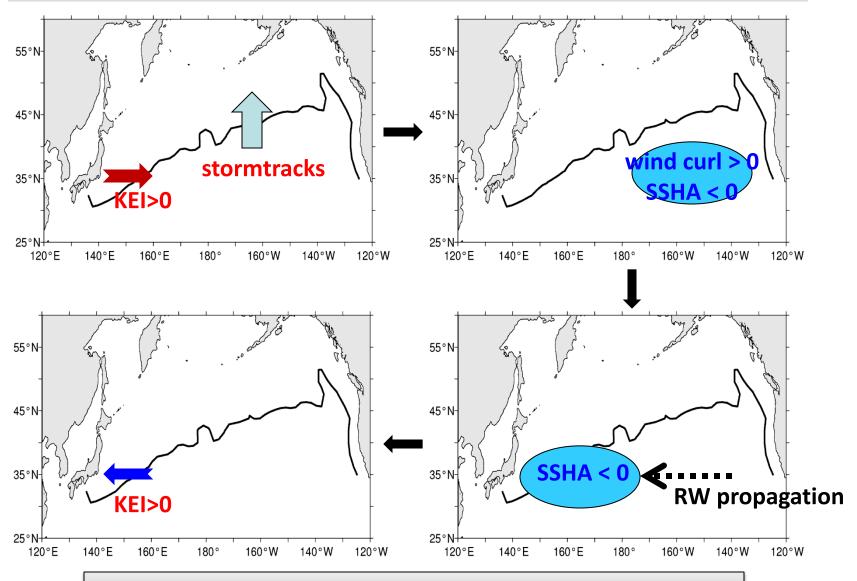


Mean square skill of the predicted KE index



- Compared to damped persistence, wave-carried SSHAs increase predictive skill with a 2~3 yr lead (Schneider and Miller 2001)
- Additional skill is gained with the 4~6 yr lead by considering the wind forcing due to KE feedback

The reason behind enhanced predictive skill with the 4~6 yr lead: delayed negative feedback mechanism



half of the oscillation cycle: ~5 yrs in the N Pacific basin

Summary

- KE dynamic state (i.e. EKE level, path, and jet/RG strengths)
 is dominated by decadal variations. It impacts the cross-front
 SST gradient and overlying stormtracks.
- SSH anomaly signals in 31-36°N, 140-165°E provide a good proxy for the decadally-varying KE system.
- A positive KE index induces overlying-high and downstreamlow pressure anomalies. This feedback favors a coupled mode with a ~10 yr timescale.
 - → Oscillatory nature of this mode enhances predictability
- Compared to Rossby wave dynamics, inclusion of the KEfeedback wind forcing increases predictive skill with a lead time from 2~3 to 4~5 yrs.
- Due to the persistent negative PDO phase, a prolonged stable
 KE dynamic state (until 2017) is predicted.