

Multiple stressors in the coastal ocean

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Overview

- **Ocean Acidification (OA):** Another CO₂ problem has emerged— yet coastal ocean more complex
- **Coastal ecosystems under multiple forcings:** temp rising + O₂ decline+ acidification within a similar time frame
- **Need consider the hydrodynamics:** e.g., Upwelling/Submarine Groundwater Discharge
- **OA observation system & multidisciplinary researches essential and consider the multiple stressors at a system level**

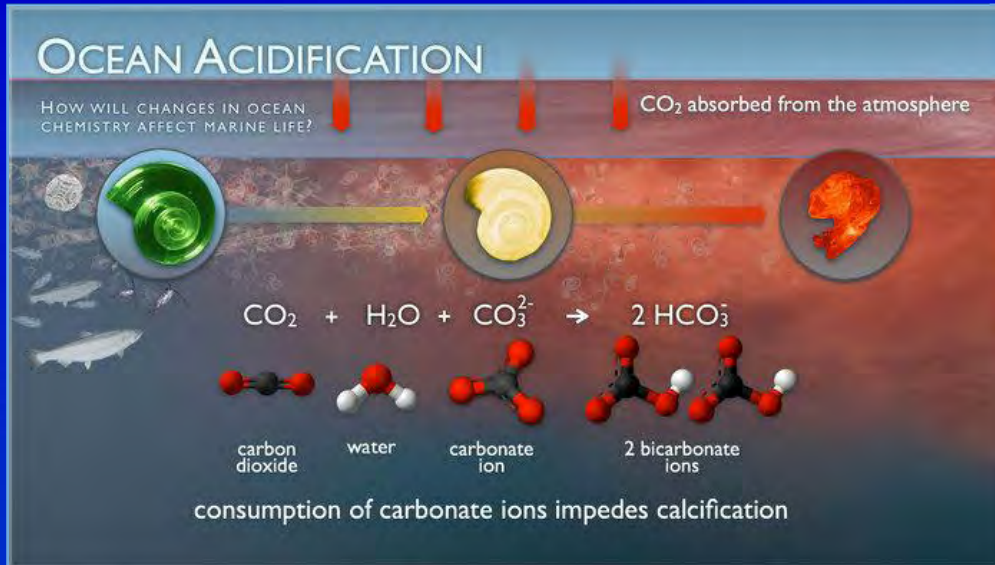


Outline

- **Coastal Ocean Acidification**
- **Multiple stressors in the Coastal Ocean**
- **Concluding Remarks**



Ocean Acidification: another CO₂ problem: increase in [H⁺] or drawdown of pH



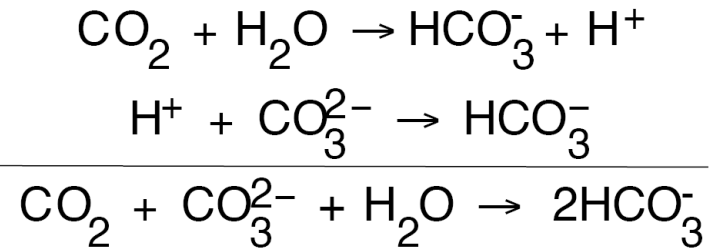
From PMEL

$$\text{pH} = -\log (\text{H}^+) = -\log \gamma_{\text{H}}\{\text{H}^+\}$$

CaCO₃ saturation state:

$$\Omega = [\text{Ca}^{2+}][\text{CO}_3^{2-}]/K_{\text{sp}}'$$

$$\Omega_a > 1 \sim \text{supersaturated}$$



When CO₂ invades sea water:

- [HCO₃⁻] increases
- [CO₃²⁻] decreases
- Ω decreases
- a small part of HCO₃⁻ formed dissociates into carbonates + H⁺ (“ocean acidification”)

Calcification rate vs. Ω_{arag} in coral reef systems

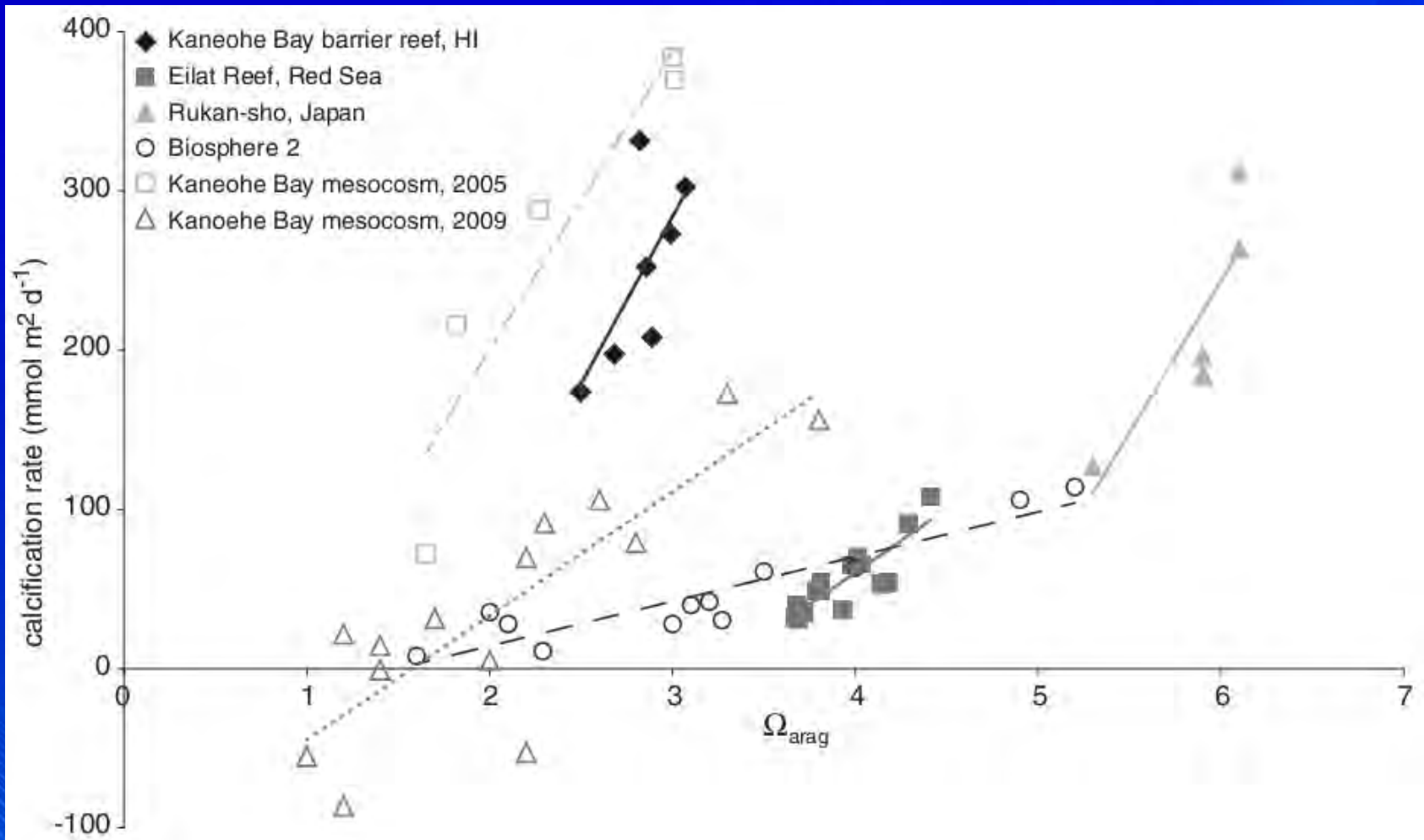
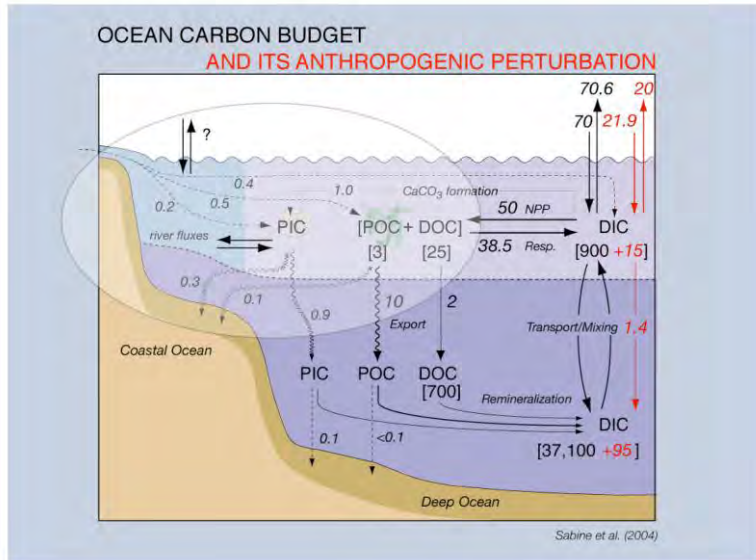


Fig. 6. Ω_{arag} versus calcification rate for coral reef ecosystems (solid symbols) and coral mesocosms (open symbols). Lines are the linear least-squares best fit of the data. References for data are: Kaneohe Bay barrier reef data (black diamonds and solid line), this study; Eilat Reef, Red Sea (dark gray squares and solid line) (Silverman et al., 2007a, b); Rukan-sho, Japan (light gray triangles and solid line) (Ohde and van Woessik, 1999); Biosphere 2 (black circles and dashed line) (Langdon et al., 2000, 2003); Kaneohe Bay mesocosm, 2005 (light gray squares and dashed-dot line); Kaneohe Bay mesocosm, 2009 (dark gray triangles and dotted line) (Andersson et al., 2009).

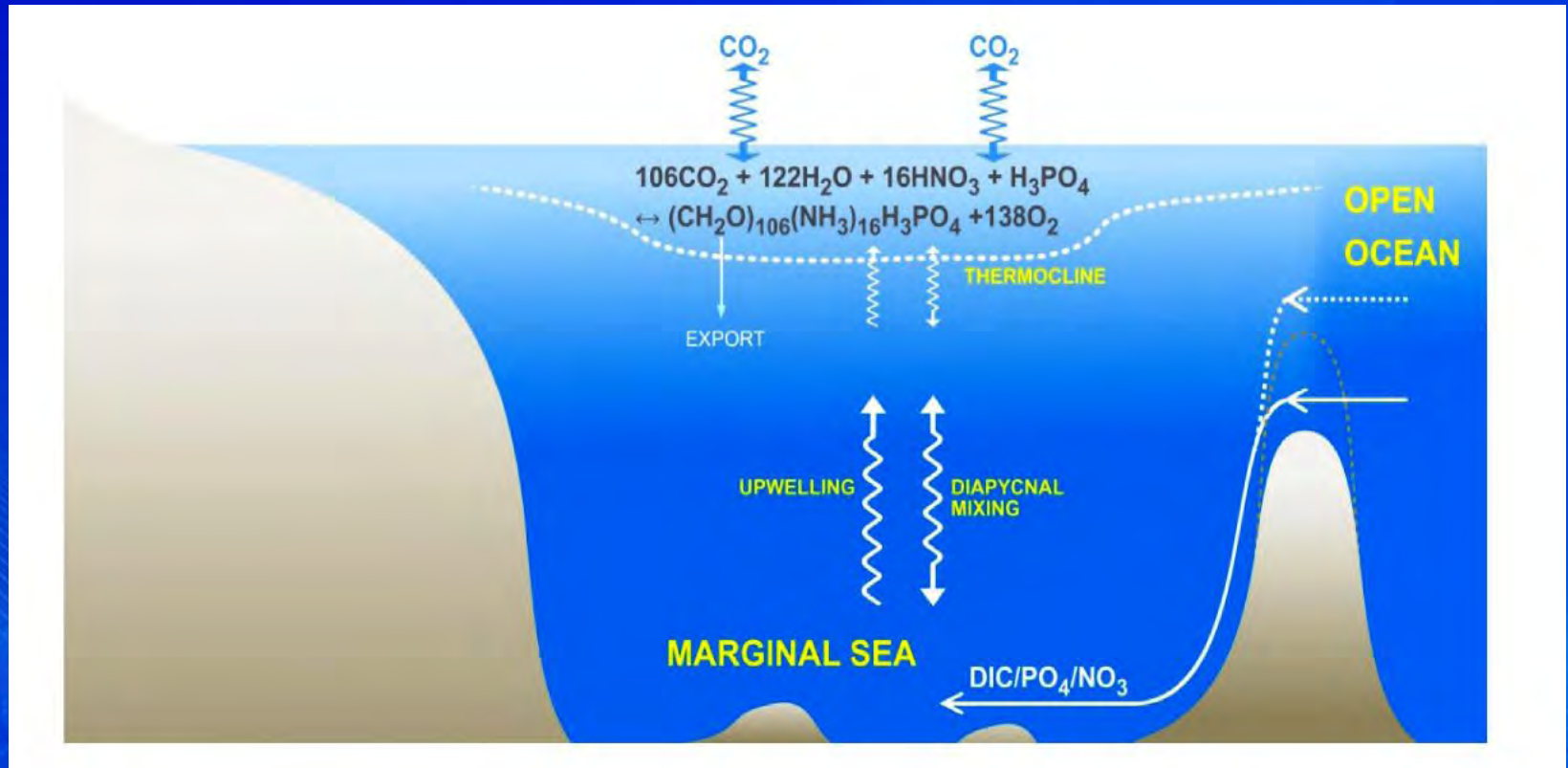
Why Coastal Ocean?



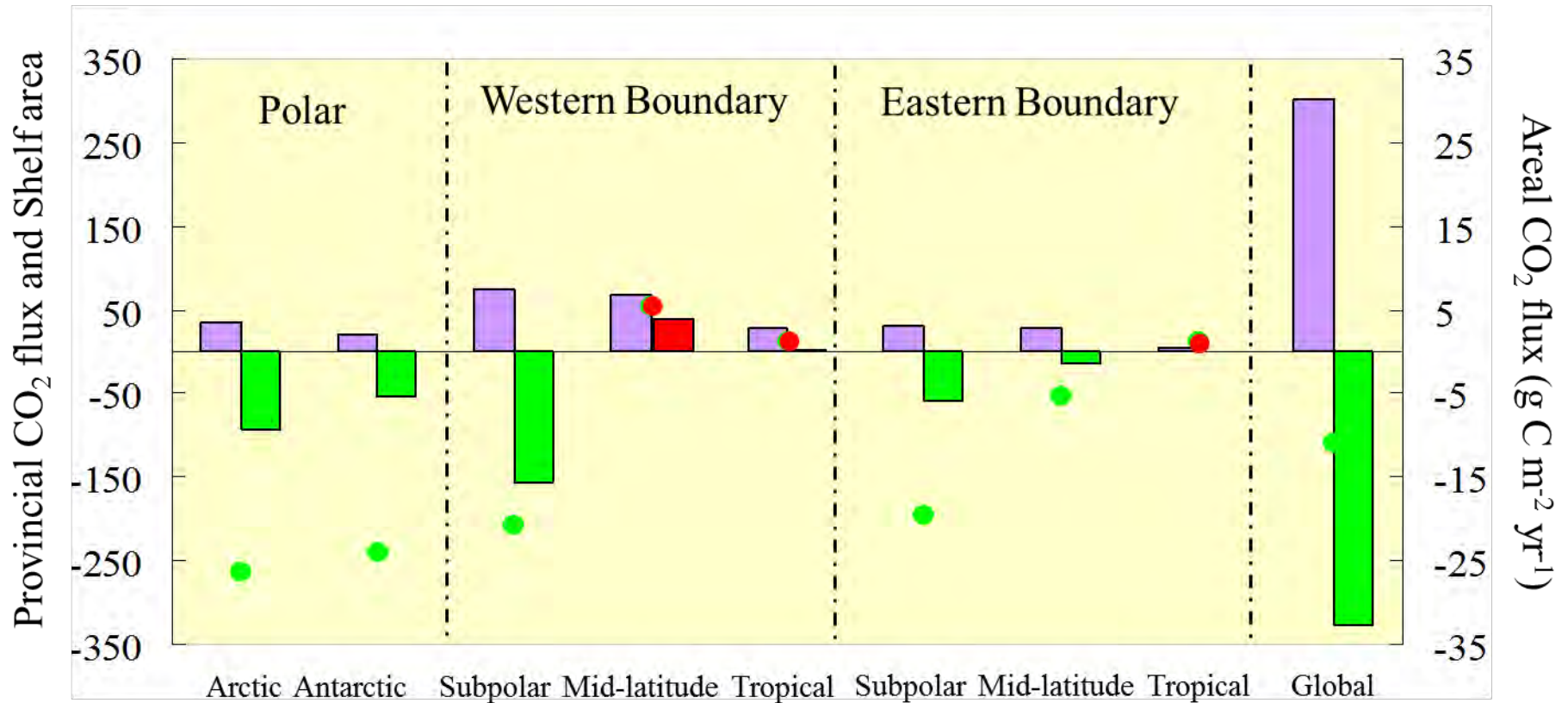
- A unique physical-biogeochemical ecosystem links the land and the open ocean but vulnerable
- Boundary processes across the land-margin and margin-ocean are key drivers
- Characterized by complex circulations, abundant river/groundwater input, dynamic sediment boundary and high productivity: large gradients chemically and biologically

Why are some marginal seas sources of atmospheric CO₂?

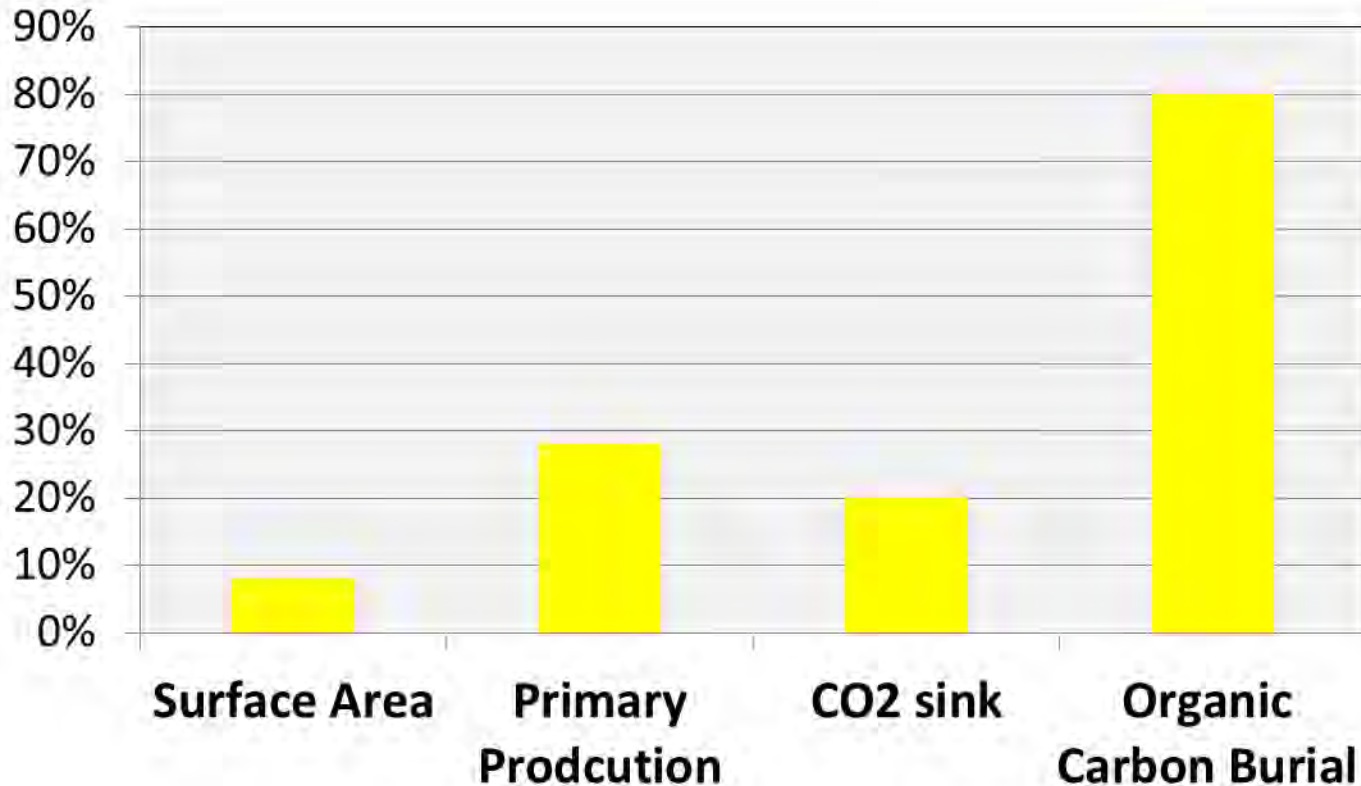
Minhan Dai,¹ Zhimian Cao,¹ Xianghui Guo,¹ Weidong Zhai,¹ Zhiyu Liu,¹ Zhiqiang Yin,¹
Yanping Xu,¹ Jianping Gan,² Jianyu Hu,¹ and Chuanjun Du¹



