

FORECASTING THE SEASONAL TO INTERANNUAL VARIABILITY OF EXTREME SEA LEVELS

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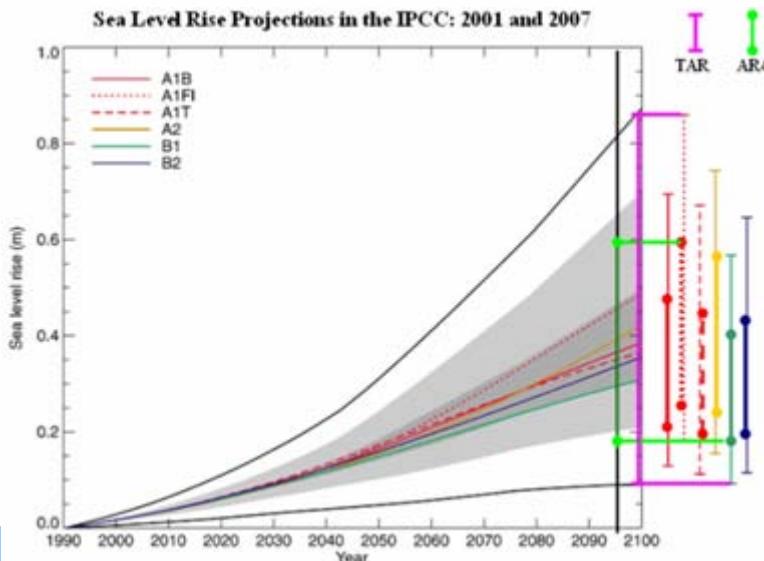
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Importance of extreme sea level values..

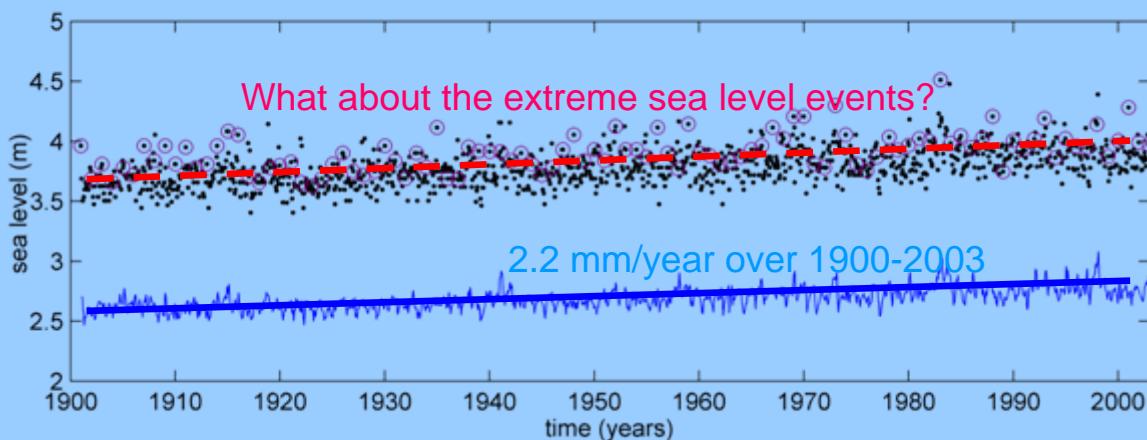
Global annual mean sea level

1.8 mm/year over 1961-2003

3.1 mm/year over 1993-2003



(San Francisco tidal gauge time series, NOAA, station #9414290)



San Francisco
tidal gauge



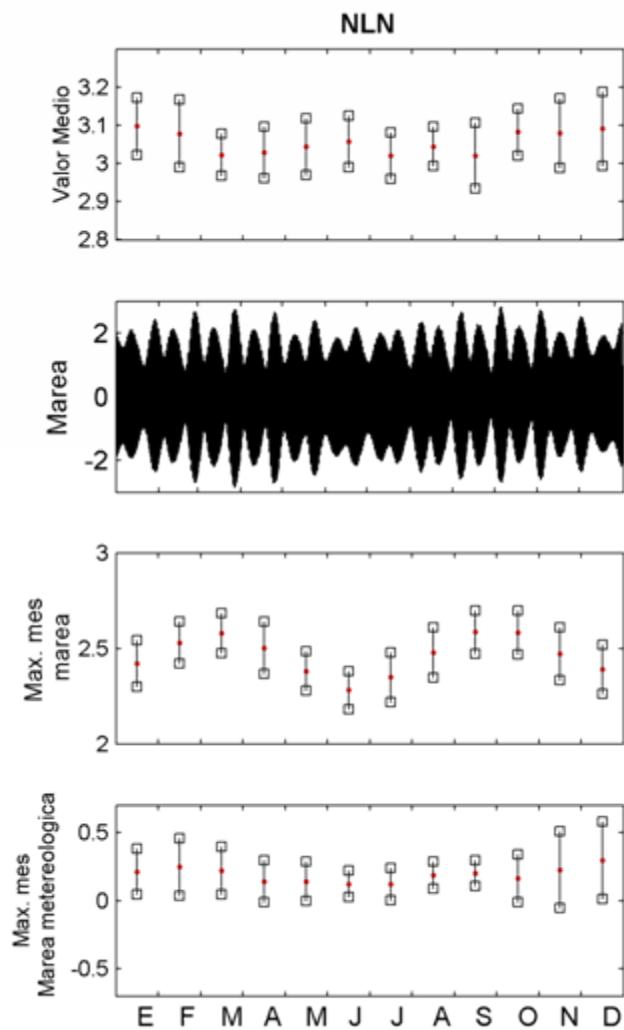
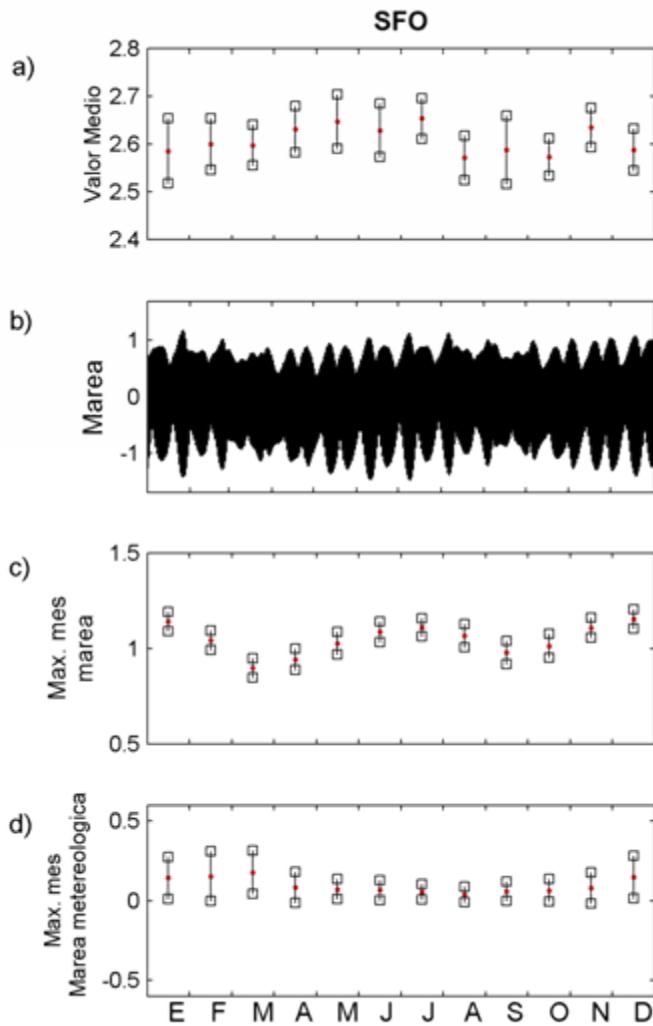
Newlyn
tidal gauge