An updated province –based global shelf air-sea CO₂ flux:



Coastal ocean mitigates more CO₂ than the open ocean



pH dynamics in different marine systems

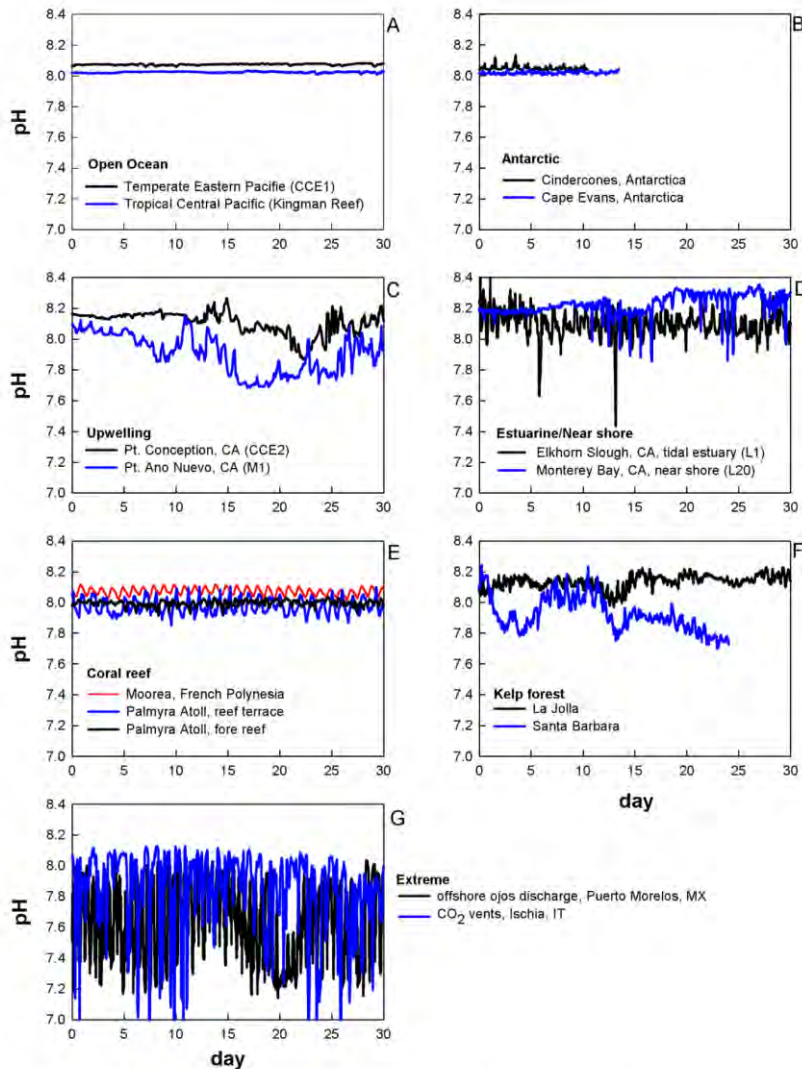


Figure 2. pH dynamics at 15 locations worldwide in 0-15 m water depth. All panels are plotted on the same vertical range of pH (total hydrogen ion scale). The ordinate axis was arbitrarily selected to encompass a 30-day period during each sensor deployment representative of each site during the deployment season. See Table 1 for details regarding sensor deployment. doi:10.1371/journal.pone.0028983.g002

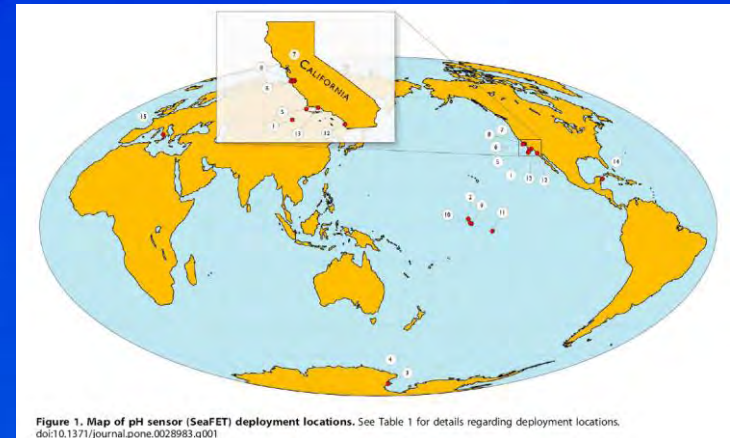


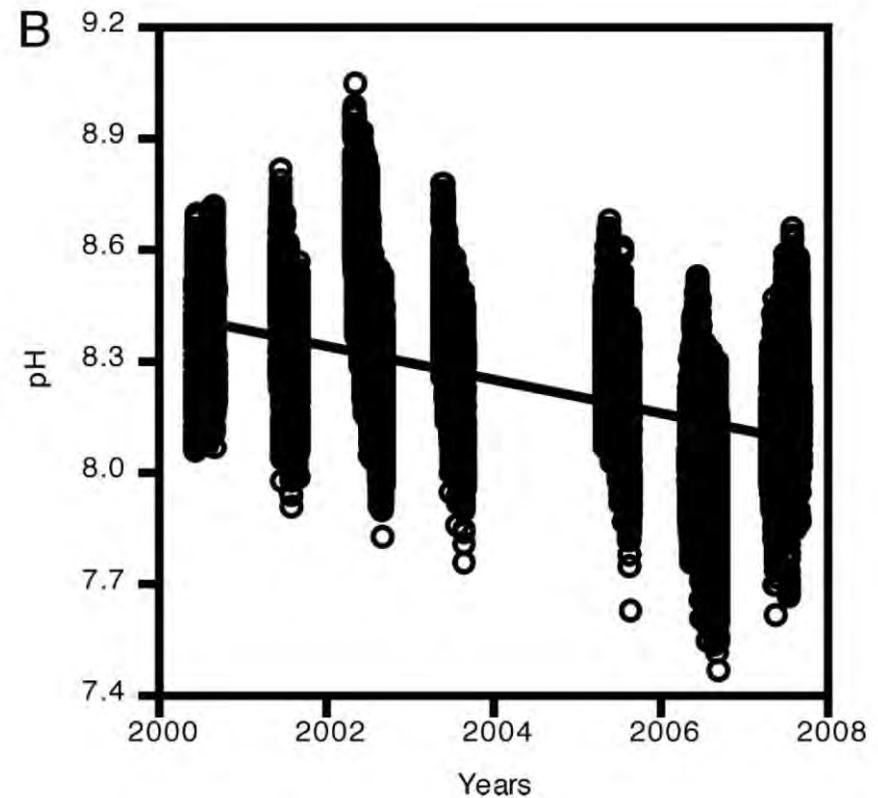
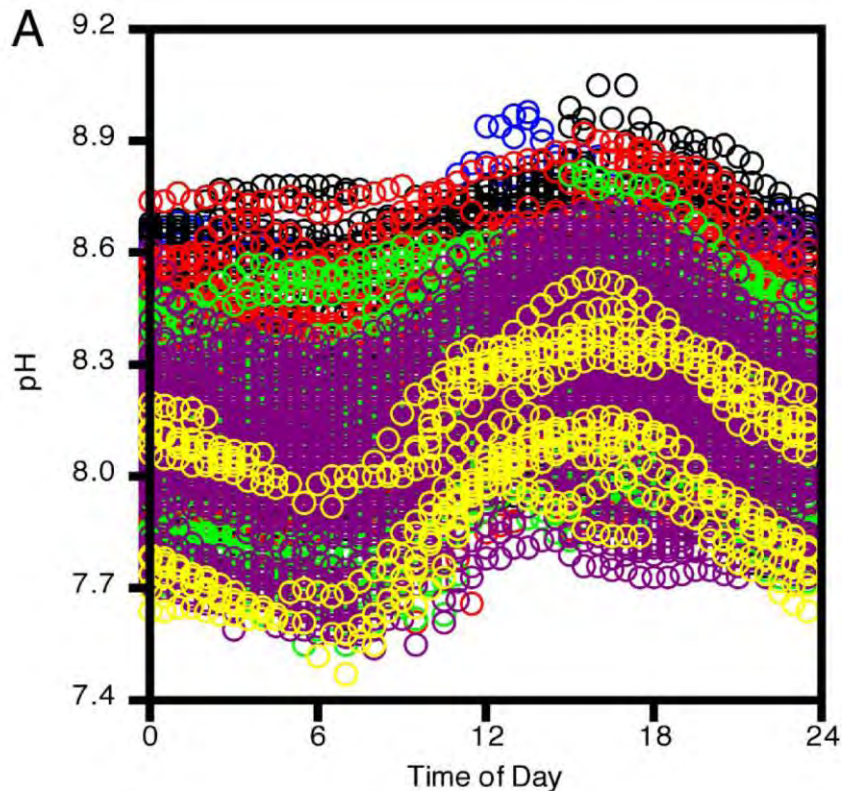
Figure 1. Map of pH sensor (SeaFET) deployment locations. See Table 1 for details regarding deployment locations. doi:10.1371/journal.pone.0028983.g001

Hofmann et al. (2011) High-Frequency Dynamics of Ocean pH: A Multi-Ecosystem Comparison. PLoS ONE 6(12): e28983.

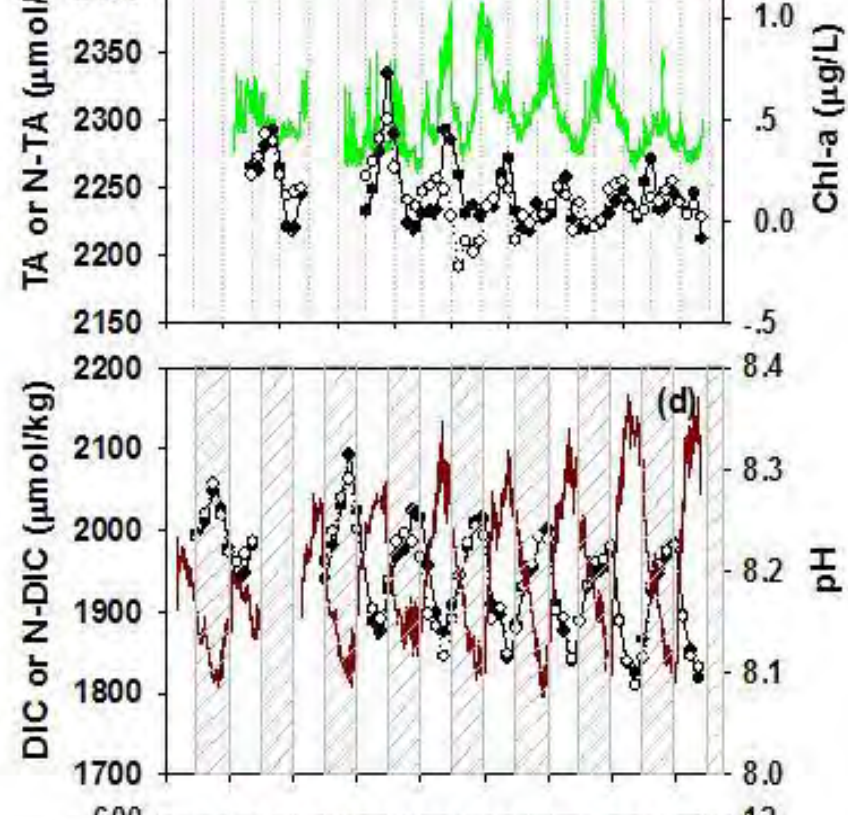
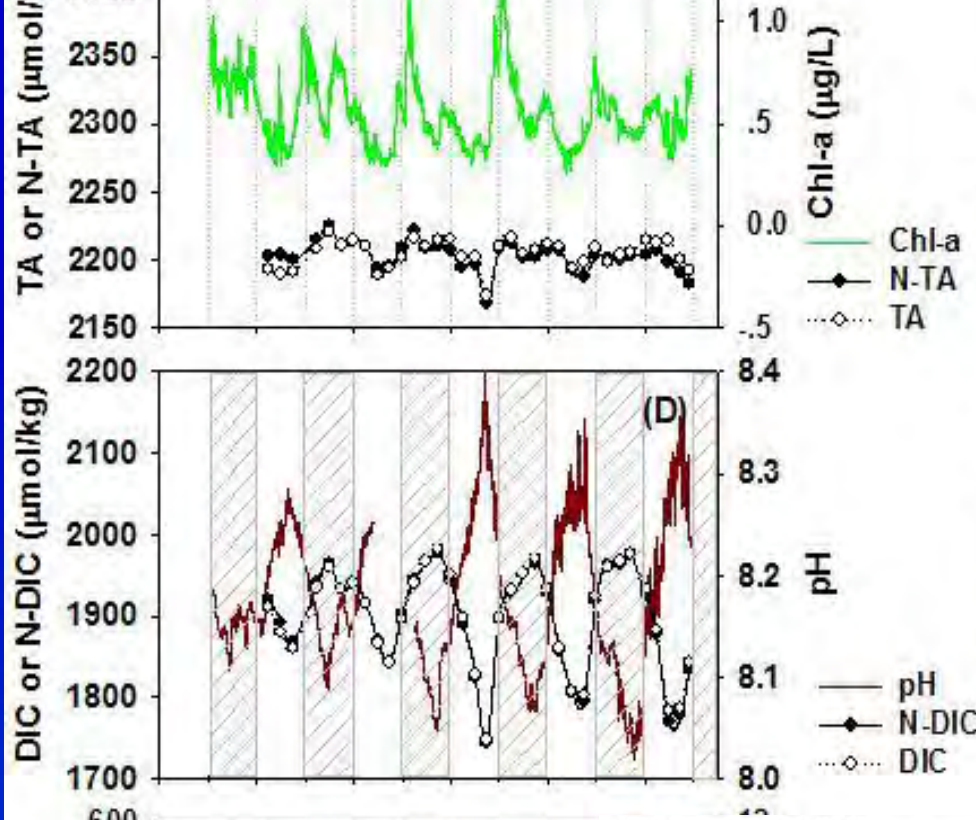
doi:10.1371/journal.pone.0028983

OA in coastal ocean more dramatic

A NE Pacific Coastal Site
(Tatoosh Island, 2000-2008)



- Wootton et al., PNAS, 2008



Limnol. Oceanogr. 56(5), 2011, 000-000
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 doi:10.4319/lo.2011.56.5.0000

Short-term dynamics of oxygen and carbon in productive nearshore shallow seawater systems off Taiwan: Observations and modeling

Zong-Pei Jiang,^{a,b} Jr-Chuan Huang,^{c,d} Minhan Dai,^{a,*} Shuh Ji Kao,^{a,c} David J. Hydes,^b Wen-Chen Chou,^e and Sen Jan^f

^aState Key Laboratory of Marine Environmental Science, Xiamen University, Xiamen, China

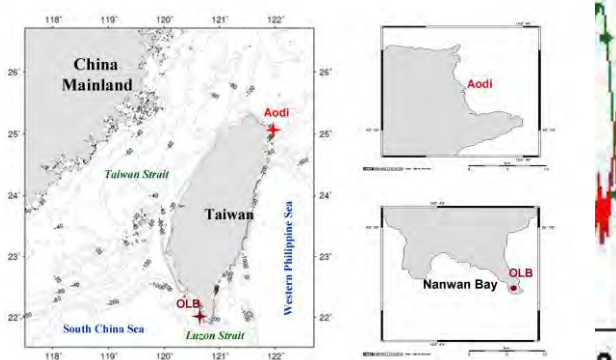
^bNational Oceanography Centre, Southampton, United Kingdom

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^eInstitute of Marine Environmental Chemistry and Ecology, National Taiwan Ocean University, Keelung

^fInstitute of Oceanography, National Taiwan University, Taipei



Spring Tide

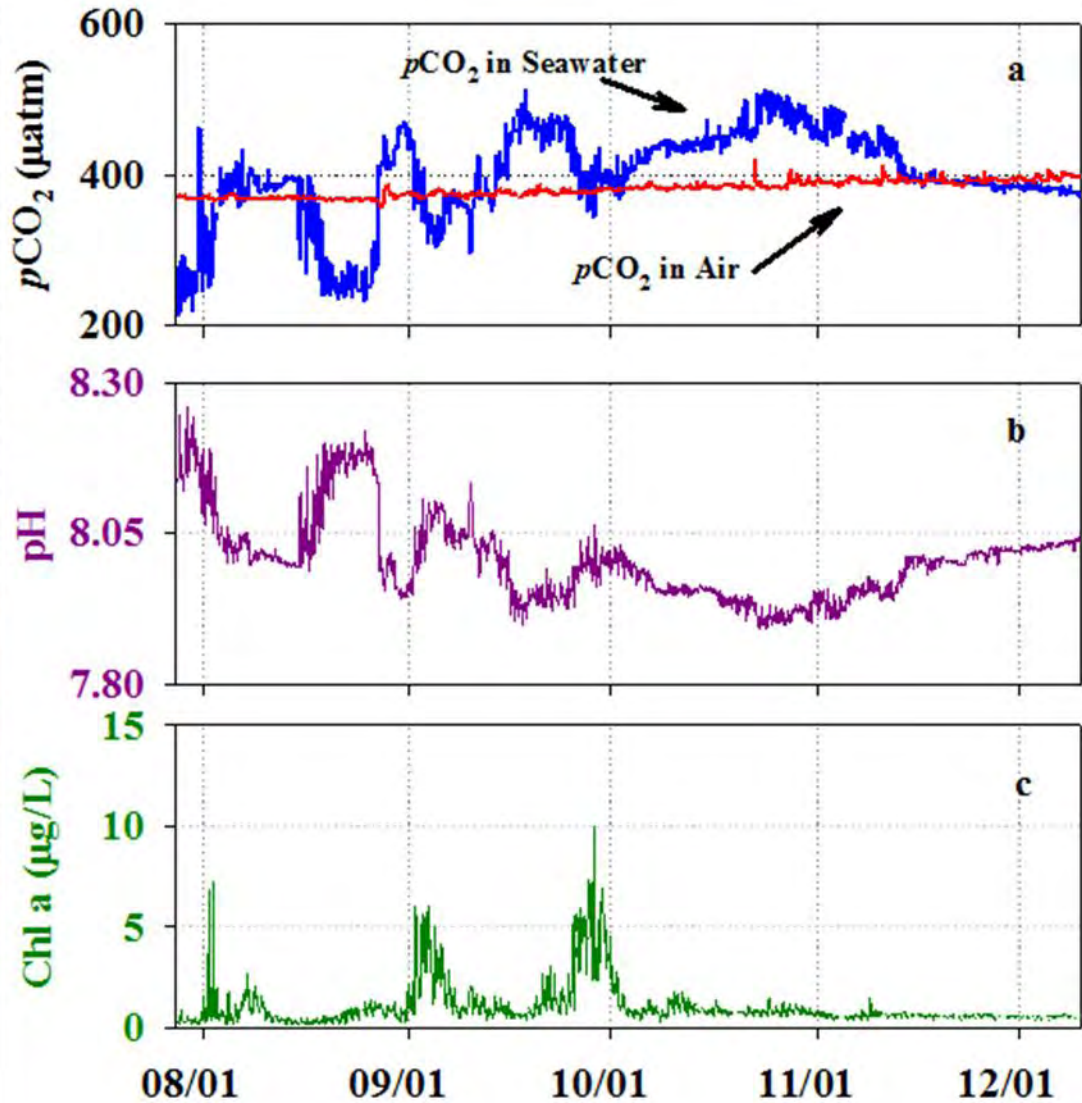
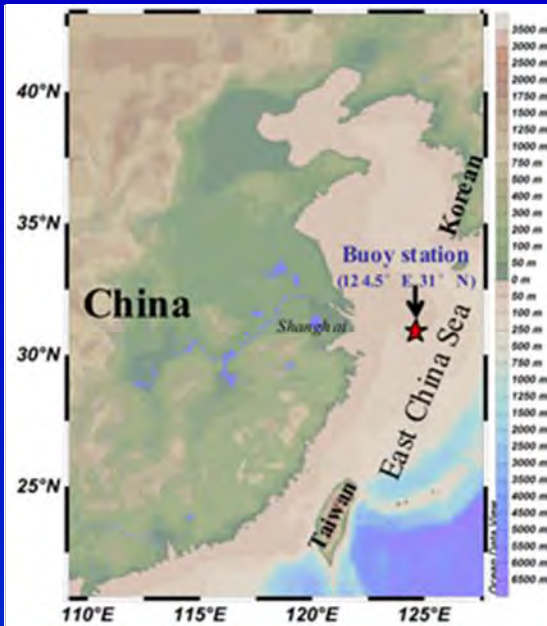
Neap Tide

Spring Tide

Neap Tide

Jiang et al., L&O, 2011

MEL-SMMC CO2-A



Cases: main drivers of OA in coastal ocean

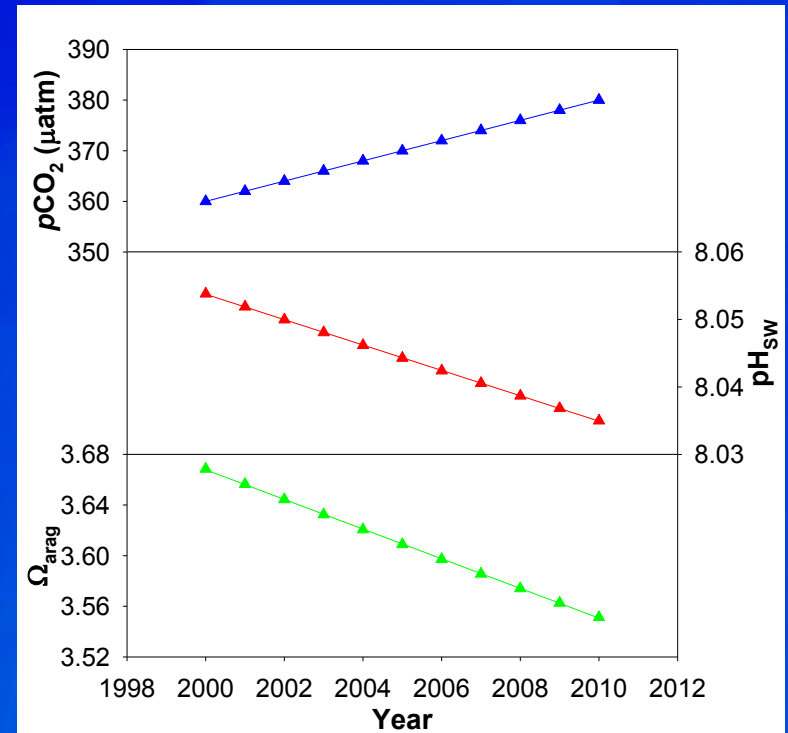
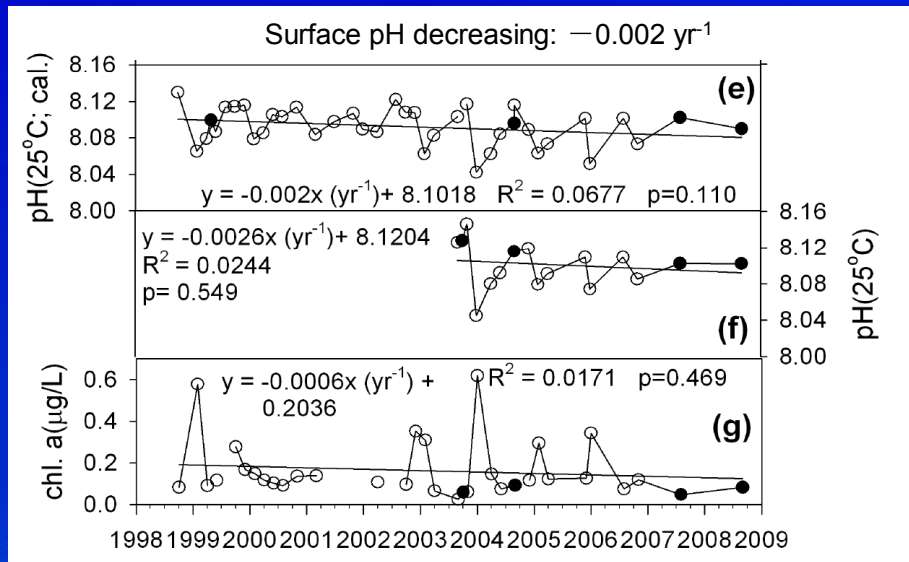
- **Anthropogenic CO₂**
- **Upwellings of low pH waters**
- **Respirations**
- **Ground water**



Anthropogenic CO₂-driven pH changes in the South China Basin

The South China Sea
(Station SEATS, 1998-2009)

Predicted anthropogenic
CO₂ induced pH decrease—
0.0019 yr⁻¹



- Arthur Chen, MEBC-SCS Conference Abstract, 2010

http://mebc-scs.marine.nsysu.edu.tw/images/Text/Abstract_Cheng.pdf

Air pCO₂ 2 μatm yr⁻¹

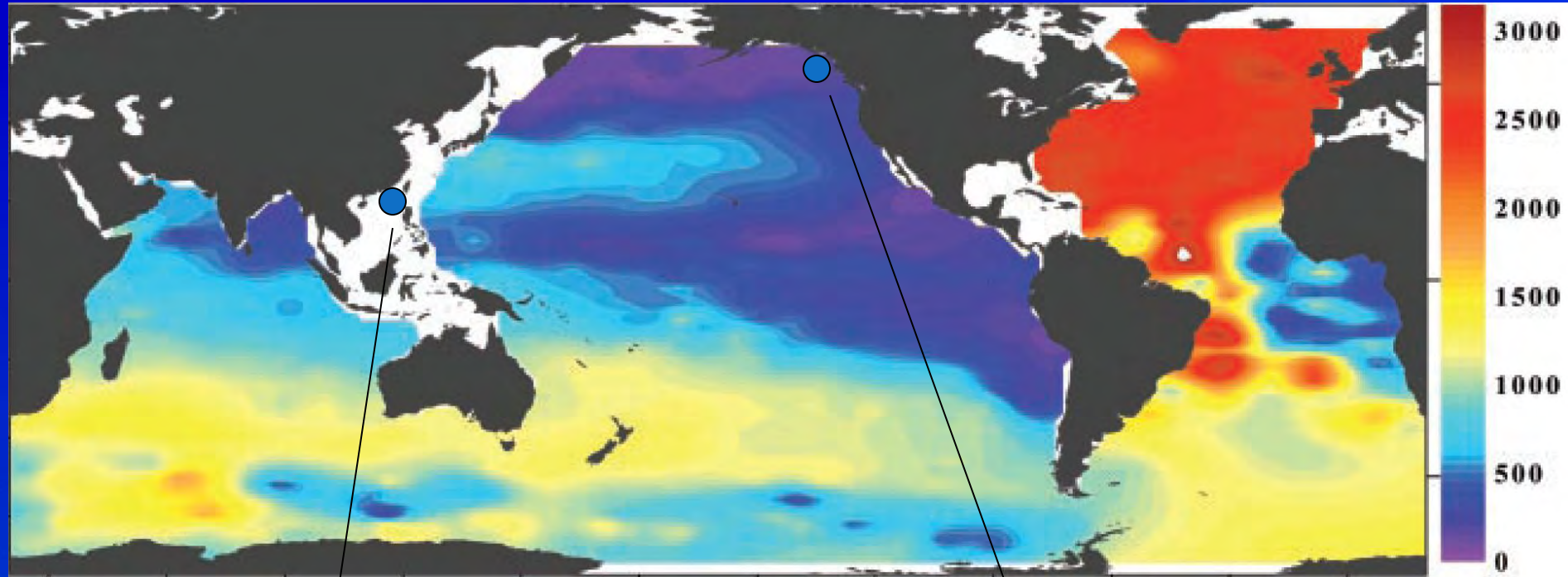
At TALK=2200 μmol kg⁻¹;
Temp=29.0 °C; S=33.5;
PO₄=0.01 μM; Si(OH)₄=2 μM

Examples: OA in coastal ocean

- Anthropogenic CO₂
- **Upwellings of low pH waters**
- Respirations
- Ground water



Aragonite Saturation Depth ($\Omega_{\text{arag}} = 1.0$) in the Global Oceans



- Feely et al., 2004

Coastal Upwelling System

Western NP vs. Eastern NP

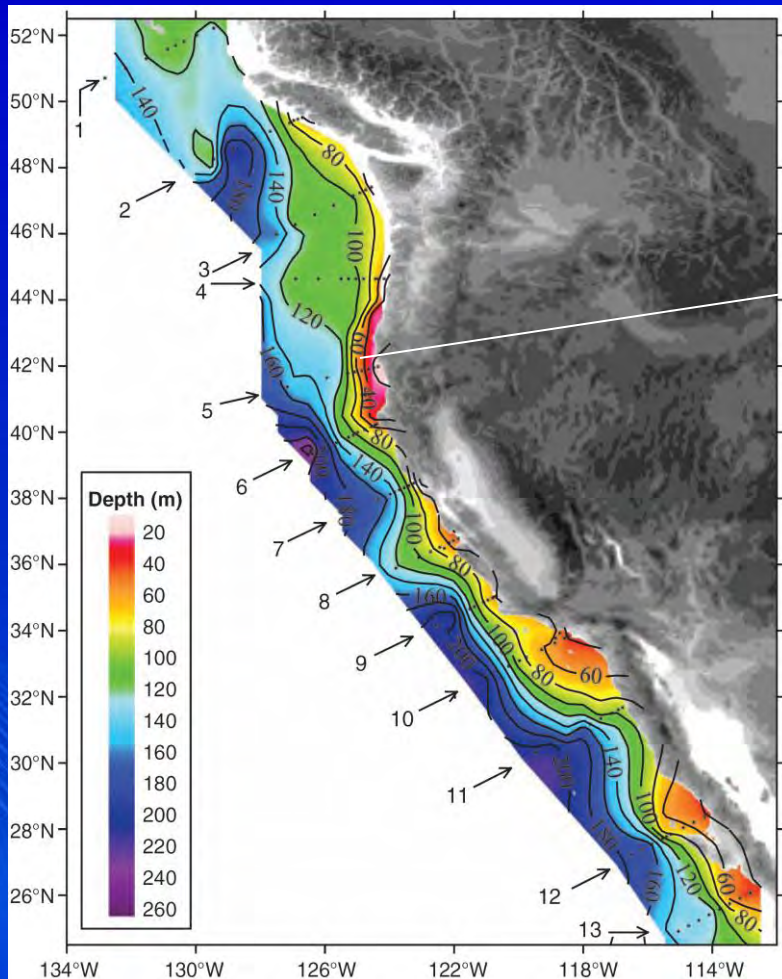
Case I: Continental shelf off the Oregon-California border (Upwelling in the Eastern NP)

Evidence for Upwelling of Corrosive “Acidified” Water onto the Continental Shelf

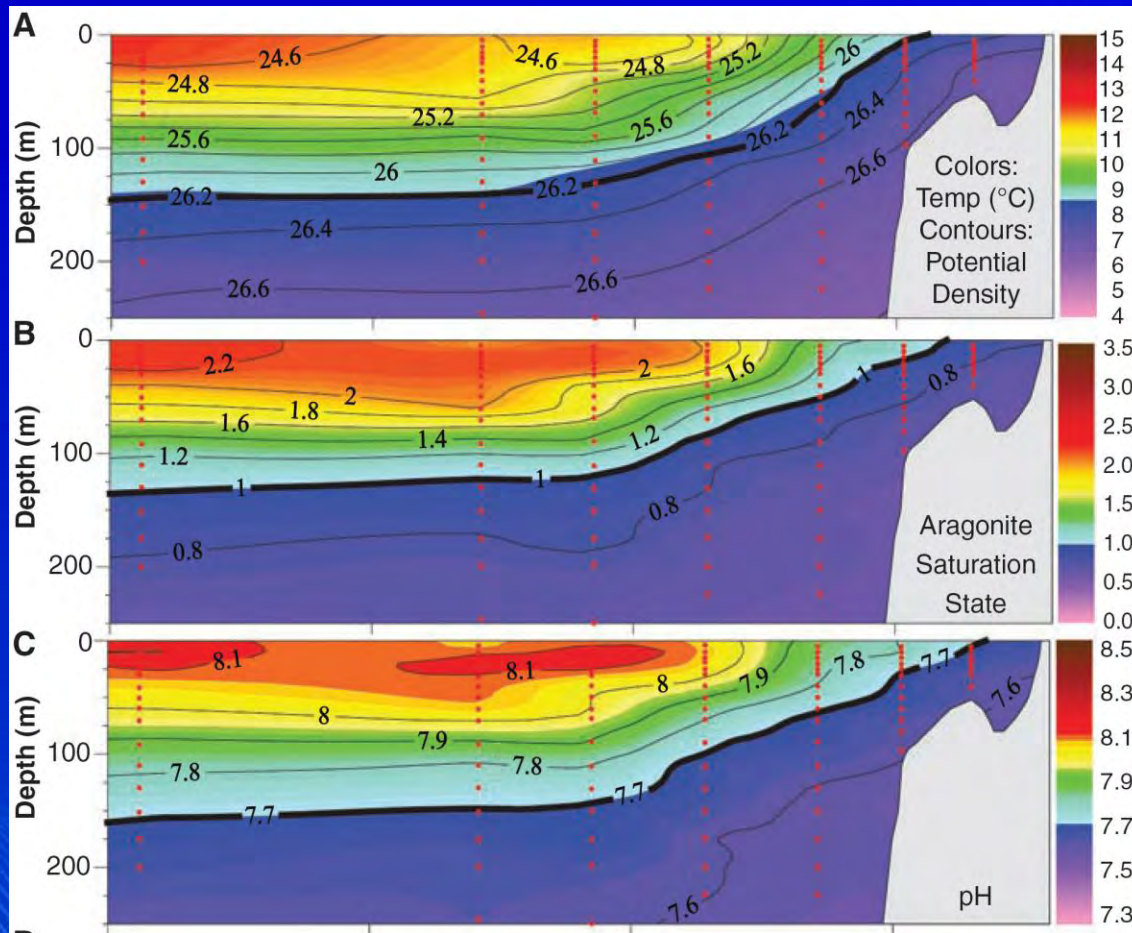
Richard A. Feely,^{1*} Christopher L. Sabine,¹ J. Martin Hernandez-Ayon,²
Debby Ianson,³ Burke Hales⁴

13 JUNE 2008 VOL 320 **SCIENCE** www.sciencemag.org

Distribution of the depths of the undersaturated water ($\Omega_{\text{arag}} < 1.0$; $\text{pH} < 7.75$) on the continental shelf of western North America.



Due to the upwelling process, the corrosive water reaches all the way to the surface in the inshore waters near the Oregon-California border.



- Upwelling from depths of 150 to 200 m onto the shelf;
- First observation of surface undersaturated waters with respect to aragonite (in open ocean surface waters until 2050);
- $\sim 30 \mu\text{mol kg}^{-1}$ of anthropogenic CO_2 lowered Ω_{arag} by ~ 0.2 units.

- Feely et al., 2008

Case II: The Northern South China Sea (NSCS) Shelf (Upwelling in the Western NP)

Dynamics of the carbonate system in a large continental shelf system under the influence of both a river plume and coastal upwelling

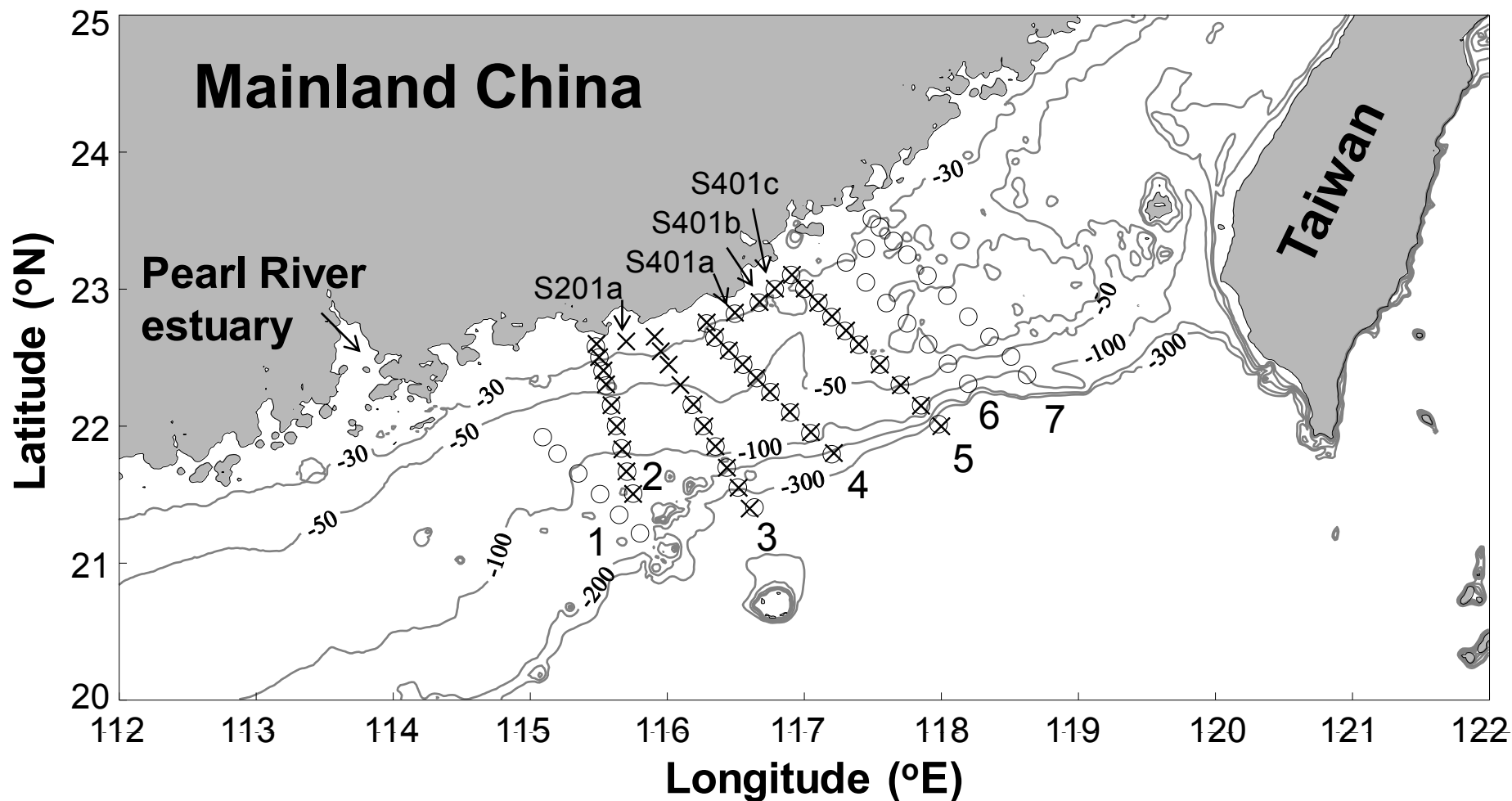
Zhimian Cao,¹ Minhan Dai,¹ Nan Zheng,¹ Deli Wang,¹ Qian Li,¹ Weidong Zhai,¹ Feifei Meng,¹ and Jianping Gan²

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 116, G02010, doi:10.1029/2010JG001596, 2011

Field work-2008 summer

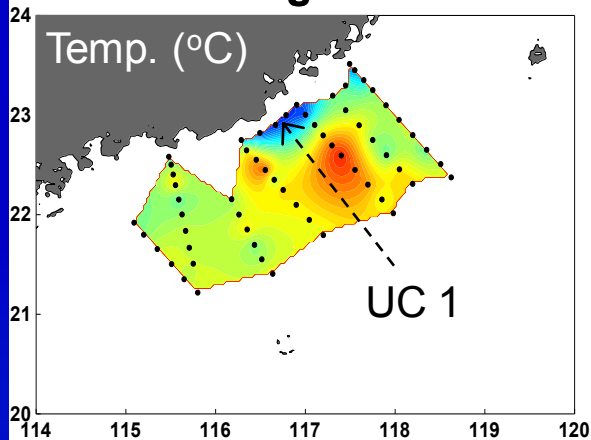
Leg 1 (Transects 1-7; Jun 30 - Jul 8)

Leg 2 (Transects 5-2; Jul 9 - Jul 12)