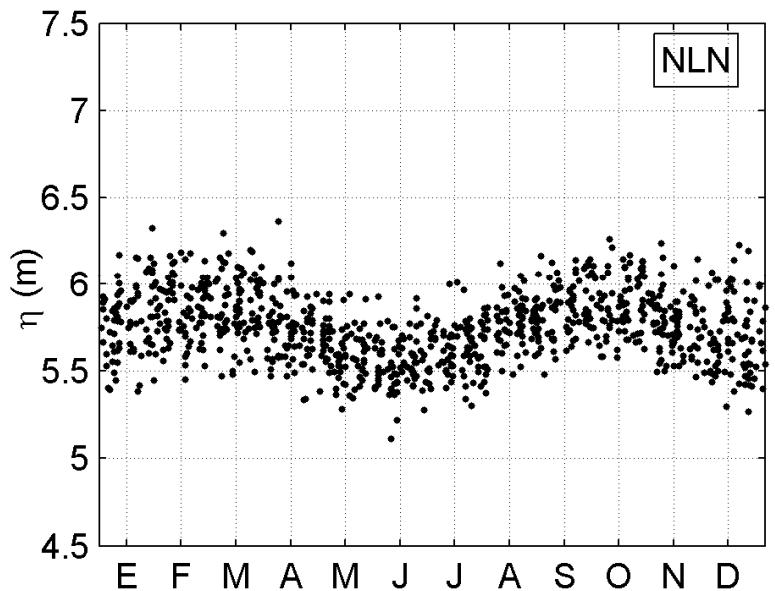
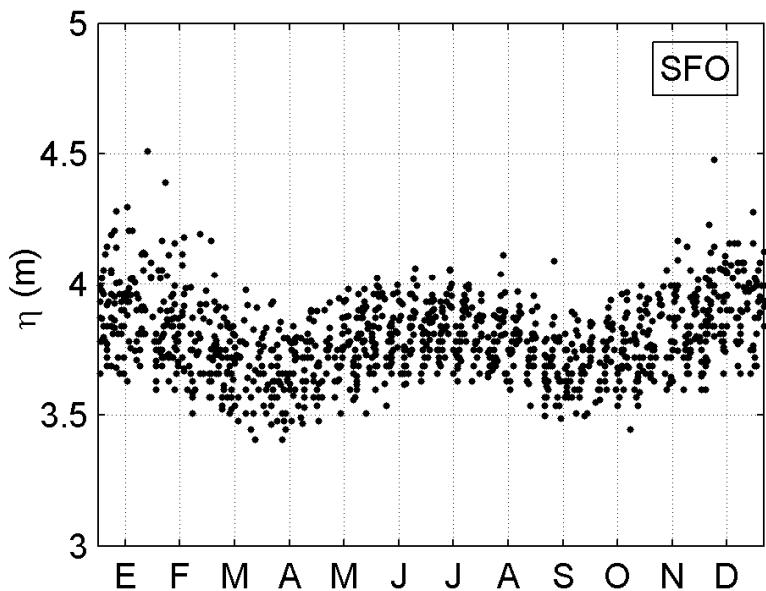
*Data set from NOAA
(California, USA)
(1901-2003)
103 years*