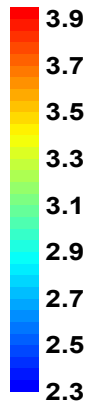
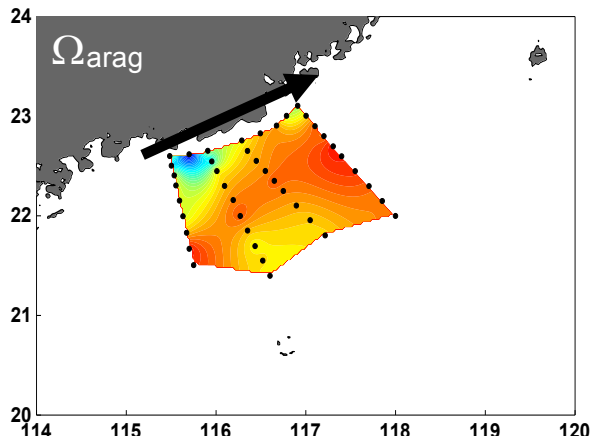
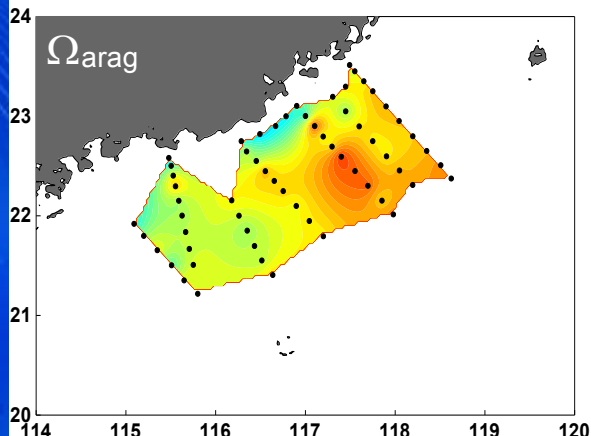
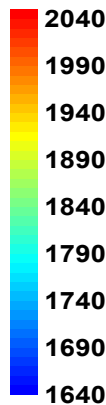
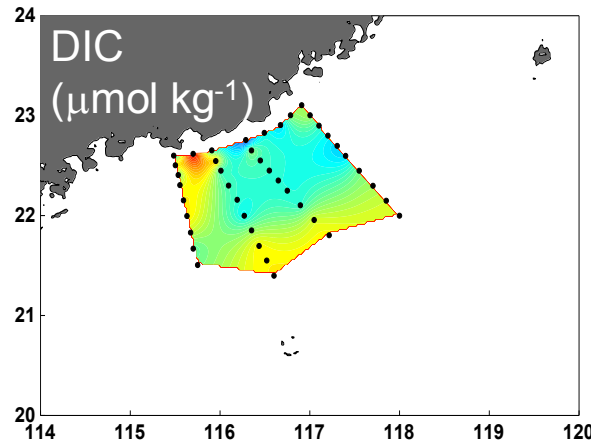
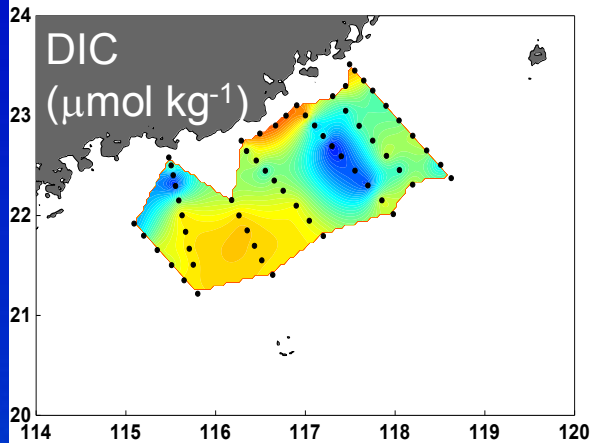
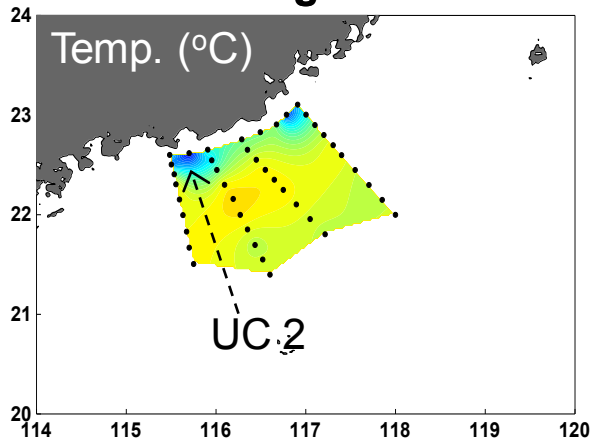


Surface

Leg 1



Leg 2



Upwelling Centers (UC):

Low Temperature

High DIC

Low Ω_{arag}

UC 2 > UC 1

Surface Ω_{arag} values:

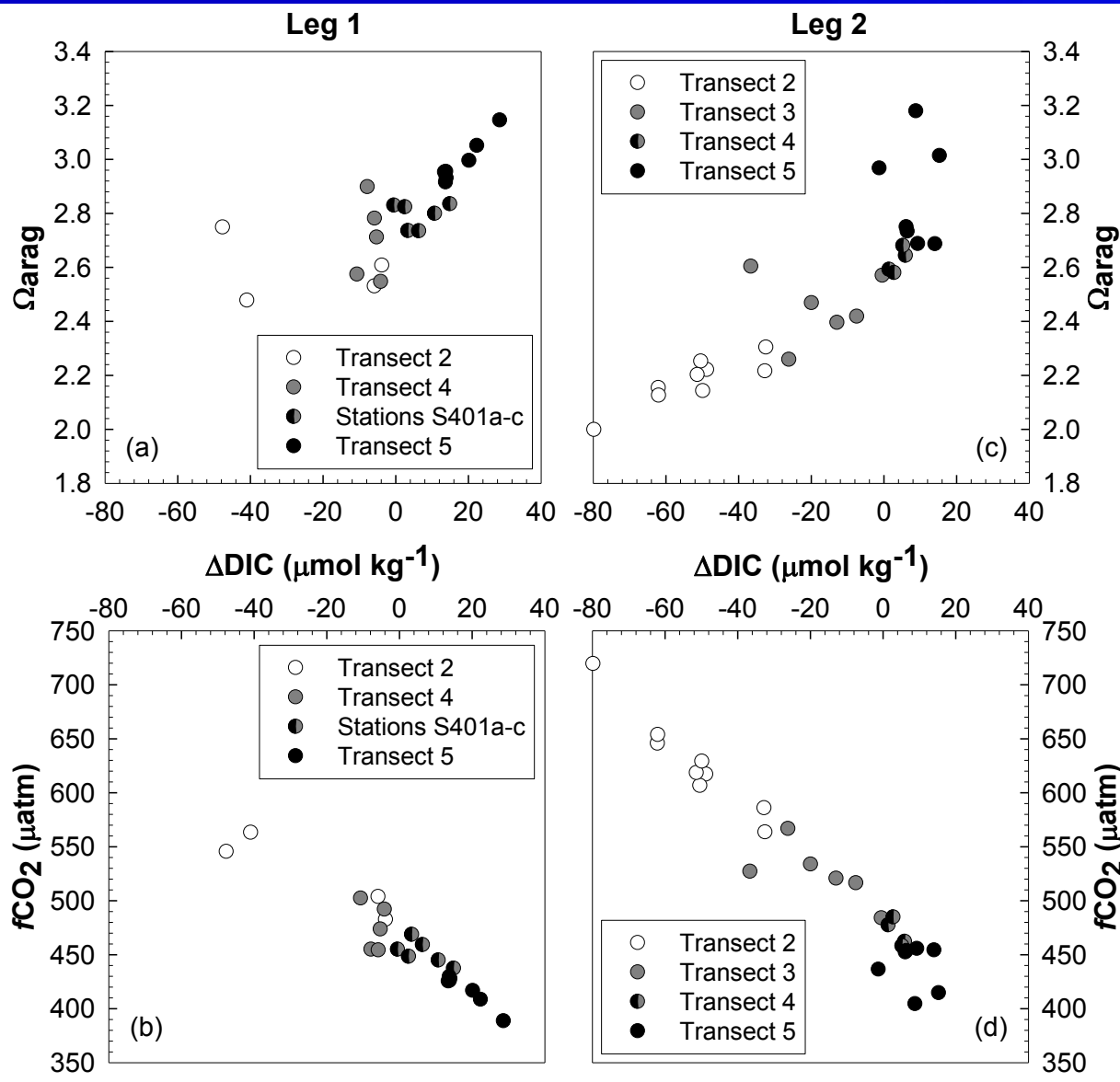
2.3 ~ 3.9

Lowest 2.3 in UC 2

Oversaturation state

Upwelled waters were transported northward from the position of UC 2 to the position of UC 1 [Gan et al., 2009].

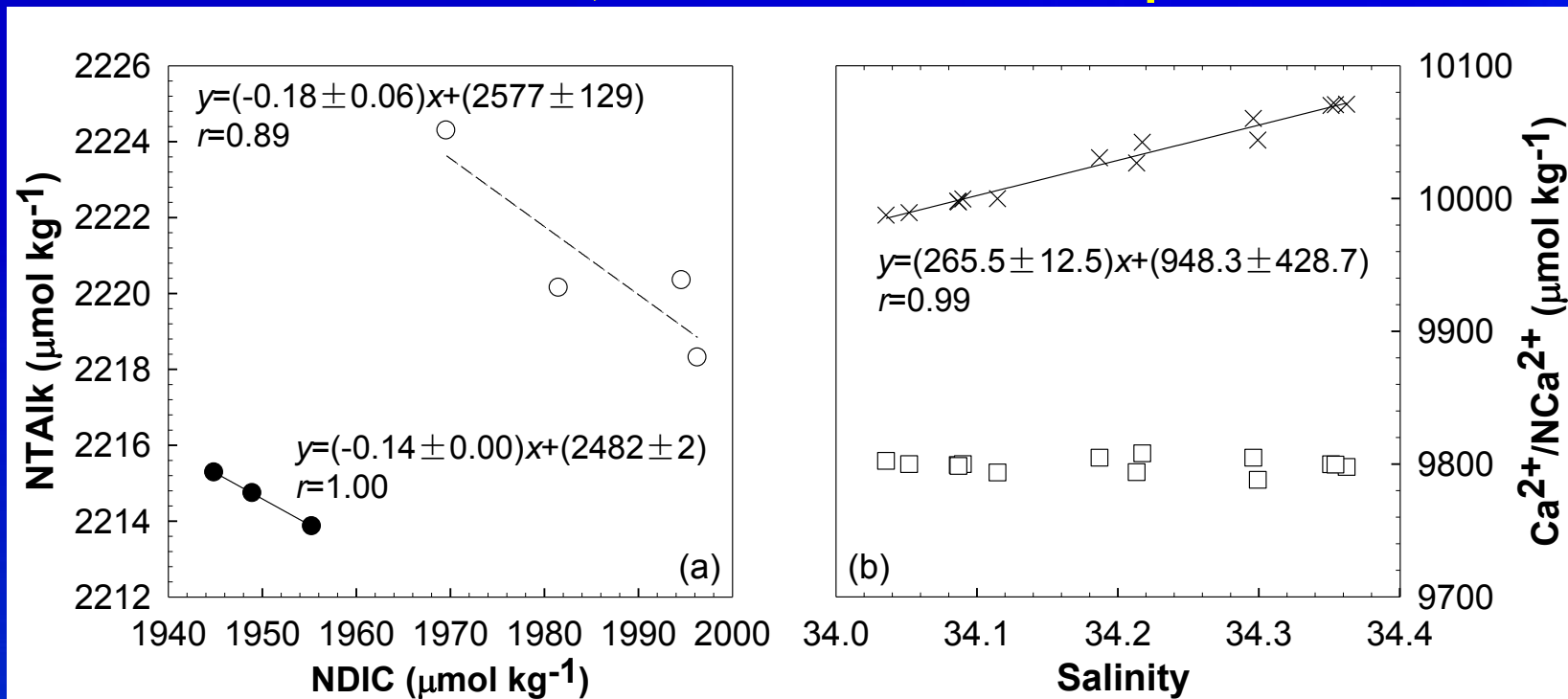
Ω_{arag} in the upwelling



Lowest values in UC 1 and UC 2 were largely imprints of upwelled subsurface waters enriched with both anthropogenic CO_2 and respired CO_2 ;

During the northward transport of the upwelled water, however, Ω_{arag} increased whereas $f\text{CO}_2$ decreased with increasing ΔDIC . The consistent biological production gradually consumed CO_2 and raised $[\text{CO}_3]$, resulting in the increasing Ω_{arag} pattern in the upwelled water.

Behaviors of DIC, TAlk and Ca²⁺ in the upwelled water



- NDIC variations $\sim 30 \mu\text{mol kg}^{-1}$; NTAlk variations $\sim 6 \mu\text{mol kg}^{-1}$
- Slopes of NTAlk against NDIC were close to -0.16
 OC metabolism: $106\text{CO}_2 + 122\text{H}_2\text{O} + 16\text{HNO}_3 + \text{H}_3\text{PO}_4 \leftrightarrow (\text{CH}_2\text{O})_{106}(\text{NH}_3)_{16}\text{H}_3\text{PO}_4 + 138\text{O}_2$, $\Delta\text{TAlk}/\Delta\text{DIC} = -0.16$
- IC metabolism: $\text{Ca}^{2+} + 2\text{HCO}_3^- \leftrightarrow \text{CaCO}_3 + \text{CO}_2 + \text{H}_2\text{O}$, $\Delta\text{TAlk}/\Delta\text{DIC} = 2$
- Conservative behavior of Ca^{2+}



No significant impact of the upwelling on the net community calcification!

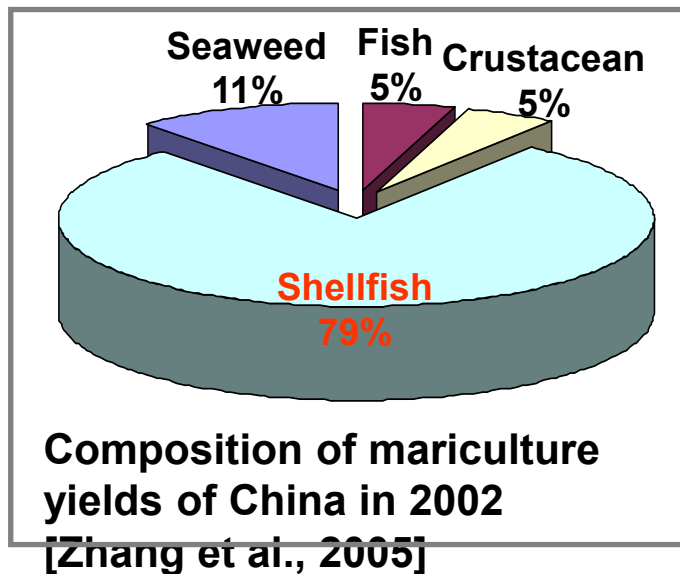
Summary for the NSCS upwelling

- Upwelling delivered low Ω_{arag} waters onto the shelf, but Ω_{arag} values are still > 1.0 ;
- Enhanced OC production during the upwelling further increased Ω_{arag} ;
- The patterns of Ω_{arag} are largely controlled by water circulation and biological activities at present;
- Oversaturation for now.



Context

- Corrosive waters with $\Omega < 1.0$ may be possible on most of wide shelves including the NSCS shelf!
- How the calcifying organisms, including many species of shellfish of economic importance to coastal regions, deal with this intermittent exposure of corrosive waters?

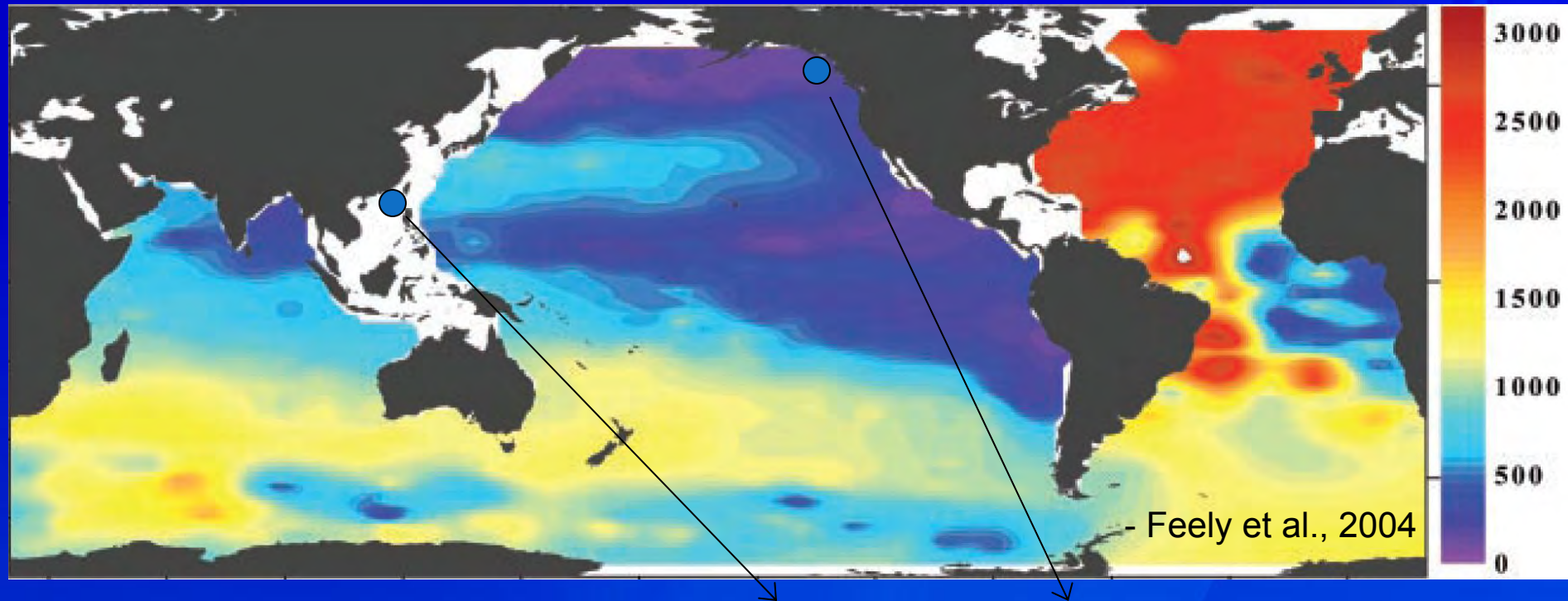


China is the largest mariculture country, contributing to 68.81% of the total global yields (FAO, 2006);

Shellfish is the most important composition.

Production of shell fish in 2008 was ~ 10 MT

Aragonite Saturation Depth ($\Omega_{\text{arag}} = 1.0$) in the Global Oceans



Source Water Depth	~150 m	~150 m
Source Water pH	~7.95	< 7.75
Source Water Ω_{arag}	~1.9	< 1.0
Aragonite Saturation Depth	~500 m in the Western NP	100-300 m in the Eastern NP

The contrasting scenarios between the two cases are related to the different “acidity” of source waters of upwelling.

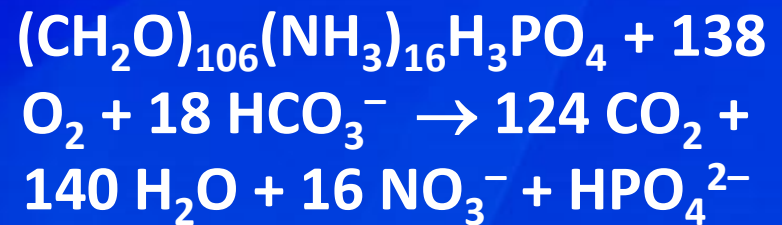
Examples: OA in coastal ocean

- Anthropogenic CO₂
- Upwellings of low pH waters
- **Respirations/nitrification**
- Ground water



Respiration + Nitrification induced pH drawdown in the Pearl River

Aerobic respiration



Nitrification

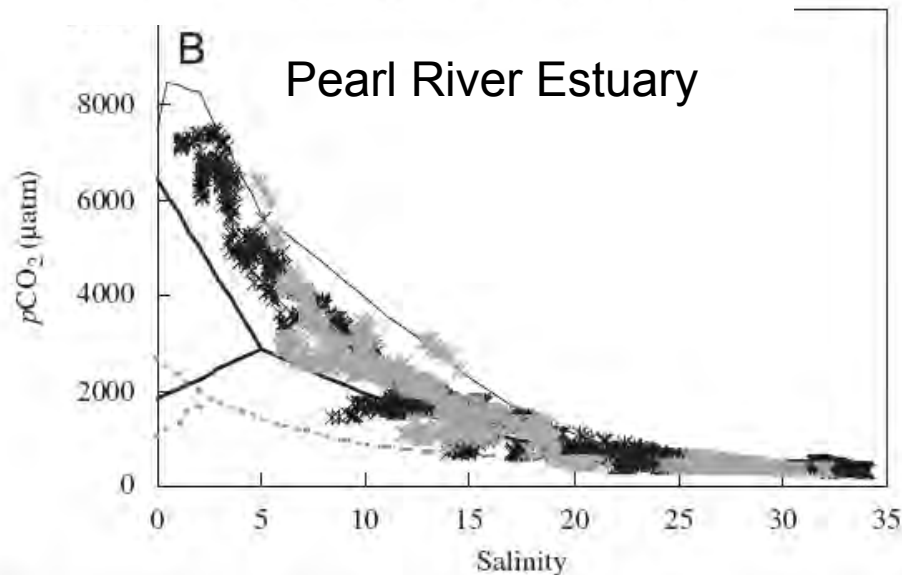
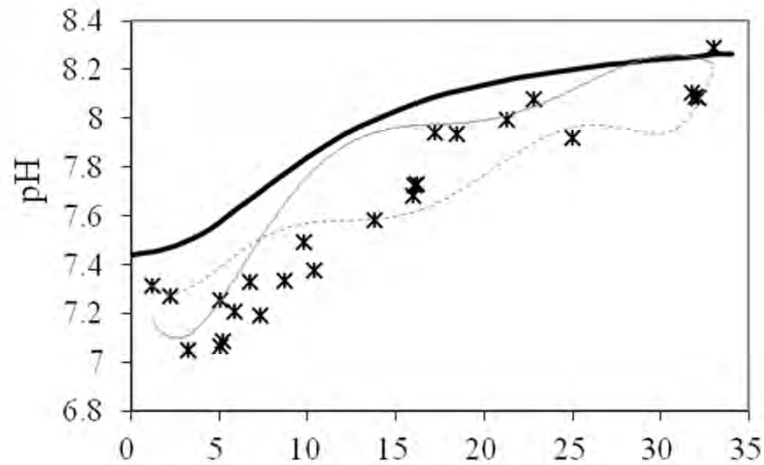
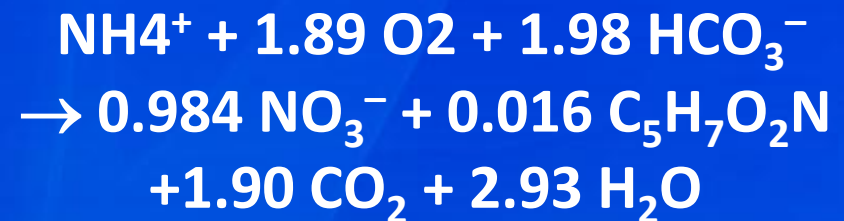


Fig. 10. A mechanistic explanation of the effect of nitrification on the formation of pH minimum (A) and $p\text{CO}_2$ maximum (B) during the dry season. For the dry season simulation, the solid black lines are the result of conservative mixing among the three end-members; the gray solid lines (same symbols as in Fig. 8) represent those with the effect of nitrification in dry season as explained in the text; the black and gray "x" symbols represent February 2004 and January 2005 data, respectively. For the wet season simulation, only conservative mixing (dashed gray lines) is presented since nitrification is not significant in the mixing zone.

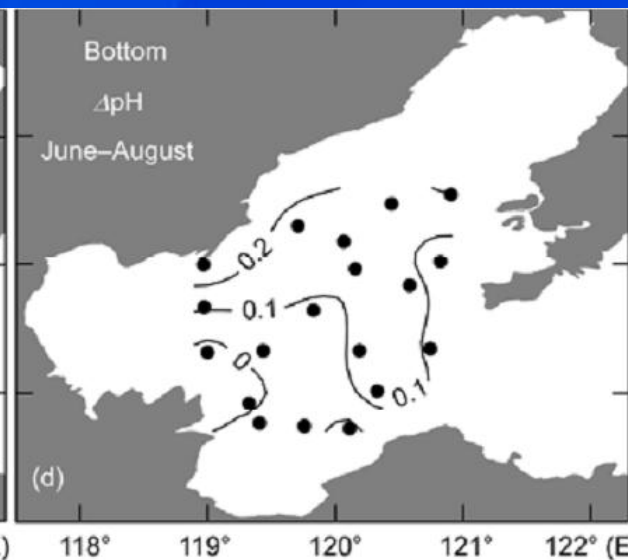
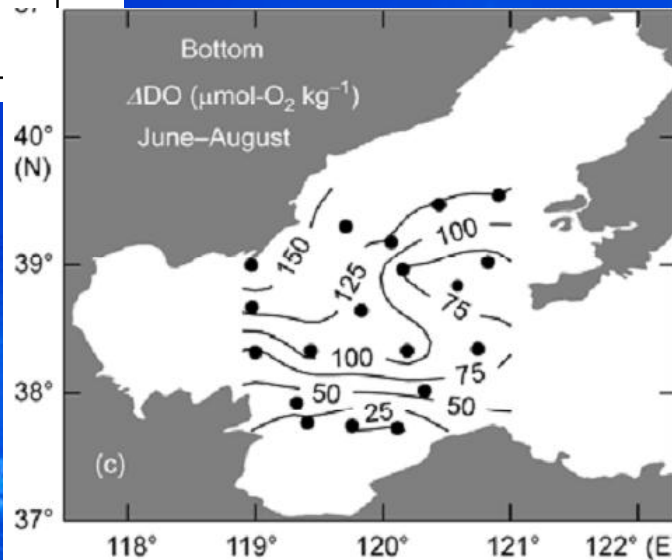
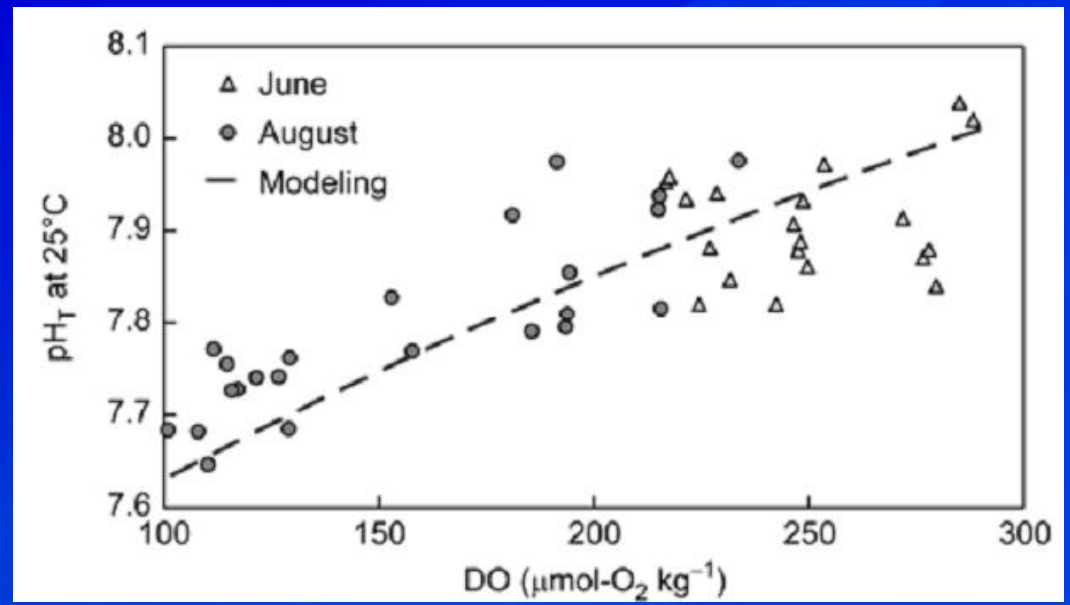
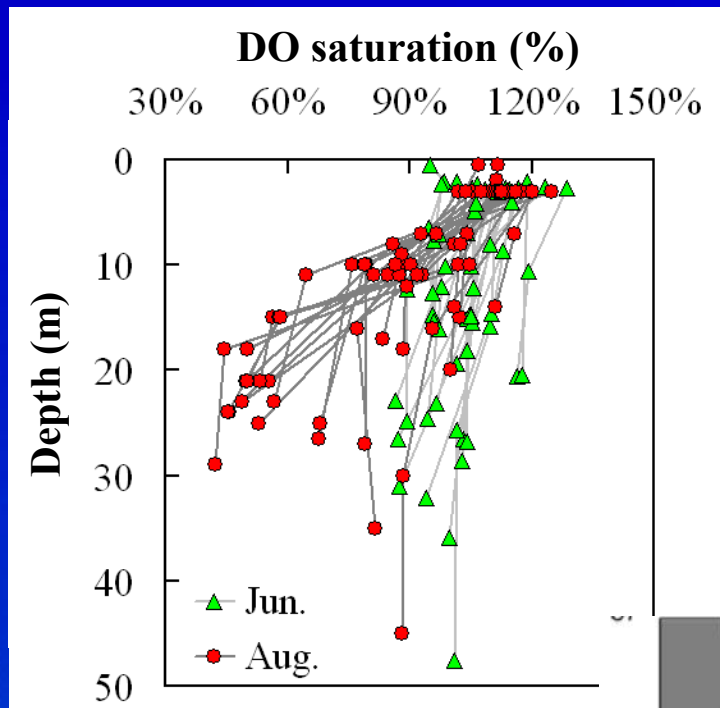
Fig A from Zhang Cao

Fig B from Guo et al., 2008, CSR

Coastal acidification in summer bottom oxygen-depleted waters in northwestern-northern Bohai Sea from June to August in 2011

ZHAI WeiDong^{1,2*}, ZHAO HuaDe^{1,2}, ZHENG Nan¹ & XU Yi²

Chin Sci Bull, 2012, 57: 1062–1068, doi: 10.1007/s11434-011-4949-2



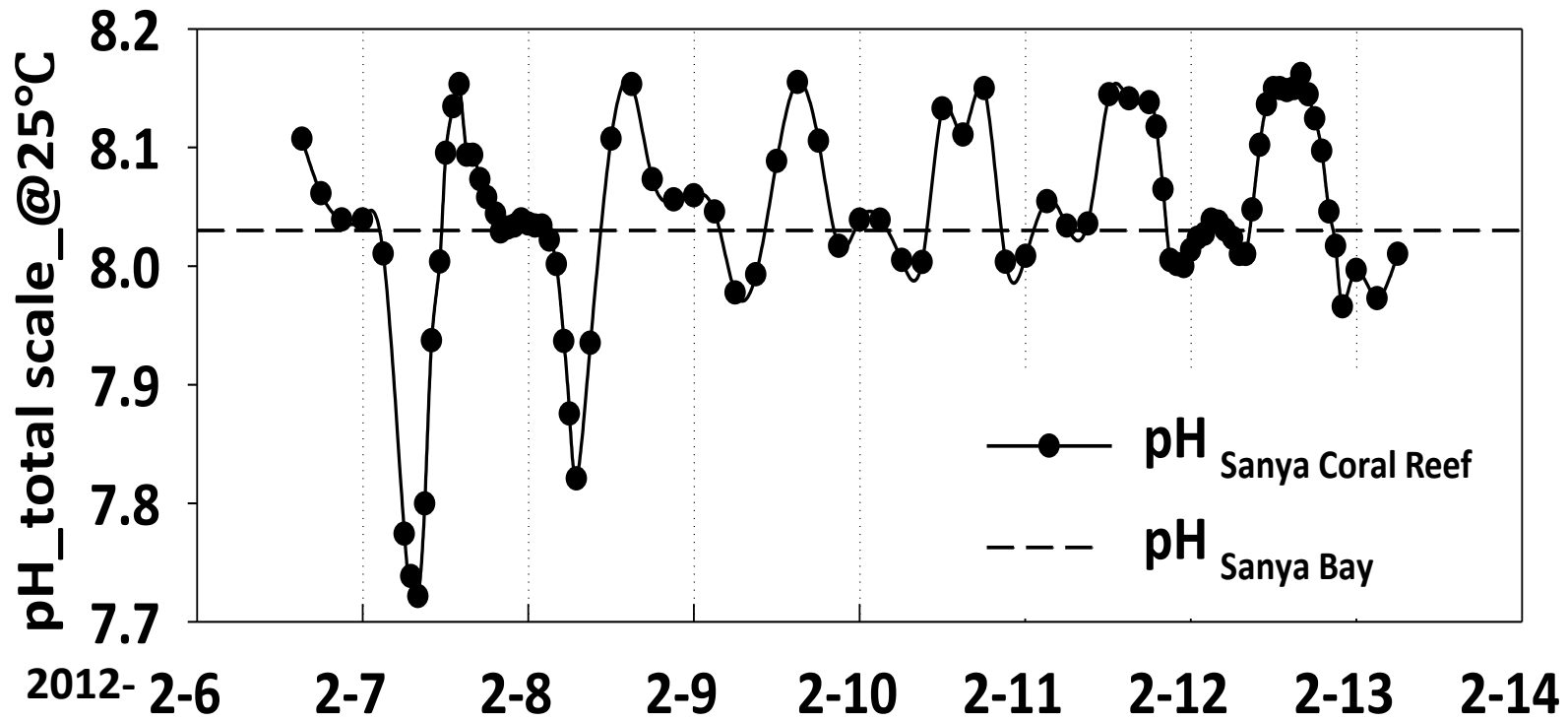
Examples: OA in coastal ocean

- Anthropogenic CO₂
- Upwellings of low pH waters
- Respirations
- **Ground water**



Acidification in Hainan Coral Reef

Low pH in Sanya fringing reef system



Major processes related to OA in the coastal ocean

Coastal ocean has large natural variability due to changes in physical and biological conditions and various anthropogenic impacts including:

River Plume

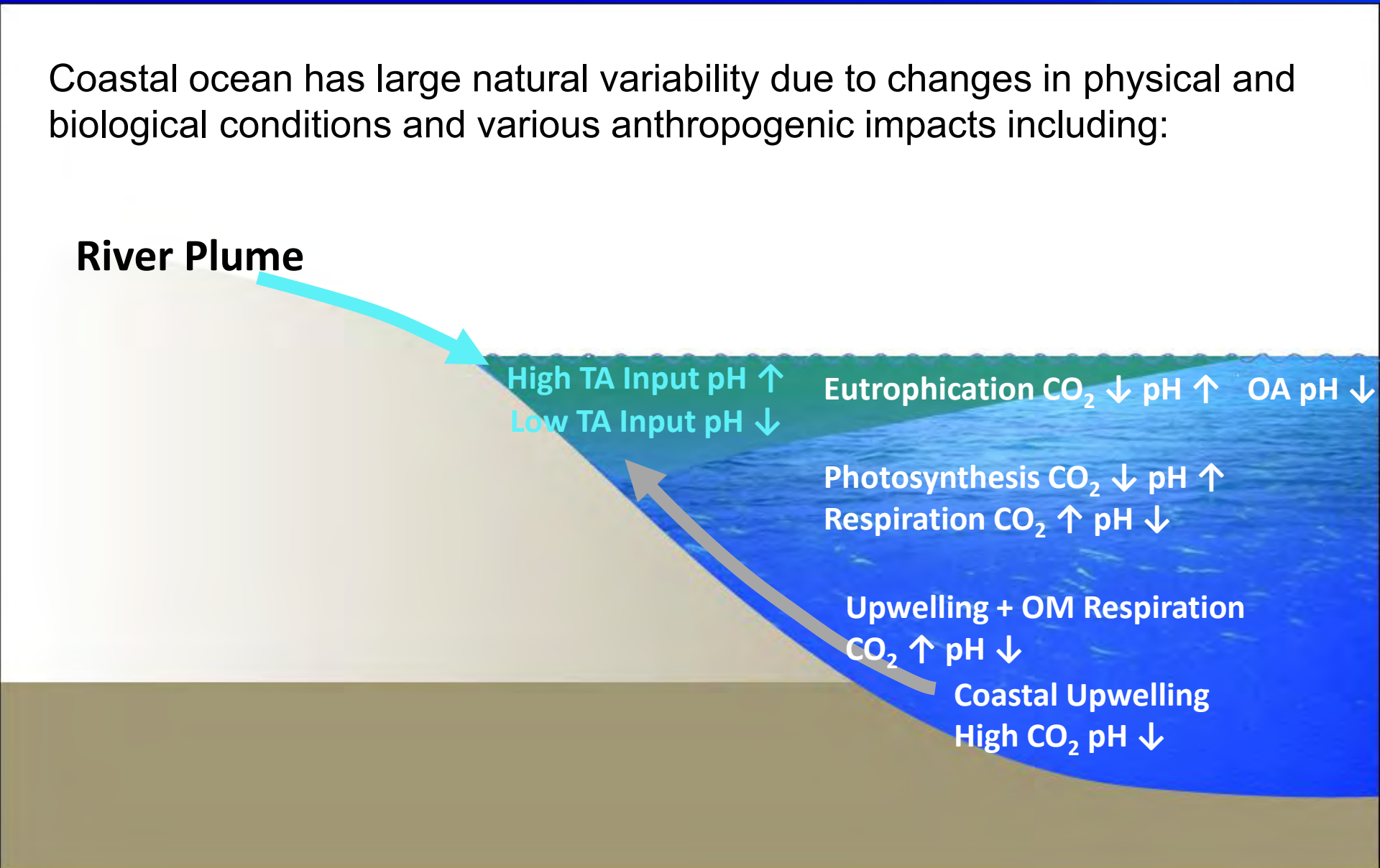
High TA Input pH \uparrow
Low TA Input pH \downarrow

Eutrophication $\text{CO}_2 \downarrow$ pH \uparrow OA pH \downarrow

Photosynthesis $\text{CO}_2 \downarrow$ pH \uparrow
Respiration $\text{CO}_2 \uparrow$ pH \downarrow

Upwelling + OM Respiration
 $\text{CO}_2 \uparrow$ pH \downarrow

Coastal Upwelling
High CO_2 pH \downarrow



Basics about coastal OA

- **OA in coastal ocean?** pH change in coastal ocean may be more complex, driven by anthropogenic CO_2 , riverine/groundwater input (eutrophication & respiration) and water circulation (upwelling of low pH/DO waters), and other biogeochemical reactions under low/no DO.
- Physical dynamics must be considered.
- OA observation is essential to set the stage of OA in coastal ocean
- Coastal ocean is particularly facing multiple stressors

Outline

- Coastal Ocean Acidification
- **Multiple stressors in the Coastal Ocean**
- Concluding Remarks



Multiple stressors in marine ecosystems

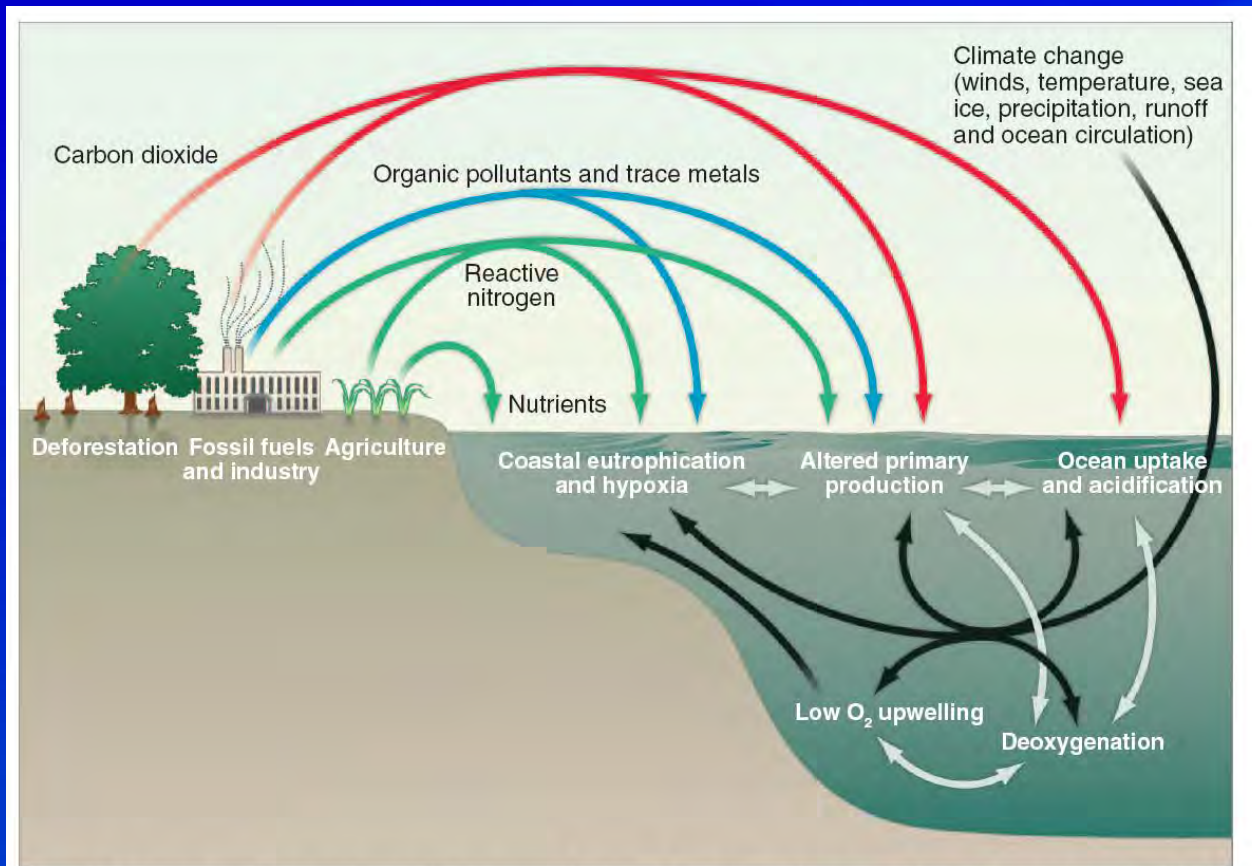


Fig. 1. Schematic of human impacts on ocean biogeochemistry either directly via fluxes of material into the ocean (colored arrows) or indirectly via climate change and altered ocean circulation (black arrows). The gray arrows denote the interconnections among ocean biogeochemical dynamics. Note that many ocean processes are affected by multiple stressors, and the synergistic effects of human perturbations is a key area for further research.