*Data set from BODC
(SW England)
(1915-2001)
87 years*

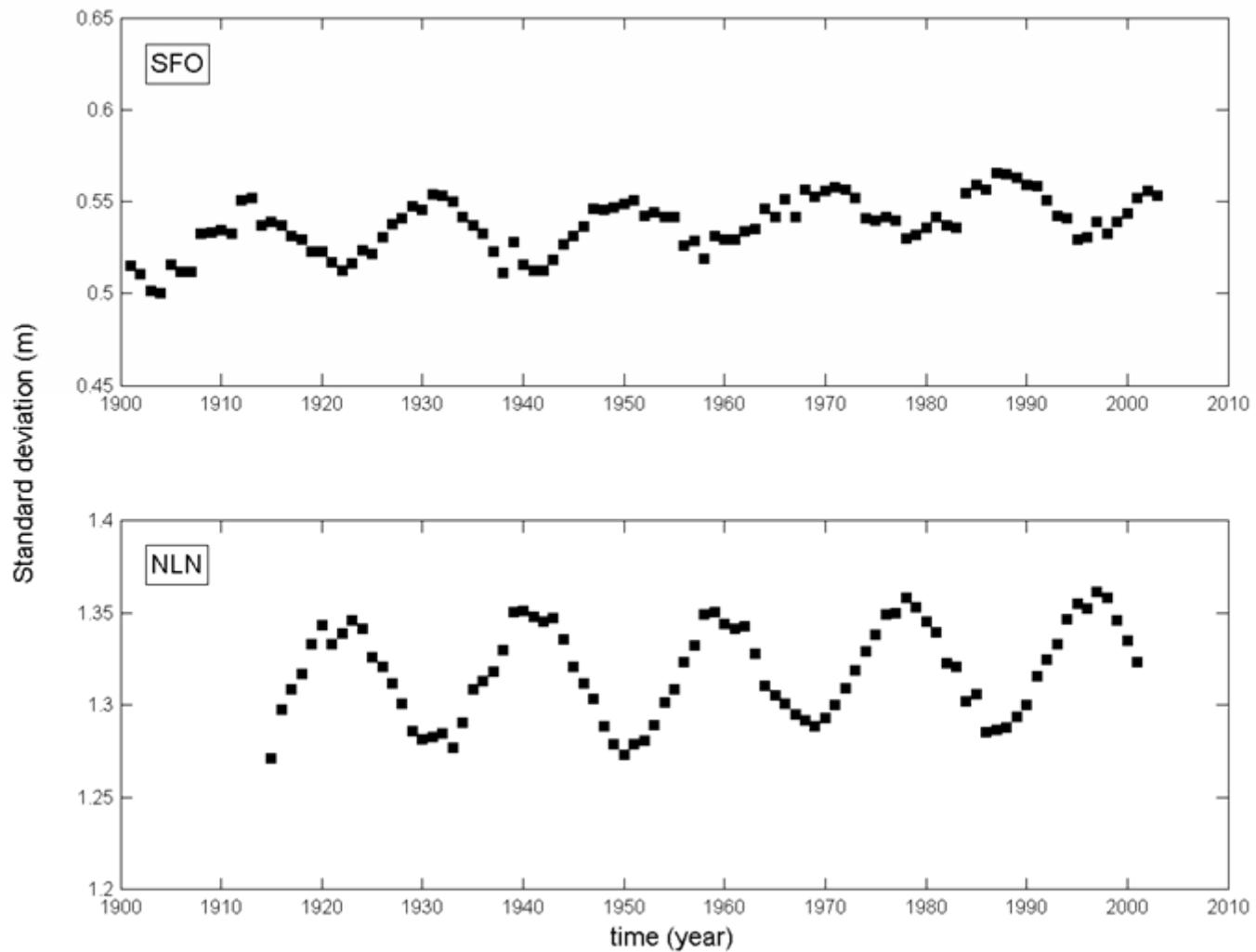
Climatic Variability..



Climatic Variability..



Climatic Variability..



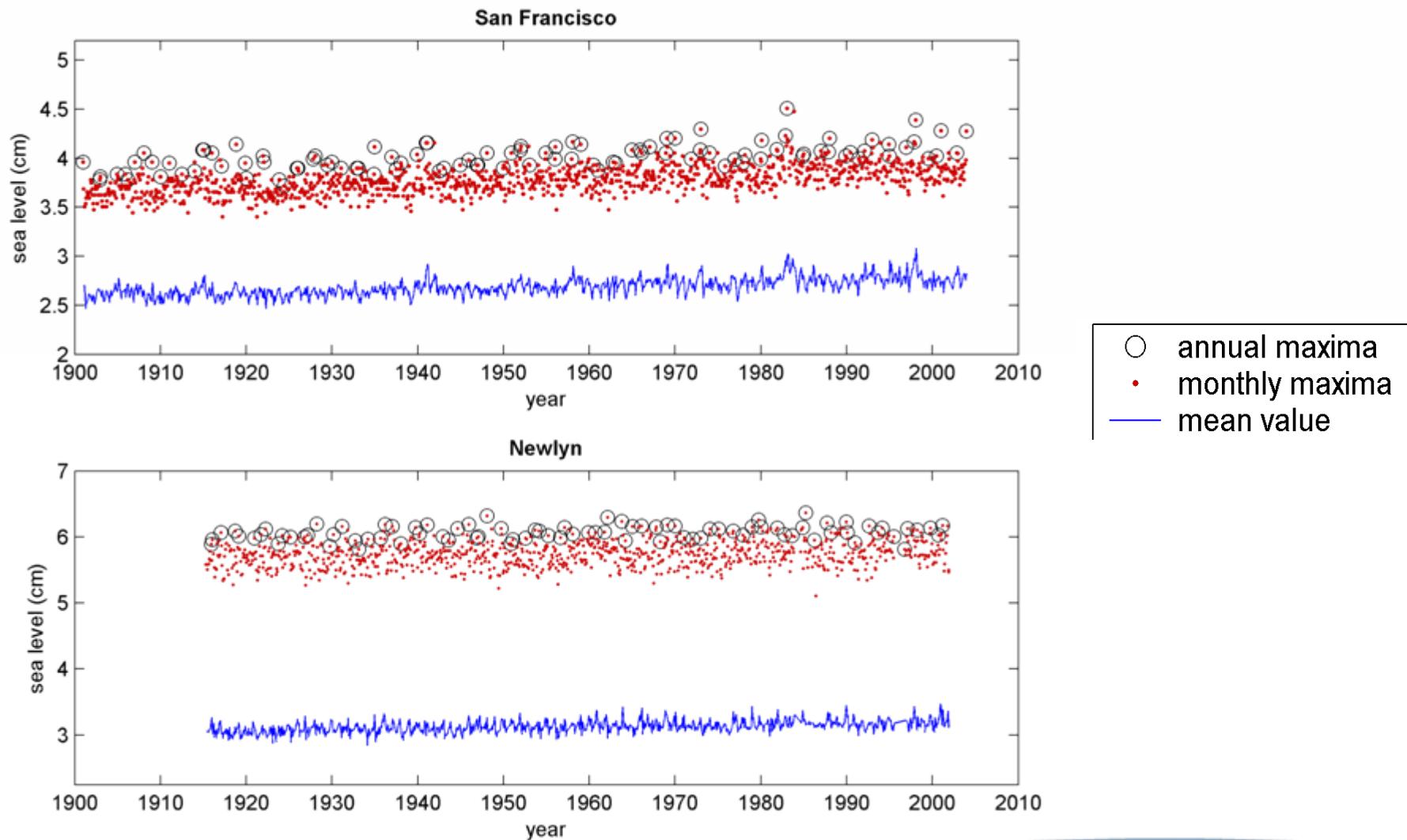
Questions to address:

- Are extreme sea levels being affected by climate change ?
- What is the influence of climate-related indices (NAO, SOI,...) on extreme sea levels?
- How is the within a year variation pattern of the extreme sea level events?
- Can we quantify the influence of each physical process in a extreme value model?