Multiple forcings & responses: offset? Amplifying?

- Eutrophication-hypoxia + OA +temp?
- SGD + upwelling +....

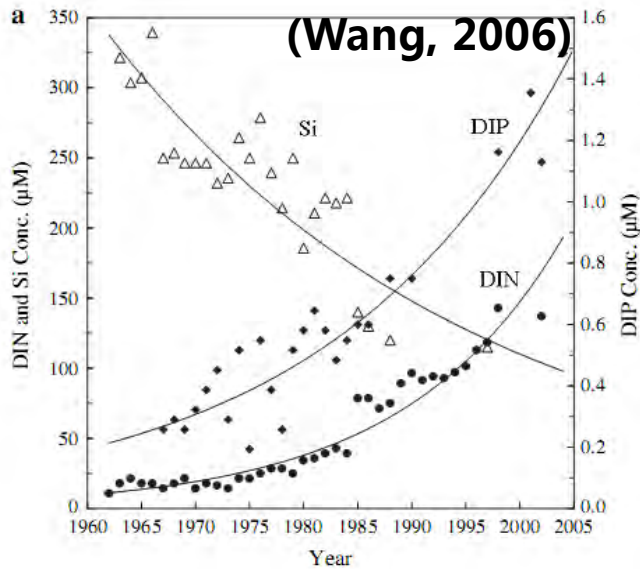


Multiple stressors: case 1

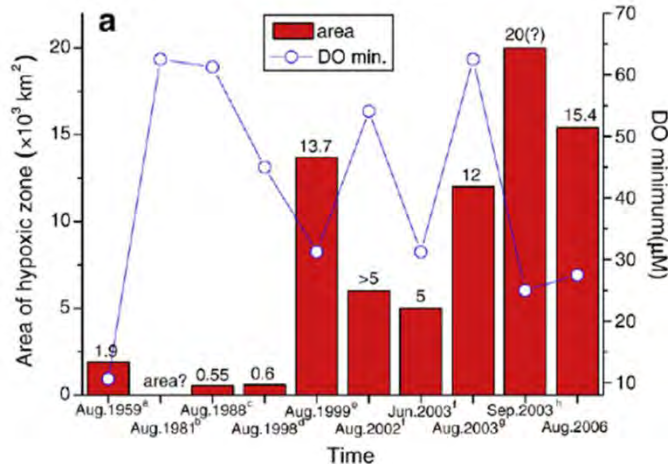
Enhanced OA by hypoxia and effect
of temp?



Hypoxia in the East China Sea off the Changjiang (Yangtze River) Estuary



Li & Zhang, 2002



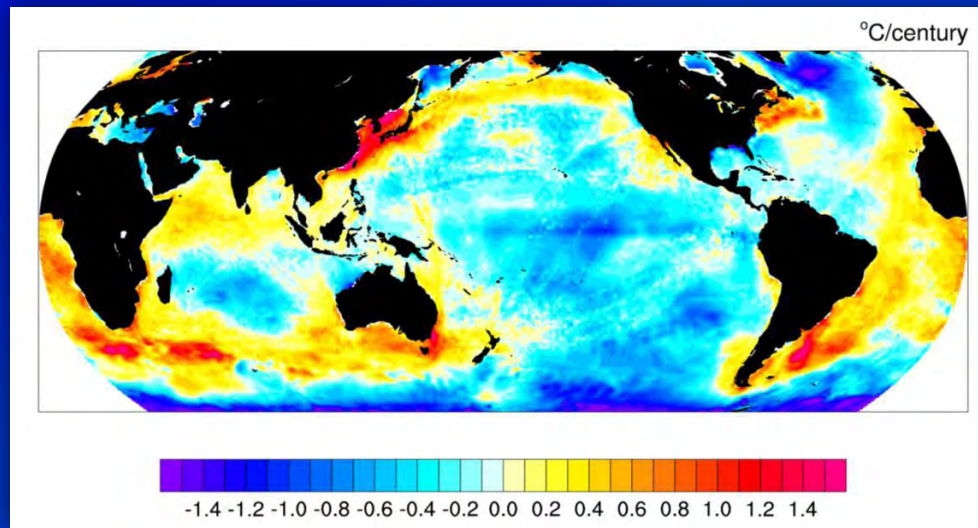
$\text{DO} < 2 \text{ mg/L}$, $\sim 10000 - 20000 \text{ km}^2$
 Taiwan: $\sim 32,260 \text{ km}^2$

(Zhu et al., 2011)

Enhanced warming over the global subtropical western boundary currents

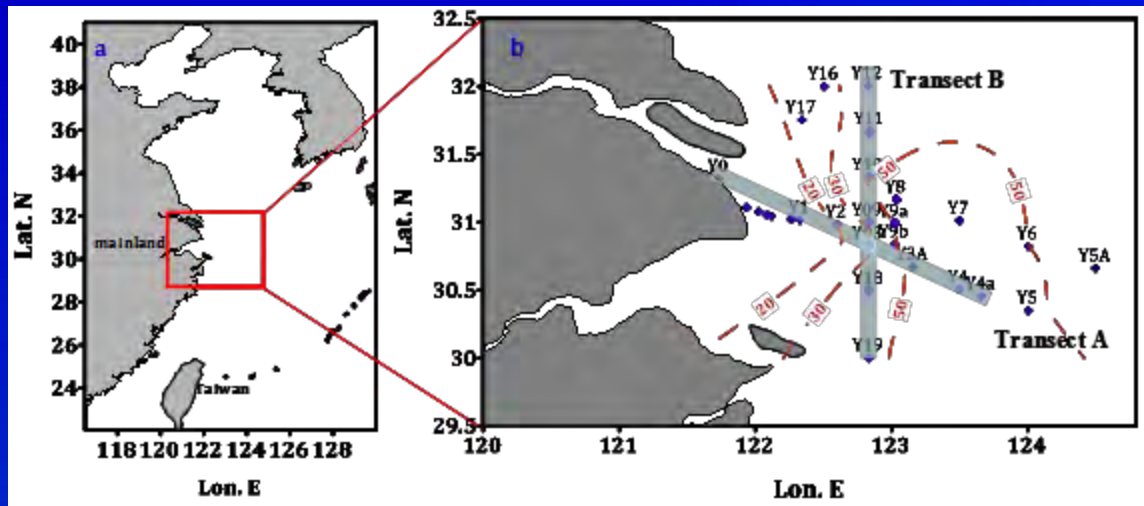
Lixin Wu^{1*}, Wenju Cai², Liping Zhang¹, Hisashi Nakamura³, Axel Timmermann⁴, Terry Joyce⁵, Michael J. McPhaden⁶, Michael Alexander⁷, Bo Qiu⁴, Martin Visbeck⁸, Ping Chang⁹ and Benjamin Giese⁹

SST anomaly of global ocean in 20th century



The enhanced warming over the global subtropical western boundary currents in the 20th century might be attributable to the poleward shift of their mid-latitude extensions and/or intensification in their strength.

Enhanced OA by respiration in hypoxic zone off the Changjiang estuary



All data points with $S > 31$ and depth > 10 m are selected.

Survey in Aug. 2011

The offshore water:

$S = 33.2716$

$T = 26.7065$

$P = 10.471$ dbar

$DO = 208.31$ $\mu\text{mol/kg}$

$DIC = 1937.25$ $\mu\text{mol/kg}$

$TA = 2227.98$ $\mu\text{mol/kg}$

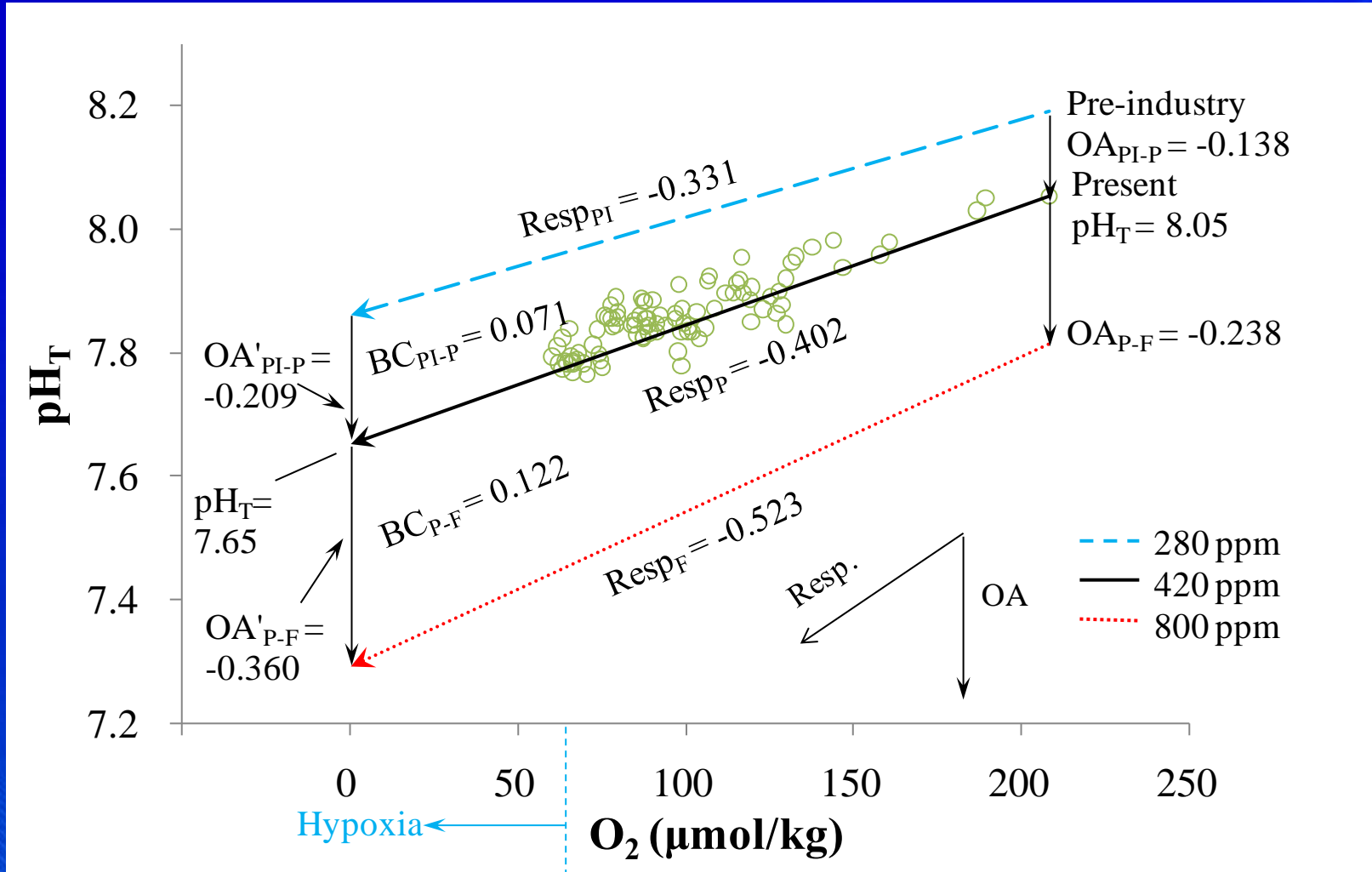
$NO_3 + NO_2 = 0$ $\mu\text{mol/kg}$

Silicate = 0 $\mu\text{mol/kg}$

Phosphate = 0 $\mu\text{mol/kg}$

$pCO_{2_Calculated} = 420$ μatm

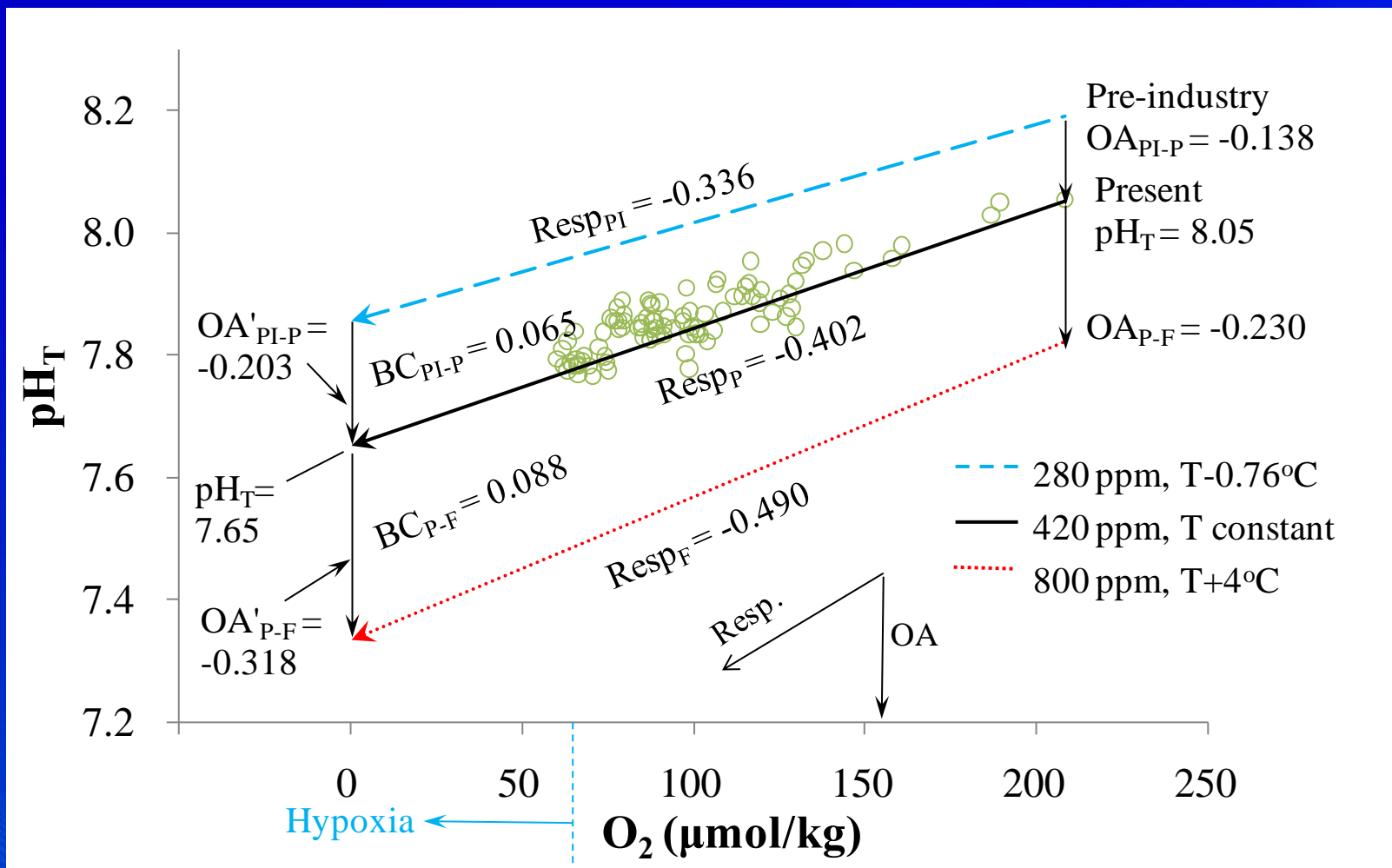
Model 1: S, T, P, and Nutrients Constant



Considering the slight TA change during respiration by $\text{TA}:\text{O} = -17/138$; $\text{DIC}:\text{DO} = 106/138$

BC: Buffering Capacity; PI: preindustrial pCO_2 (280 ppm); P: present pCO_2 (420 ppm); F: future pCO_2 (year 2100, 800 ppm)

Model 2: S, P, Nutrients Constant, T Changes



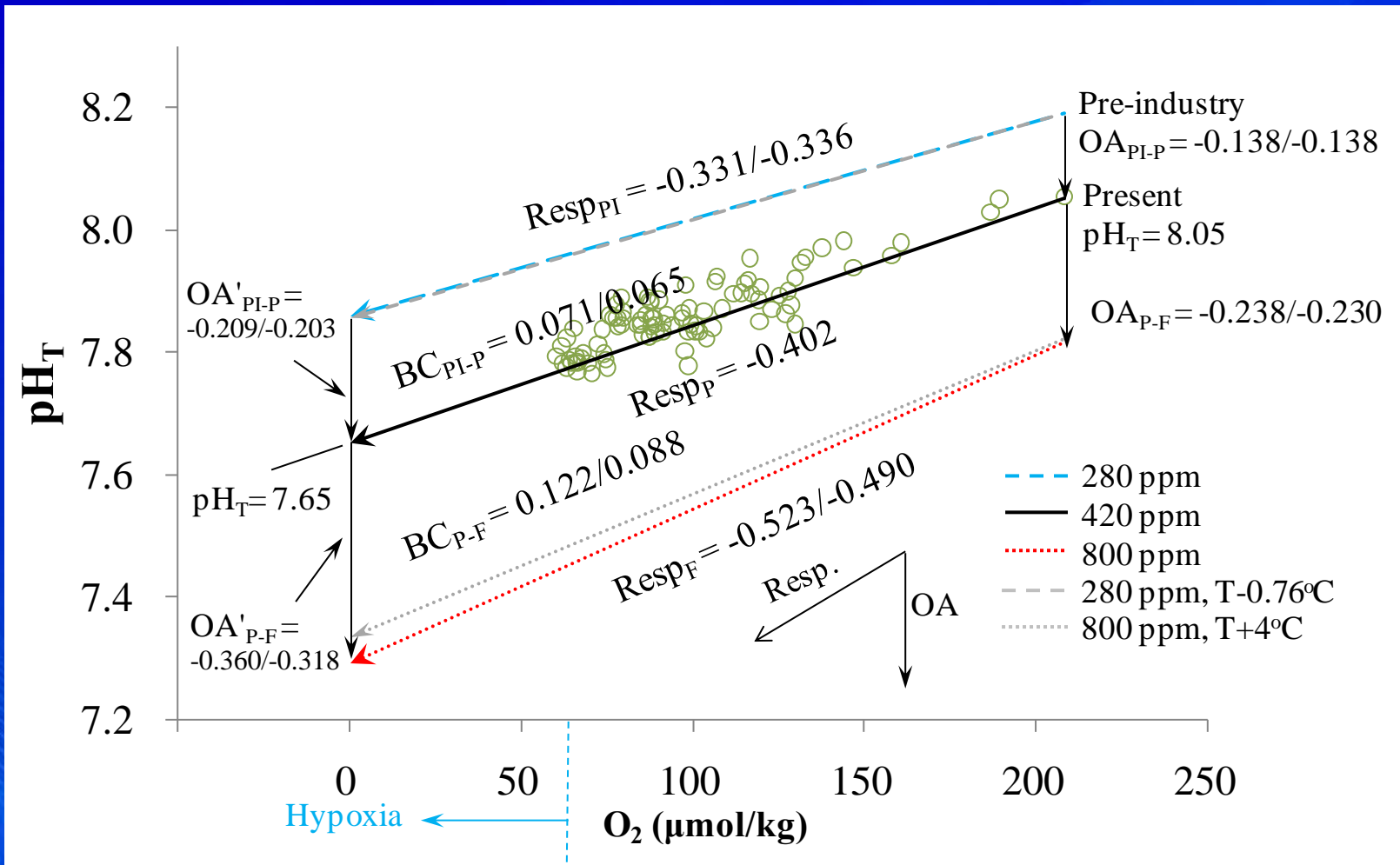
Consider the slight TA change during respiration by $\text{TA}:\text{O} = -17/138$; $\text{DIC}:\text{O} = 106/138$