OBJECTIVES

- To develop a Statistical Model to characterize rigorously the extreme values of sea level.
- To account for the relevant seasonal-to-interannual variability of extreme sea levels.

❖ Identifying extreme events..



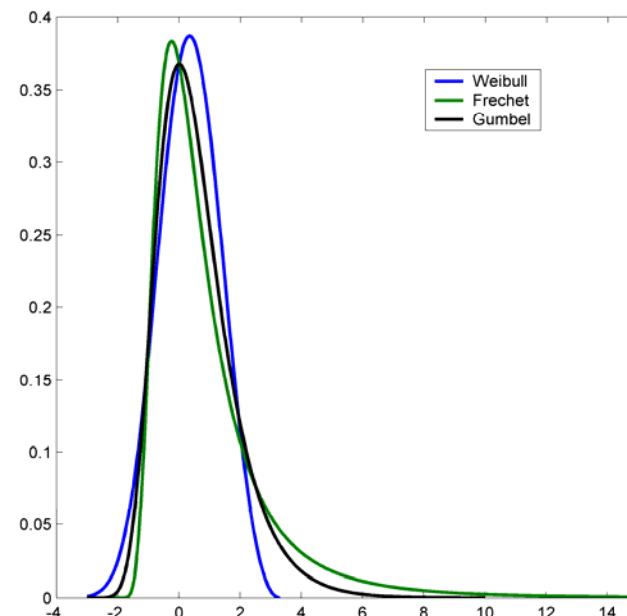
❖ Statistical Model

Extreme Value Theory

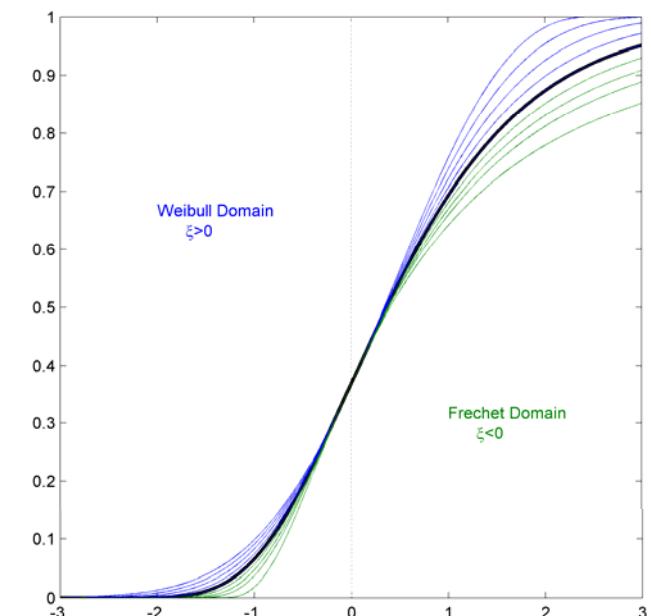
Generalized Extreme Value distribution, GEV

$$F(x; \theta) = \exp \left\{ - \left[1 + \xi \left(\frac{x - \mu}{\psi} \right) \right]^{-1/\xi} \right\}$$