Consider the temperature change but no effect of temperature on respiration

$T_{\text{preindustry}}: T-0.76^\circ\text{C}$

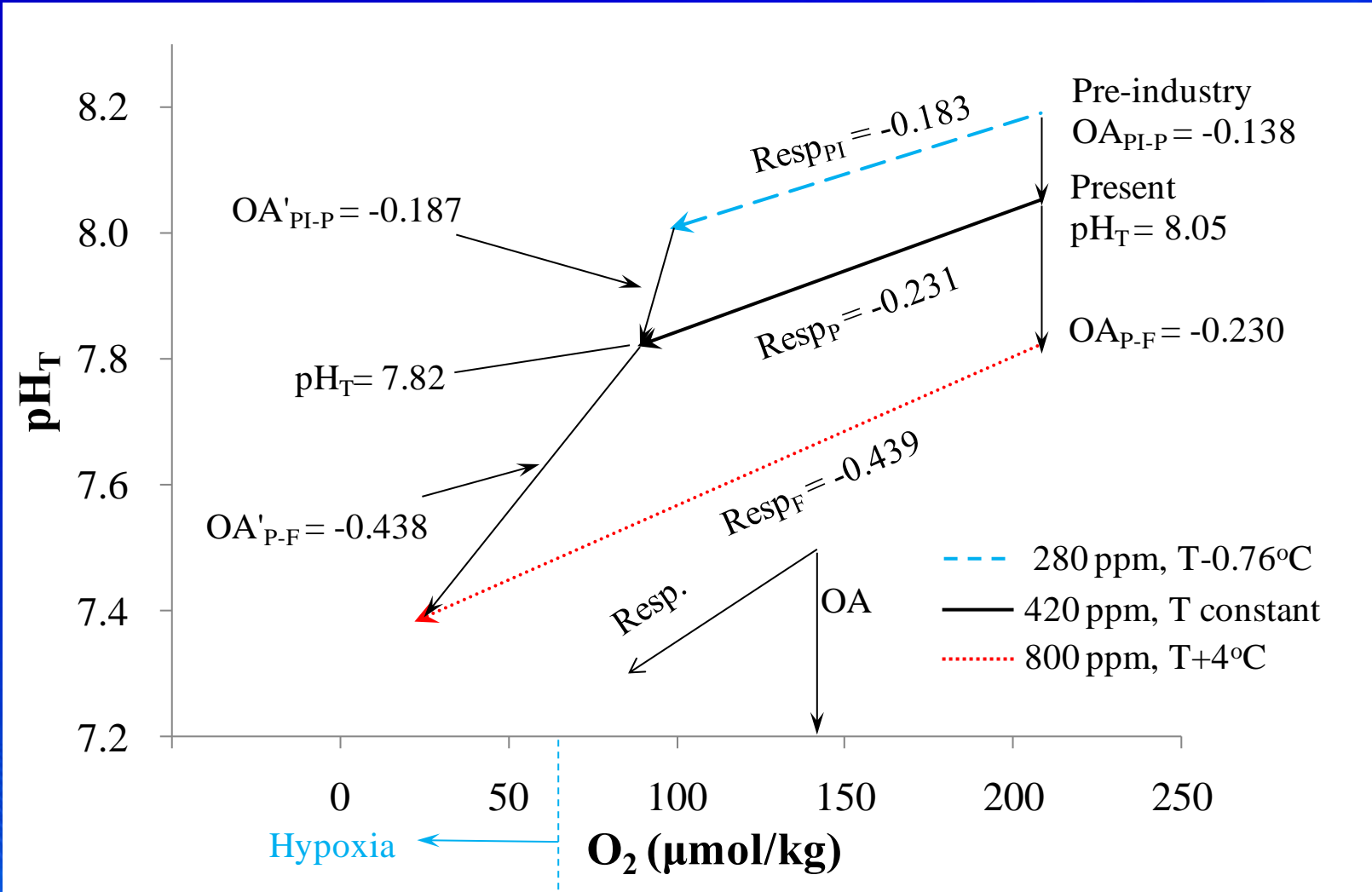
$T_{\text{year 2100}}: T+4^\circ\text{C}$

Model 1 vs 2: Higher temperature buffers the effect of ocean acidification



The values beside the "/" come from the model 1 (left) and model 2 (right)

Considering the effect of temperature on respiration



Because the effect of temperature on CR, pH will drop 0.21 unit ($\text{Resp}_{\text{F}} - \text{Resp}_{\text{P}}$) in the future than at present

Reminder

Subsurface enhancement of oxygen consumption and/or ocean acidification will depend on the seasonal hydrodynamics (if transport to the open ocean)

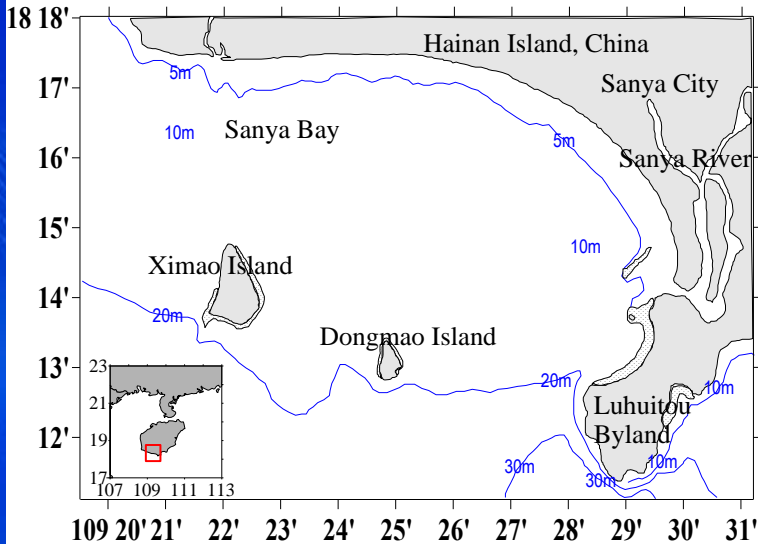
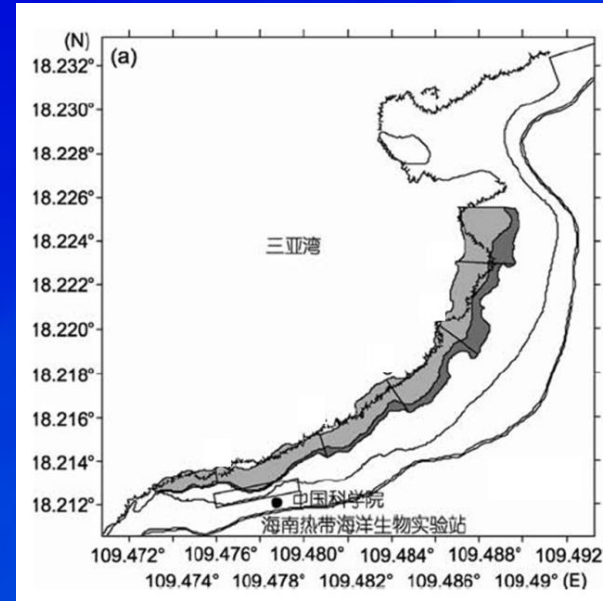
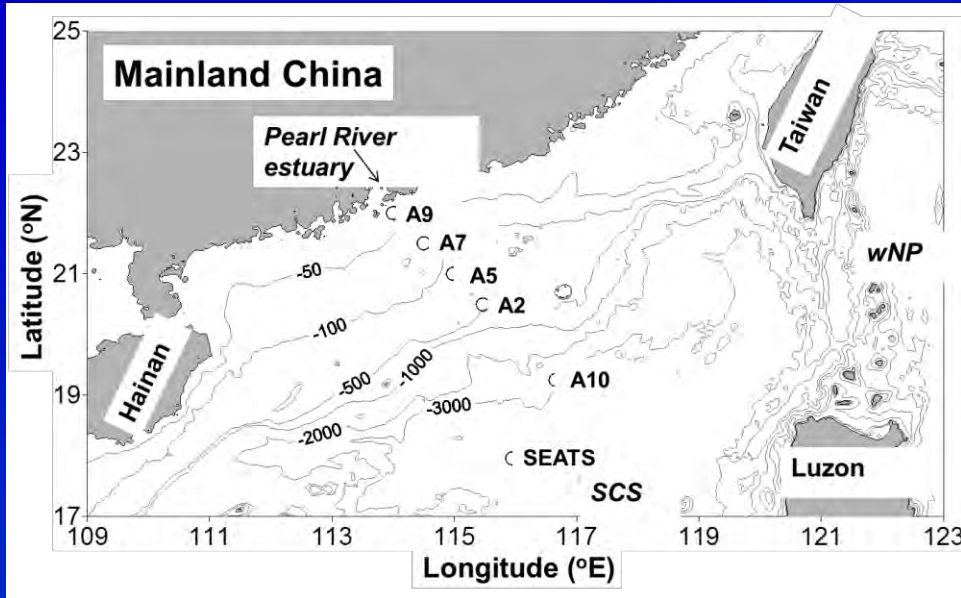


Multiple stressors: case 2

**Coastal Coral system under the
impact of groundwater, upwelling
and anthropogenic CO₂?**



Acidification in Coral Reef Environment

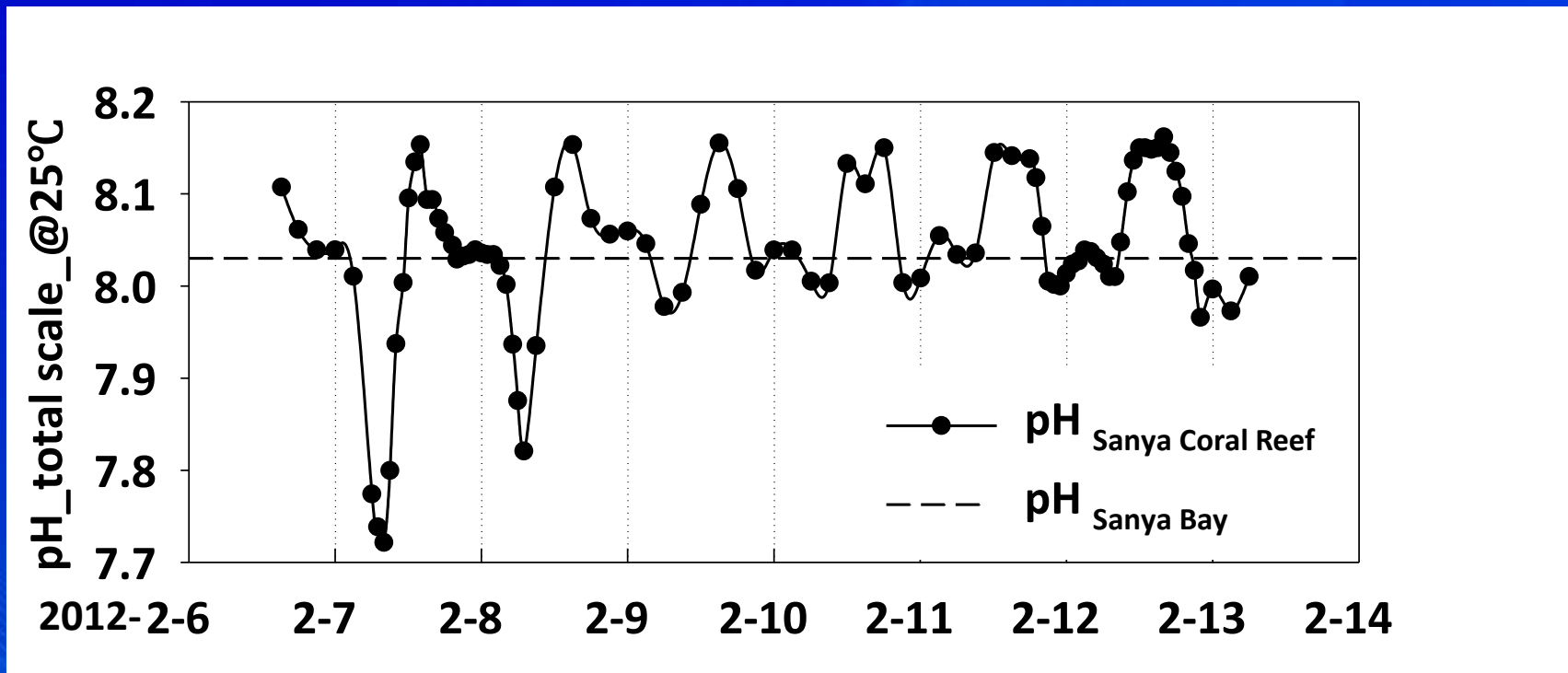


Sanya National Coral Reef Reserves:

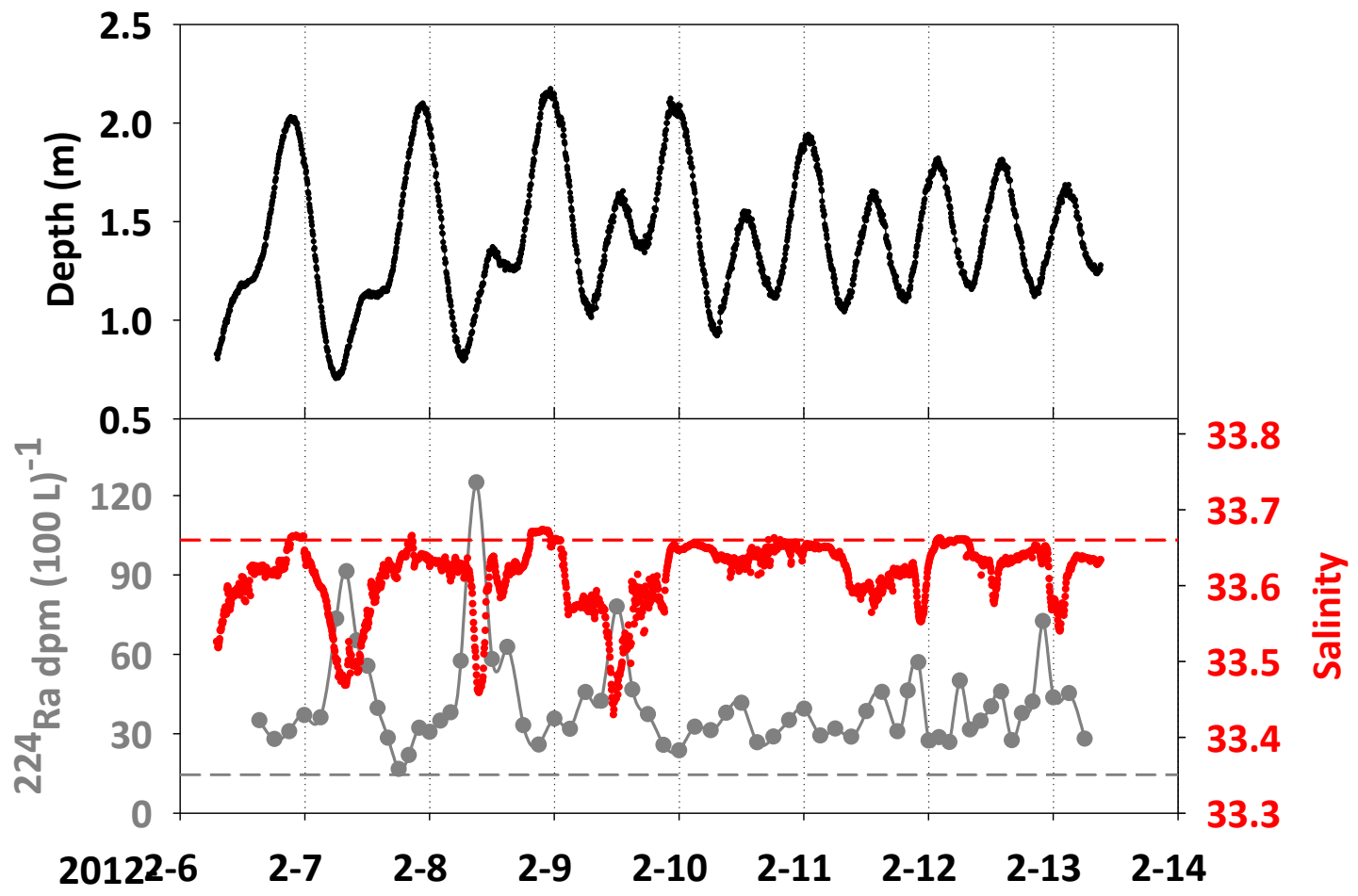
- fringing reef
- length ~ 3 km and width ~ 250~500 m

Acidification in Hainan Coral Reef

Low pH in Sanya fringing reef system



Tidal-driven Submarine Groundwater Discharge (SGD) in Sanya fringing reef



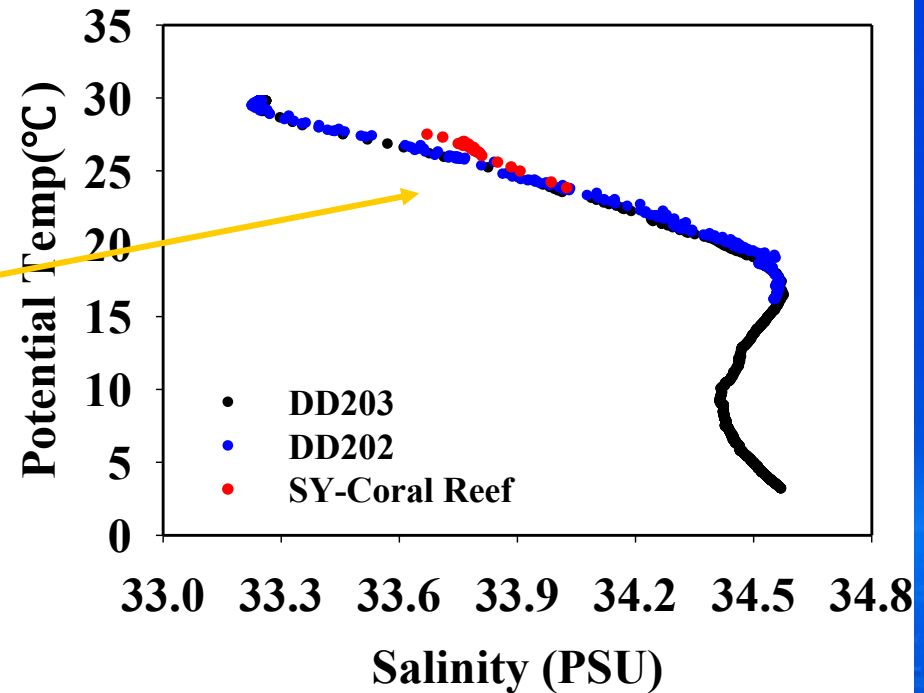
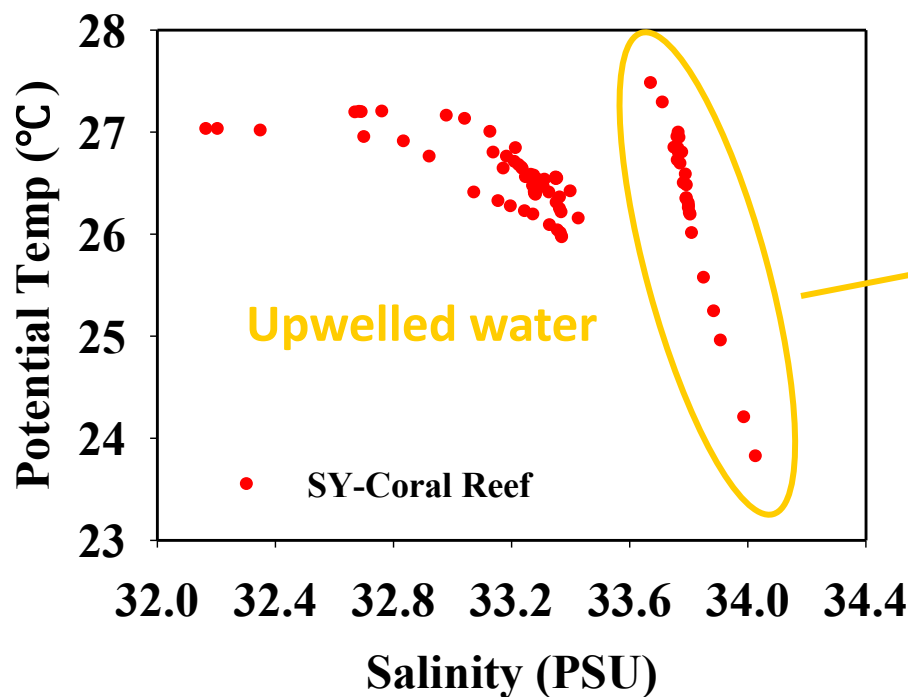
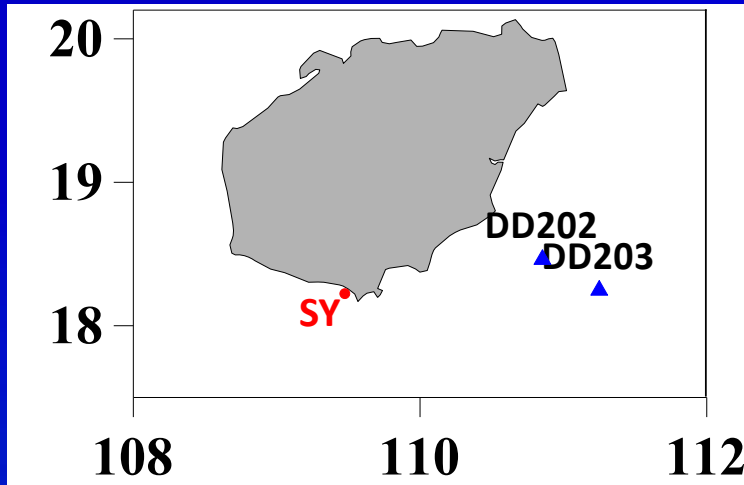
Low salinity and high ^{224}Ra during spring tide indicates the intrusion of SGD

^{224}Ra Sanya Bay
Salinity Sanya Bay

Upwelling in Sanya Coral Reef

SY station is located at the edge of coral reef ecosystem.

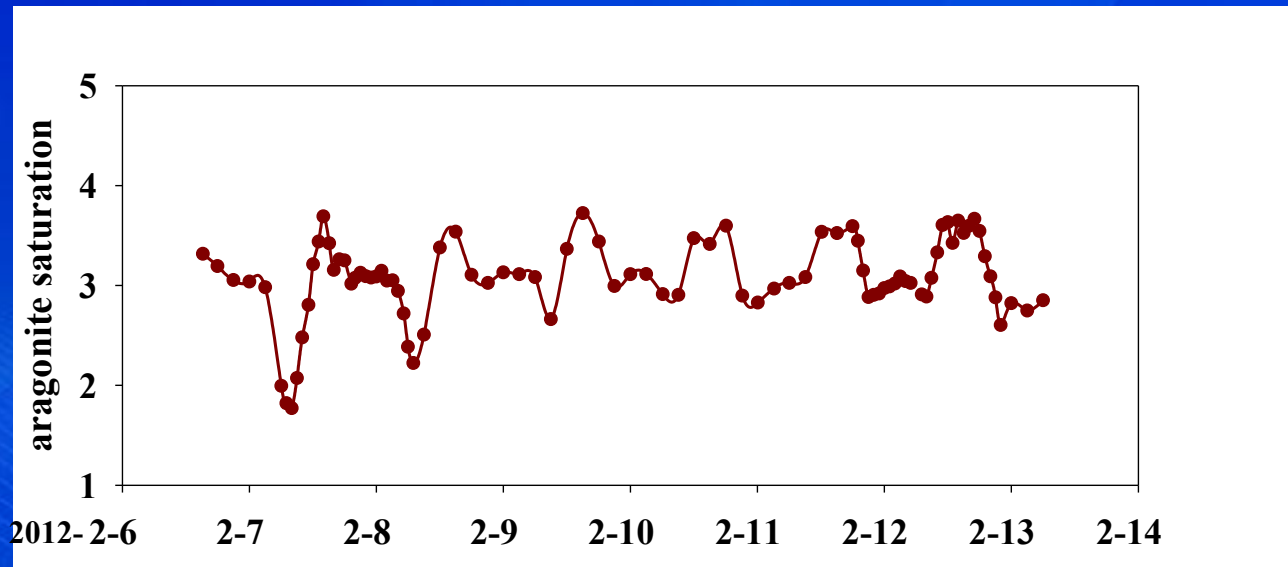
Average depth of SY station: 14m



Contribution of SGD to acidification

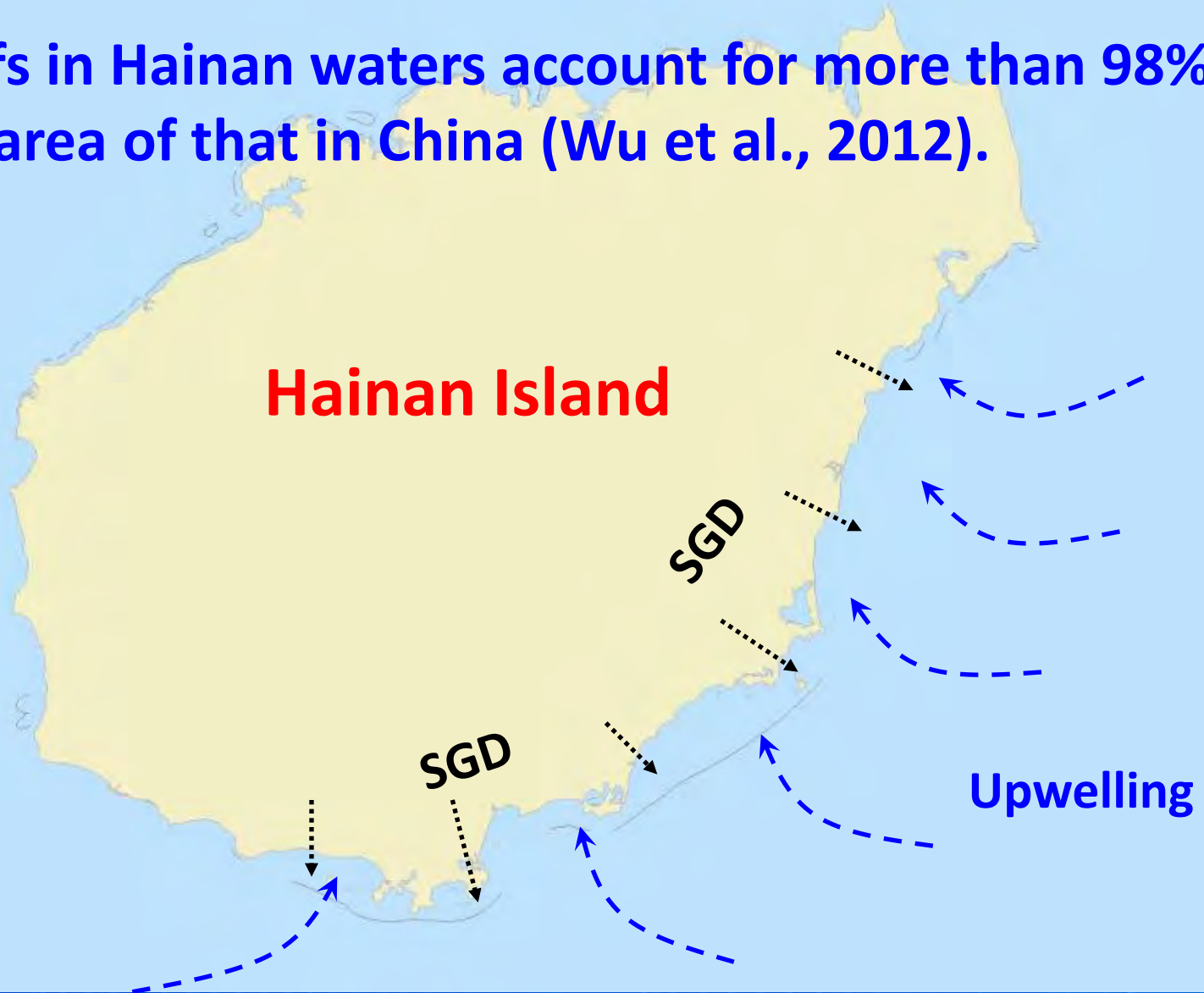
Large diurnal changes of pH and $p\text{CO}_2$ in spring tide when SGD was significant: up to 70% of the changes.

The aragonite saturation decreased to 1.77 during low tide of spring tide, a potential threat to coral reef systems.

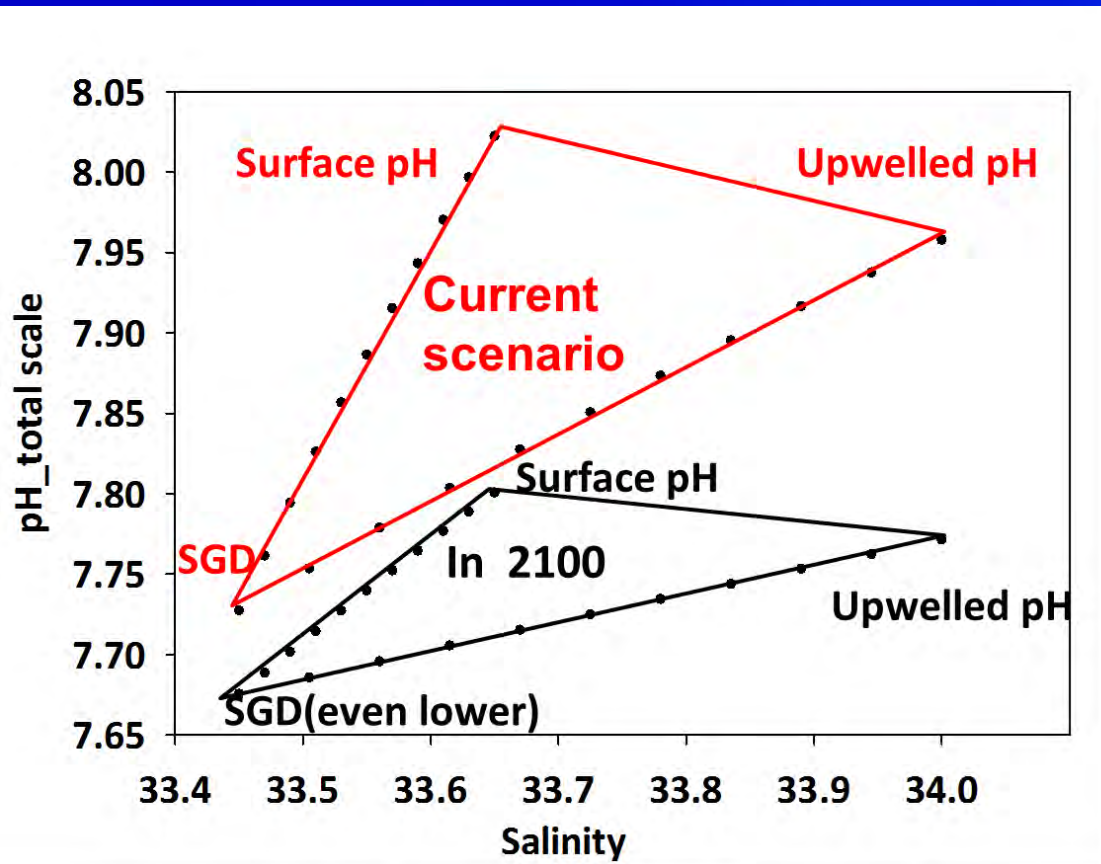


Threats to coral reef around Hainan

Coral reefs in Hainan waters account for more than 98% of the total area of that in China (Wu et al., 2012).



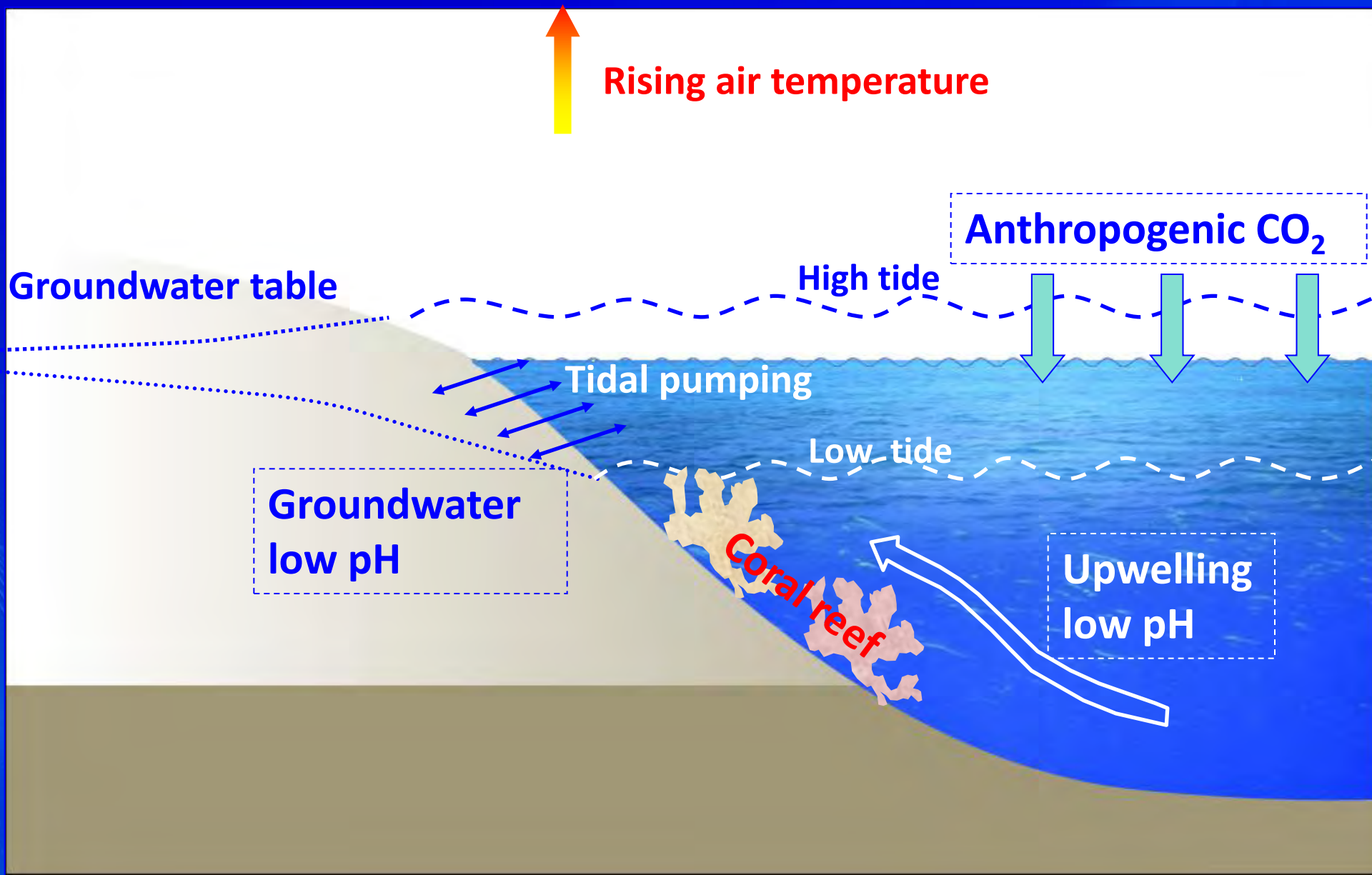
Current and Future Scenarios



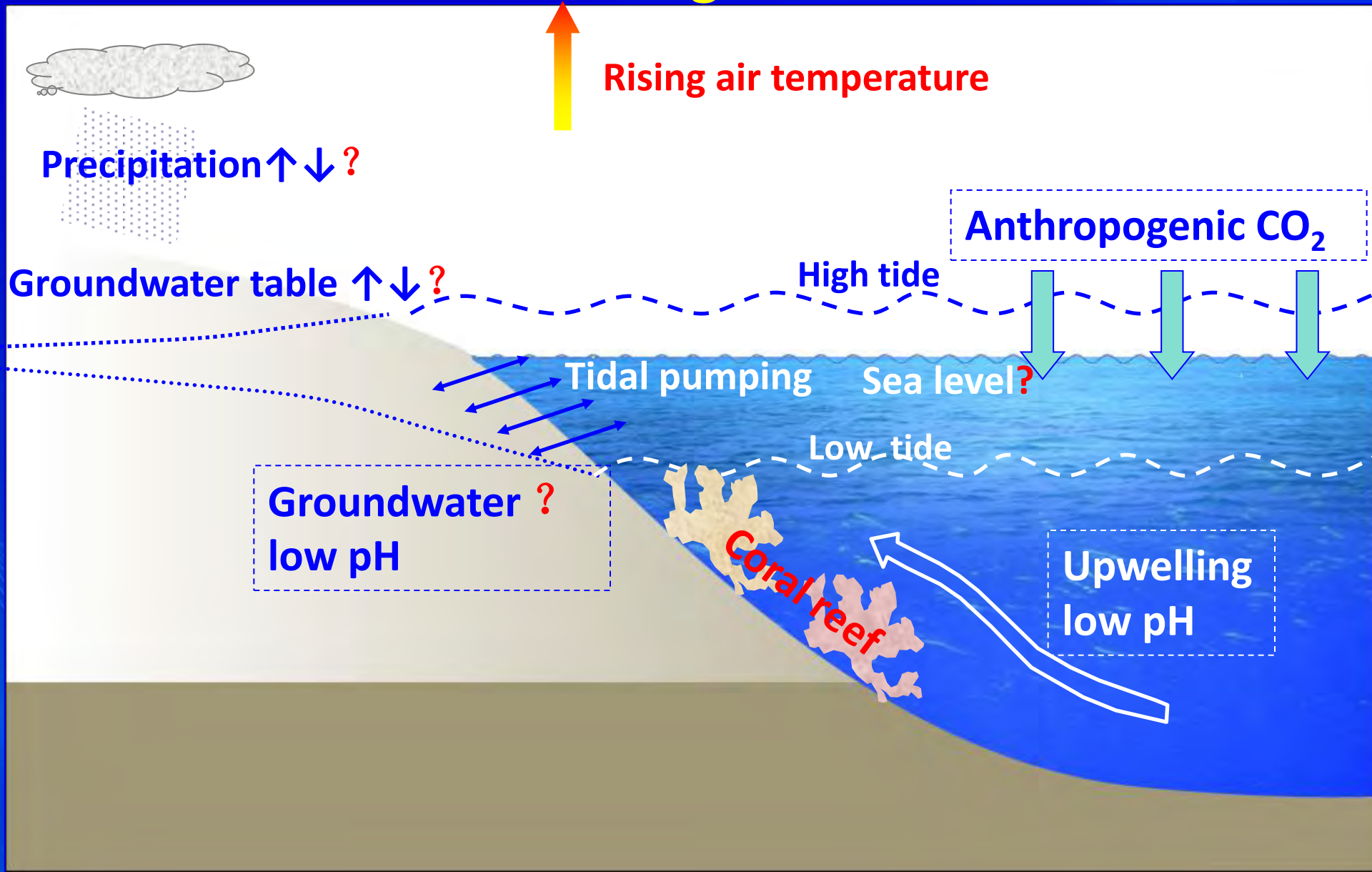
In 2100

- air $p\text{CO}_2 = 750 \mu\text{atm}$
- $\text{pH} = 7.6 \sim 7.8$
- CaCO_3 saturation down to $1.62 \sim 2.06$.

Multiple stresses to coastal coral reefs



Multiple stresses to coastal coral reefs: future changes?



Gordian Knot?



from Boyd, 2013, Nature Climate Change, 3, 530-533

Summary

- Coastal ecosystem are clearly under multiple stressors
 - Hypoxic zones
 - Coral reef systems
- The ecosystem responses/feedbacks to multiple stressors are complex.



Concluding remarks

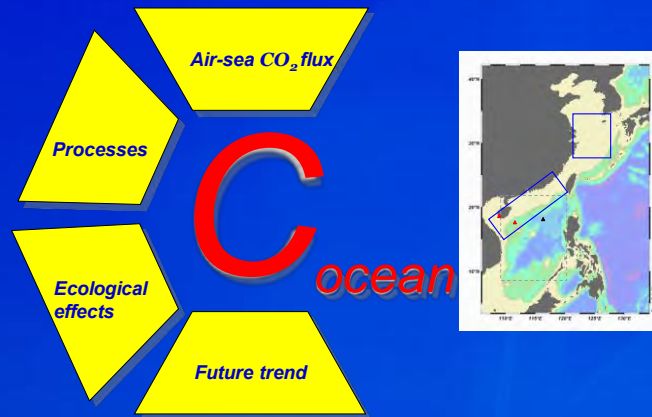
- Penetration of CO₂ in the coastal ocean has emerged
- The geochemistry of acidification is based on very well defined knowledge – yet coastal ocean more complex
- Coastal ecosystems under multiple forcings: temp rising + O₂ decline+ acidification within a similar time frame
- Need consider the hydrodynamics: e.g., Upwelling/SGD
- OA observation system & multidisciplinary researches essential

Thank you for your audience!

CHOICE-C: Carbon Cycling in China Seas - budget, controls and ocean acidification

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<http://973oceancarbon.xmu.edu.cn>

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