μ → location
 ψ → scale
 ξ → shape



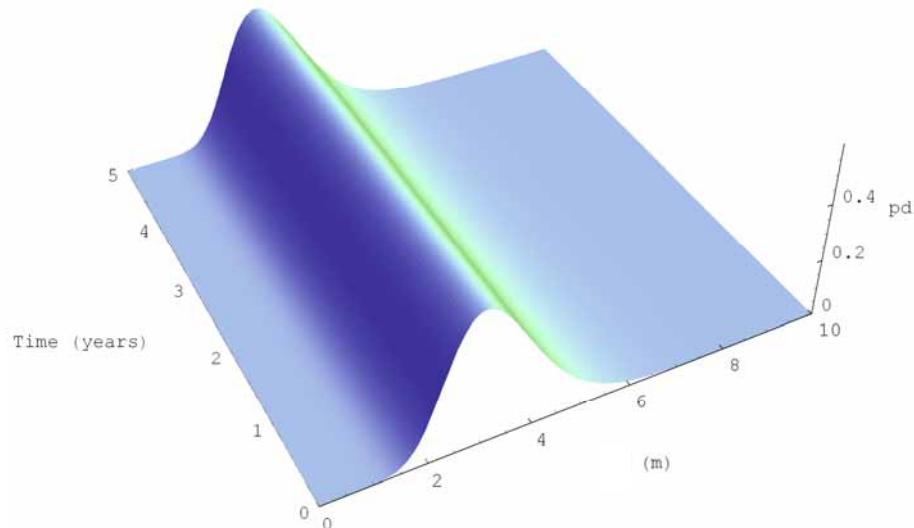
pdf



cdf

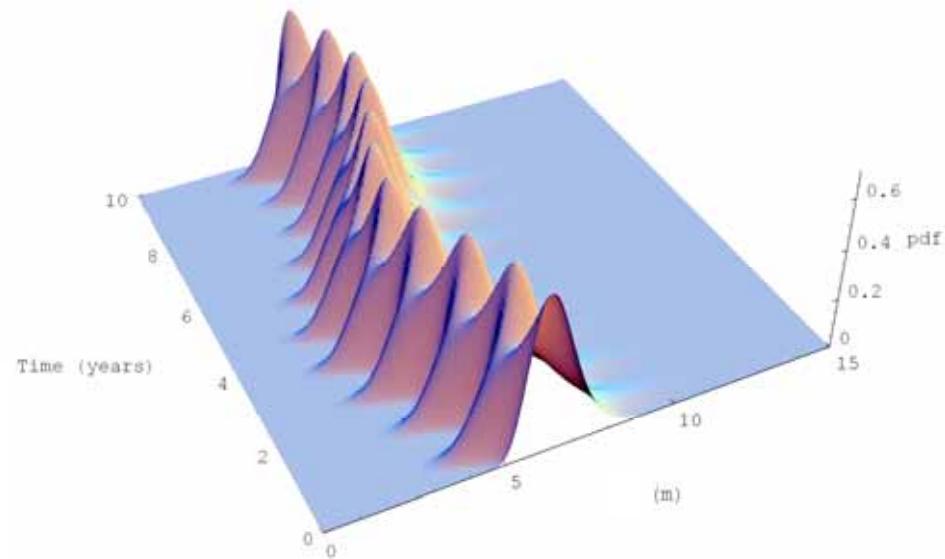
❖ Statistical Model

Stationary process



$$\begin{pmatrix} \mu & = & c t e \\ \psi & = & c t e \\ \xi & = & c t e \end{pmatrix}$$

Non-Stationary process



$$\begin{pmatrix} \mu = \mu_t = f(\text{time}) \\ \psi = \psi_t = f(\text{time}) \\ \xi = \xi_t = f(\text{time}) \end{pmatrix}$$

The probability of an extreme event of a certain magnitude, varies through time

Parameterization

$$\left\{ \begin{array}{l} \mu(t) = \mu_{SLR}(t) + \mu_s(t) \exp[\mu_{LT}(t)] + \mu_N(t) + \mu_{CLI}(t) \\ \psi(t) = \psi_s(t) \exp[\psi_{LT}(t)] + \psi_N(t) + \psi_{CLI}(t) \\ \xi(t) = \xi_s(t) \exp[\xi_{LT}(t)] \end{array} \right.$$

<i>Physical processes</i>	<i>Location</i>	<i>Parameters</i>	<i>Scale</i>	<i>Shape</i>
Seasonality	$\mu_s(t) = \beta_0 + \sum_{i=1}^{P_\mu} [\beta_{2i-1} \cos(2i\pi t) + \beta_{2i} \sin(2i\pi t)]$ ($\beta_0, \beta_1, \beta_2, \beta_3, \beta_4, \dots$)	$\psi_s(t) = \alpha_0 + \sum_{i=1}^{P_\psi} [\alpha_{2i-1} \cos(2i\pi t) + \alpha_{2i} \sin(2i\pi t)]$ ($\alpha_0, \alpha_1, \alpha_2, \alpha_3, \alpha_4, \dots$)		$\xi_s(t) = \gamma_0 + \sum_{i=1}^{P_\xi} [\gamma_{2i-1} \cos(2i\pi t) + \gamma_{2i} \sin(2i\pi t)]$ ($\gamma_0, \gamma_1, \gamma_2, \gamma_3, \gamma_4$)
Sea Level Rise	$\mu_{SLR}(t) = \beta_{SLR}t$ (β_{SLR})		-	-
Nodal Cycle	$\mu_N(t) = \beta_{N_1} \cos(2\pi t / T_N) + \beta_{N_2} \sin(2\pi t / T_N)$ (β_{N_1}, β_{N_2})	$\psi_N(t) = \alpha_{N_1} \cos(2\pi t / T_N) + \alpha_{N_2} \sin(2\pi t / T_N)$ ($\alpha_{N_1}, \alpha_{N_2}$)		-
Climatic patterns	$\mu_{CLI}(t) = \begin{cases} \beta_{SOI} SOI(t) \\ \beta_{NAO} NAO(t) \end{cases}$ (β_{SOI}, β_{NAO})	$\psi_{CLI}(t) = \begin{cases} \alpha_{SOI} SOI(t) \\ \alpha_{NAO} NAO(t) \end{cases}$ ($\alpha_{SOI}, \alpha_{NAO}$)		-
Secular Trends	$\mu_{LT}(t) = \beta_{LT}t + \beta_{LT_2}t^2$ (β_{LT}, β_{LT_2})	$\psi_{LT}(t) = \alpha_{LT}t$ (α_{LT})		$\xi_{LT}(t) = \gamma_{LT}t$ (γ_{LT})

Parameterization

$$\mu(t) = [\beta_0 + \beta_1 \cos(2\pi t) + \beta_2 \sin(2\pi t) + \beta_3 \cos(4\pi t) + \beta_4 \sin(4\pi t)]$$

Annual cycle

1 0

Codification

$$\exp(\beta_{LT}t + \beta_{LT2}t^2) + \beta_{SOI}SOI(t)$$

Semiannual cycle

1 0

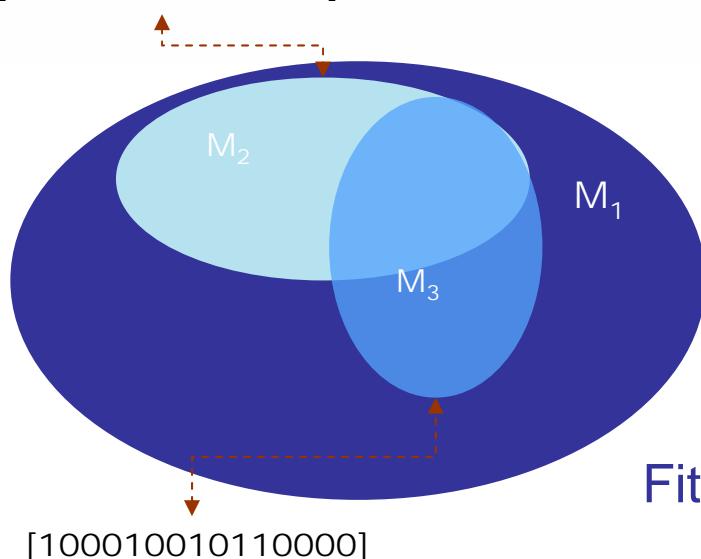
$$1 0 \quad 1 0$$

Parabolic trend El Niño

Linear trend

1 0

[110010010110000]



[100010010110000]

A possible model: [1100110110111100]
(binary chromosomes)

Automatic Selection
STEPWISE

Fitness: Maximum Likelihood Estimation

$$l(\theta | t_i, z_i) = -\sum_{i=1}^m \left\{ \log \psi(t_i) + (1 + 1/\xi(t_i)) \log \left[1 + \xi(t_i) \left(\frac{z_i - \mu(t_i)}{\psi(t_i)} \right) \right] + \left[1 + \xi(t_i) \left(\frac{z_i - \mu(t_i)}{\psi(t_i)} \right) \right]^{-1/\xi(t_i)} \right\}$$

Statistical Information criterion (BIC & HQ)

Stationary model

Best model

MLE (SE)	SFO-MM0	SFO-MM1	SFO-MM2	SFO-MM3	SFO-MM4	SFO-MM5	SFO-MM7	SFO-MM9	SFO-MM11	SFO-MM12
β_0 (m)	3.728 (0.004)	3.6248 (0.004)	3.626 (0.004)	3.6239 (0.004)	3.6326 (0.003)	3.6353 (0.003)	3.6354 (0.003)	3.6154 (0.005)	3.6146 (0.005)	3.6189 (0.005)
β_1 (m)	—	—	—	—	0.0214 (0.004)	0.0296 (0.004)	0.0269 (0.004)	0.0265 (0.004)	0.0261 (0.004)	0.0271 (0.004)
β_2 (m)	—	—	—	—	0.0302 (0.004)	0.0286 (0.004)	0.0262 (0.004)	0.027 (0.004)	0.0276 (0.004)	0.0283 (0.004)
β_3 (m)	—	—	—	—	0.0823 (0.004)	0.0856 (0.004)	0.0858 (0.004)	0.0842 (0.004)	0.0855 (0.004)	0.0854 (0.003)
β_4 (m)	—	—	—	—	0.0497 (0.004)	0.048 (0.003)	0.048 (0.003)	0.0489 (0.003)	0.0475 (0.003)	0.0488 (0.003)
β_{LT} (yr^{-1})	—	—	—	—	—	—	—	0.000 09 (0.000 02)	0.000 09 (0.000 02)	0.000 09 (0.000 02)
β_{N_1} (m)	—	—	0.0064 (0.005)	—	0.0087 (0.004)	0.0096 (0.003)	0.0095 (0.003)	0.0103 (0.003)	0.0101 (0.003)	0.0071 (0.003)
β_{N_2} (m)	—	—	0.0199 (0.005)	—	0.0201 (0.004)	0.0203 (0.003)	0.021 (0.003)	0.0214 (0.003)	0.0198 (0.003)	0.0236 (0.003)
β_{SOI} (m unit $^{-1}$)	—	—	—	0.0304 (0.004)	—	—	—	—	—	0.0237 (0.003)
α_0 (m)	0.1364 (0.003)	0.1144 (0.003)	0.1133 (0.002)	0.1121 (0.003)	0.0865 (0.002)	0.0858 (0.002)	0.0857 (0.002)	0.085 (0.002)	0.0776 (0.003)	0.0735 (0.003)
α_1 (m)	—	—	—	—	—	0.0274 (0.002)	0.0261 (0.003)	0.0253 (0.002)	0.023 (0.002)	0.0186 (0.002)
α_2 (m)	—	—	—	—	—	0.0072 (0.002)	0.0084 (0.003)	0.008 (0.003)	0.0083 (0.002)	0.0072 (0.003)
α_{LT} (yr^{-1})	—	—	—	—	—	—	—	0.0018 (0.000 76)	0.002 (0.000 75)	0.002 (0.000 75)
γ_0	0.1319 (0.013)	0.1112 (0.012)	0.1077 (0.014)	0.1193 (0.012)	0.0743 (0.015)	0.1177 (0.015)	0.1307 (0.016)	0.1339 (0.016)	0.1294 (0.017)	0.1328 (0.017)
γ_1	—	—	—	—	—	—	0.0461 (0.02)	0.0613 (0.02)	0.0686 (0.022)	0.0768 (0.021)
γ_2	—	—	—	—	—	—	0.0519 (0.026)	0.0441 (0.025)	0.0421 (0.026)	0.0325 (0.025)
ℓ	315.71	818.61	828.17	848.4	1138.51	1207.67	1211.63	1224.14	1227.66	1267.21
p	3	3	5	4	9	11	13	14	15	16
AIC	625.42	1631.22	1646.34	1688.80	2259.02	2393.34	2397.26	2420.28	2425.32	2502.42
BIC	610.06	1615.87	1620.74	1668.32	2212.97	2337.05	2330.74	2348.65	2348.55	2420.53

Best model

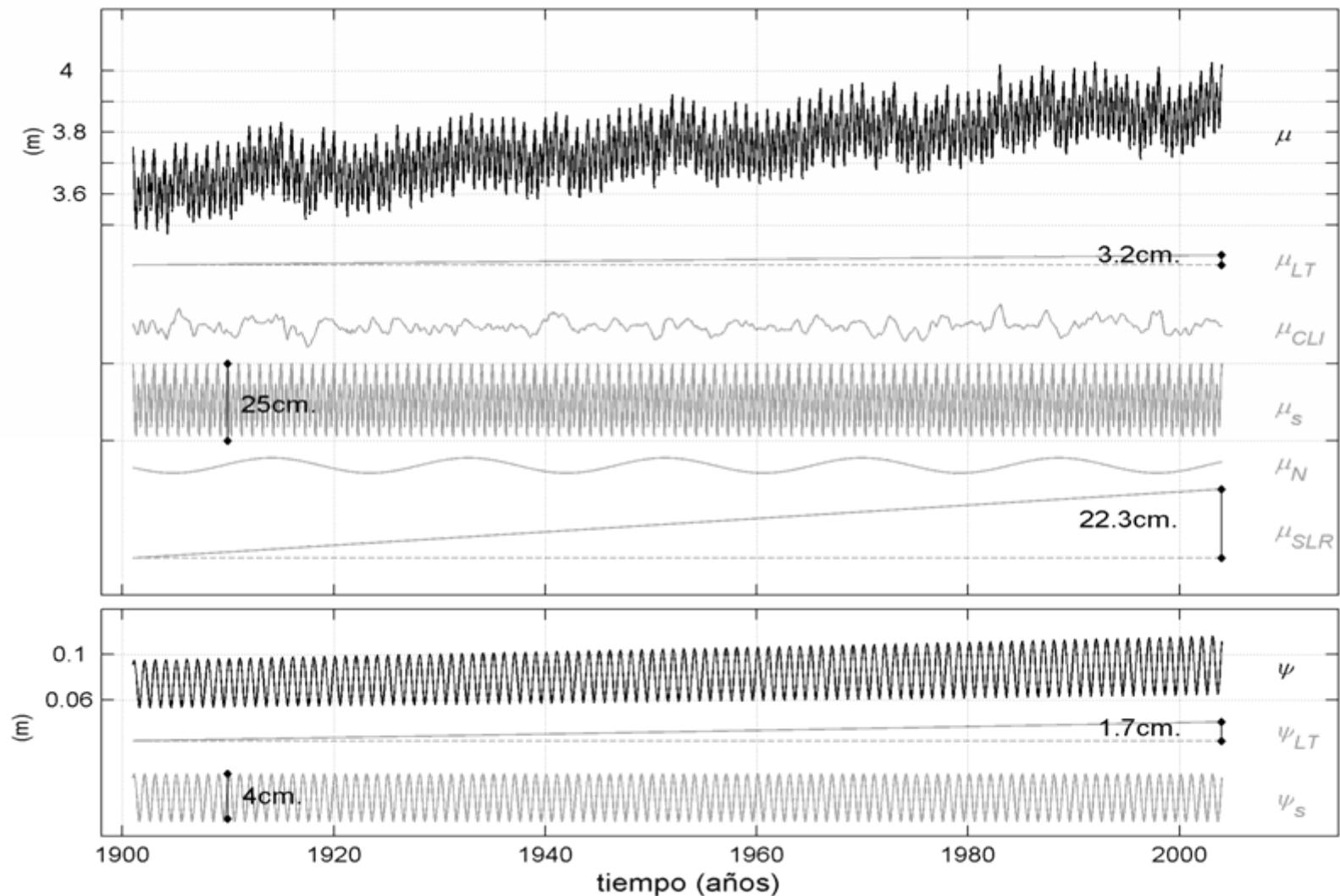
$$\mu(t) = \beta_{SLR} t + [\beta_0 + \beta_1 \cos(2\pi t) + \beta_2 \sin(2\pi t) + \beta_3 \cos(4\pi t) + \beta_4 \sin(4\pi t)] e^{\beta_{LT} t} + \beta_{N_1} \cos(2\pi t/T_N)$$

$$+ \beta_{N_2} \sin(2\pi t/T_N) + \beta_{SOI} \text{SOI}(t)$$

$$\psi(t) = [\alpha_0 + \alpha_1 \cos(2\pi t) + \alpha_2 \sin(2\pi t)] e^{\alpha_{LT} t}$$

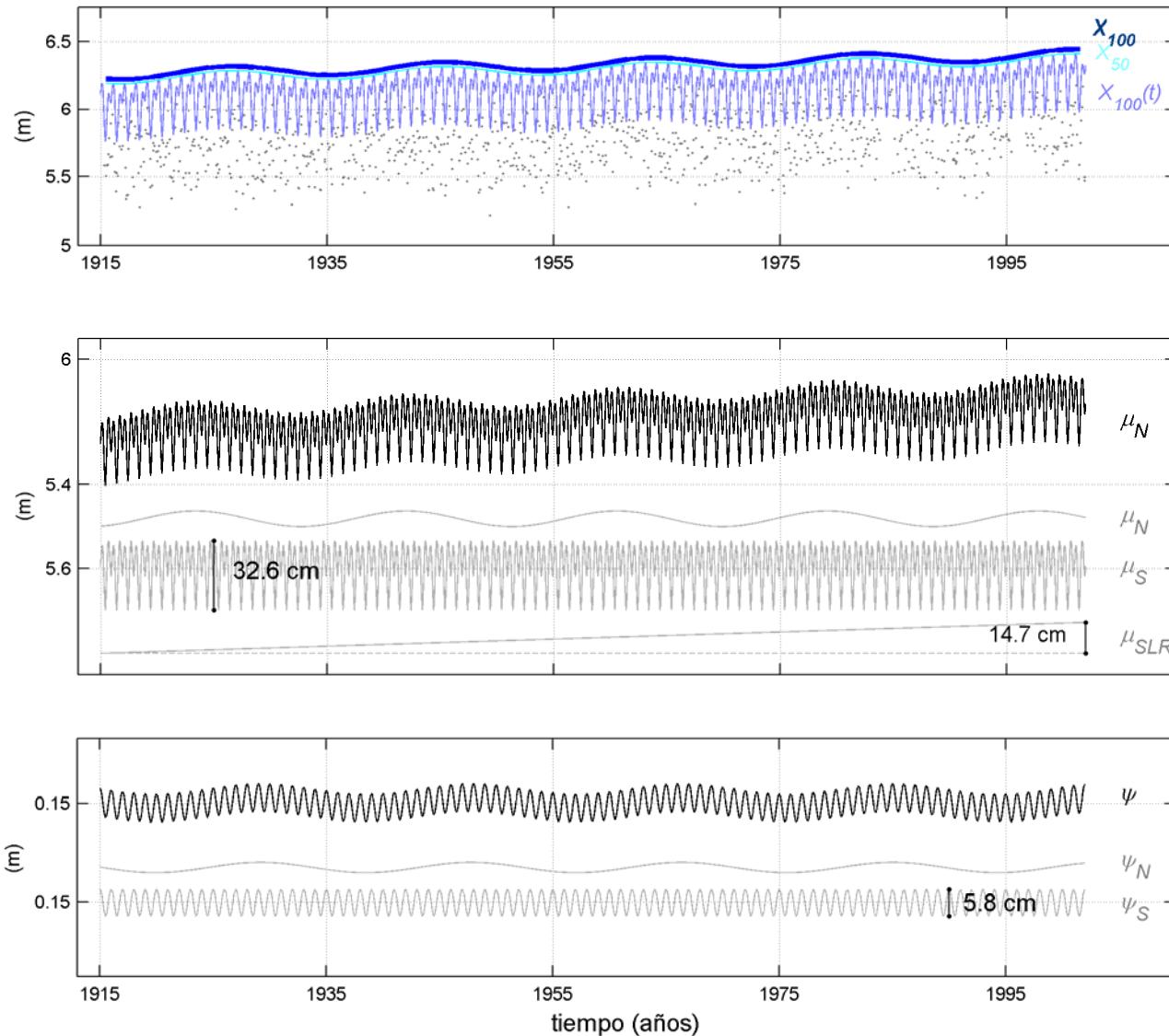
$$\xi(t) = [\gamma_0 + \gamma_1 \cos(2\pi t) + \gamma_2 \sin(2\pi t)].$$

San Francisco

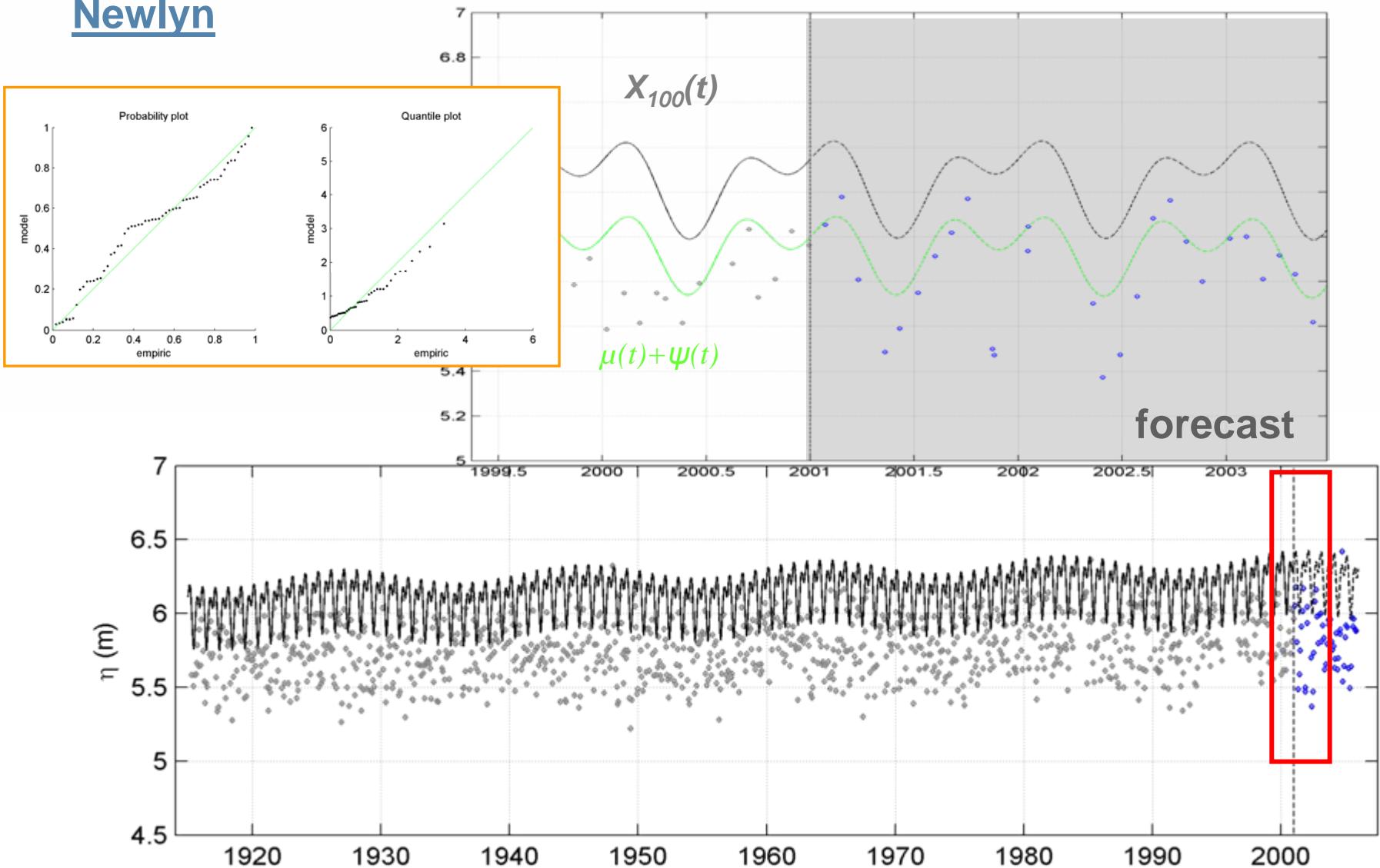


- Increase of 3.2 cm (above the 22.3 cm of SLR)
- Increase of 23% of the variability in the last century!!

Newlyn



Newlyn



- The model is valid for a complete analysis of seasonal-to-interannual sea level extremes providing time dependent quantiles and confidence intervals
- The modelling of the different time scales makes for a better understanding of recent secular trends of extreme climate events, which are one of the main concerns nowadays
- The statistical model is able to predict in the short-term time scale (for example in the next 12 months) the probability of exceedence of a certain extreme sea level value.

References:

Méndez, F.J., Menéndez, M., Luceño, A., Losada, I.J. (2007) *Analyzing monthly extreme sea levels with a time-dependent GEV model*, *Journal of Atmospheric and Oceanic Technology*, 24, 5, 894–911.



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